

Greenhouse Warming and Ocean Acidification in the Past: Lessons for the Future

James C. Zachos

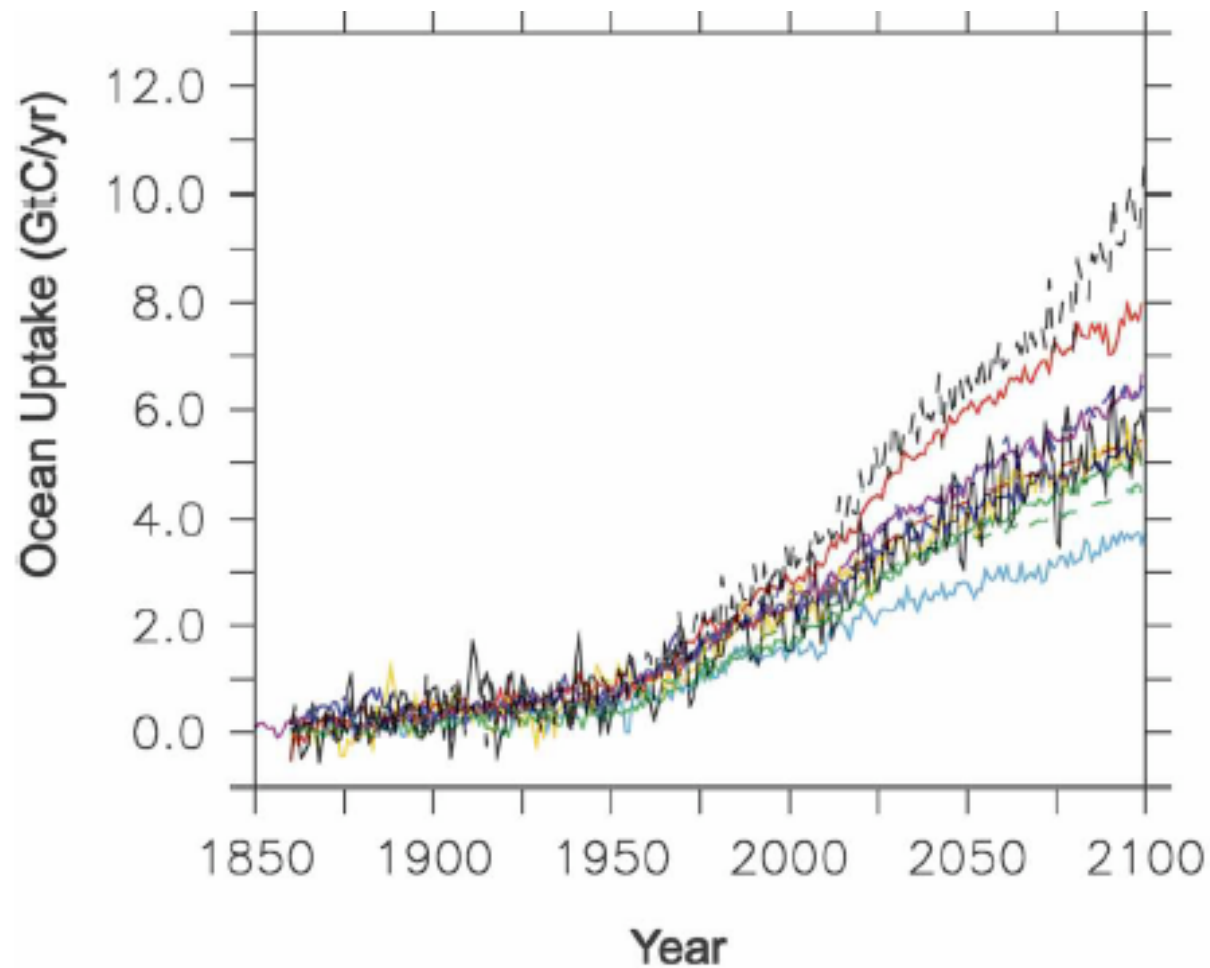
Earth & Planetary Sciences

University of California, Santa Cruz, CA



Climate–Carbon Cycle Feedback Analysis: Results from the C⁴MIP Model Intercomparison

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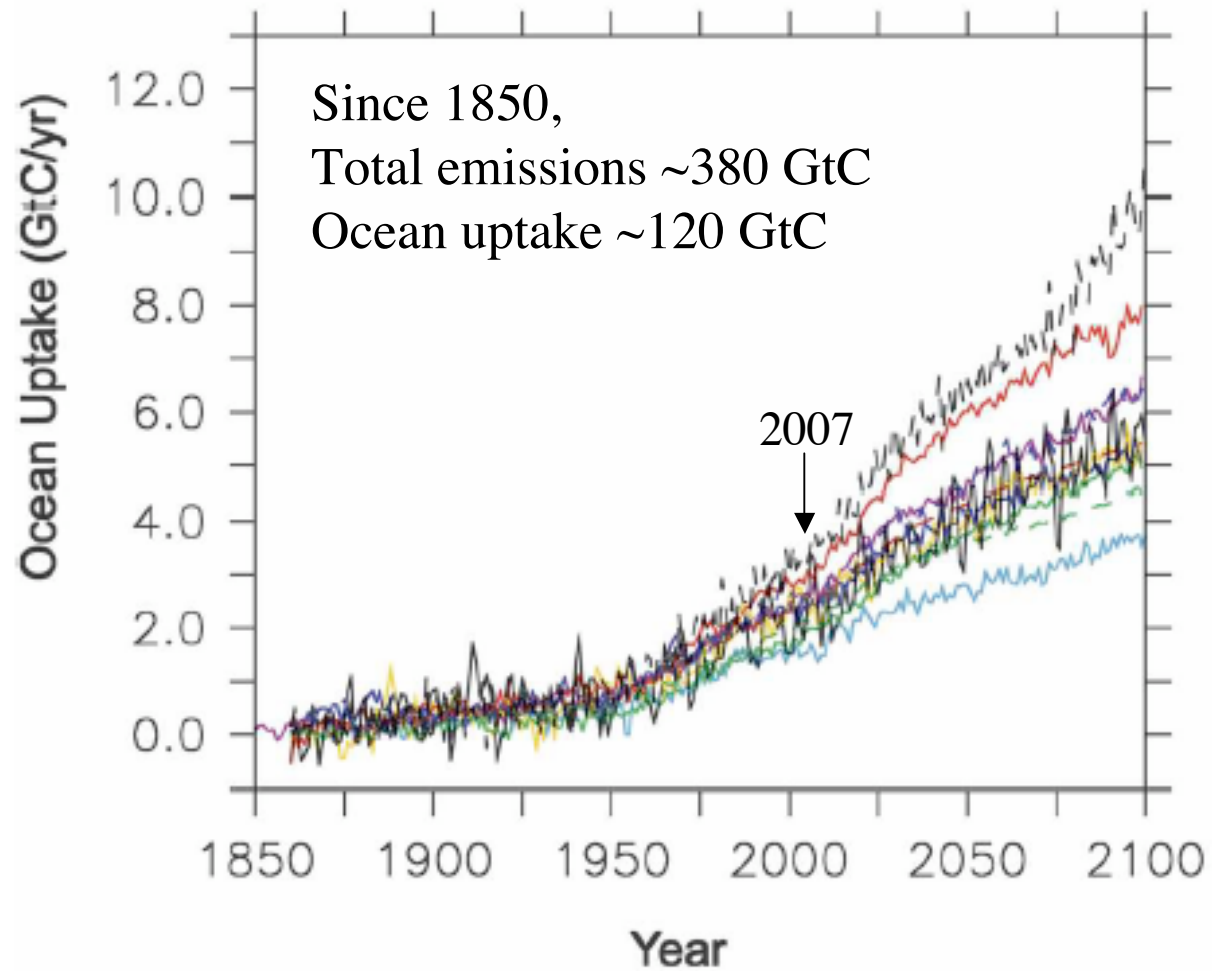


Short- and Long-Term Changes in $p\text{CO}_2$ & Global Warming

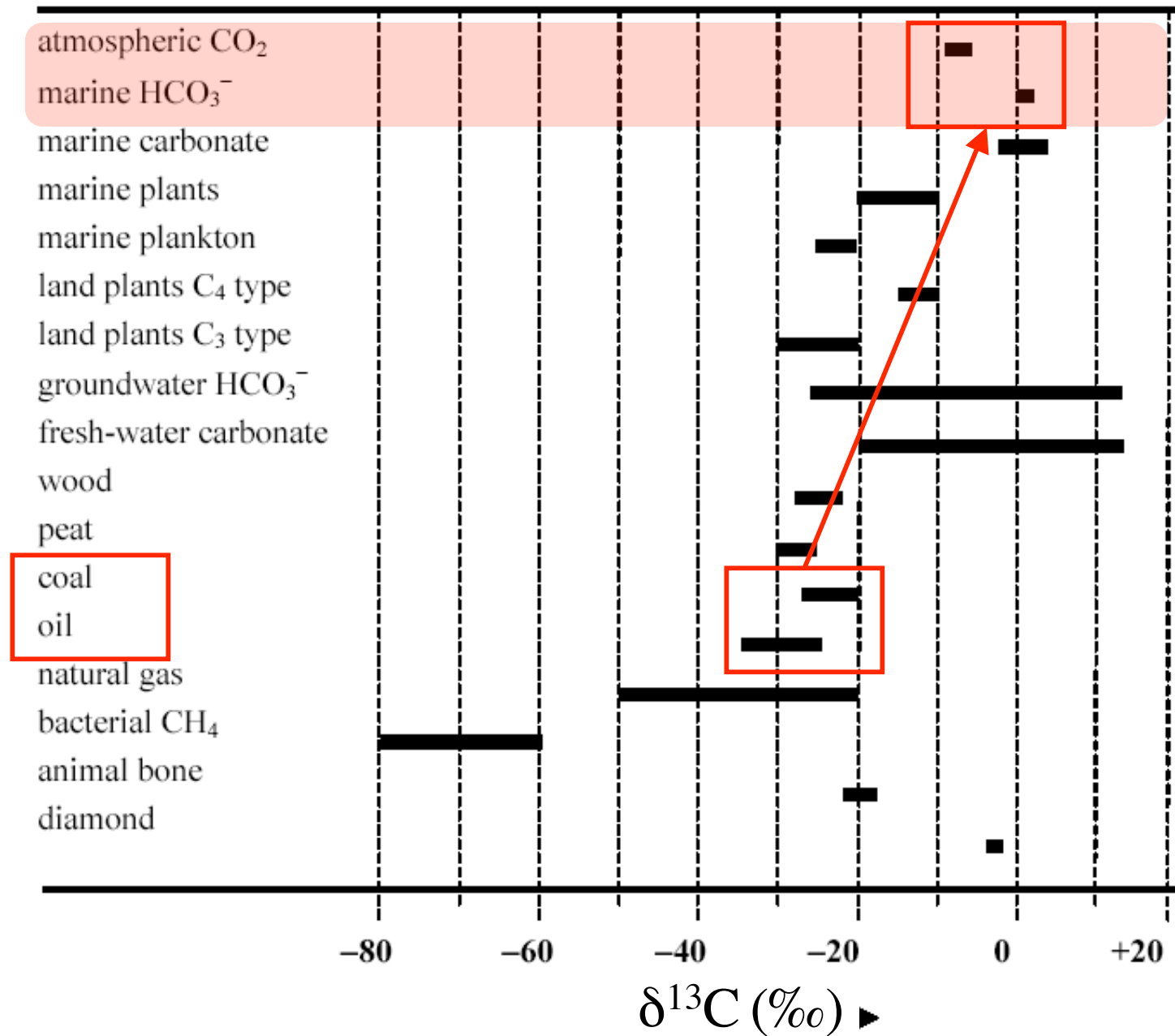
- Are there **positive** (amplifying) feedbacks that might accelerate the rate of rise in $p\text{CO}_2$?
 - ✓ *Stratification/ocean mixing ($\sim 10^2\text{y}$)*
 - ✓ *Methane hydrates ($\sim 10^3\text{y}$)*
- How fast will **negative** (damping) feedbacks sequester carbon & restore steady state?



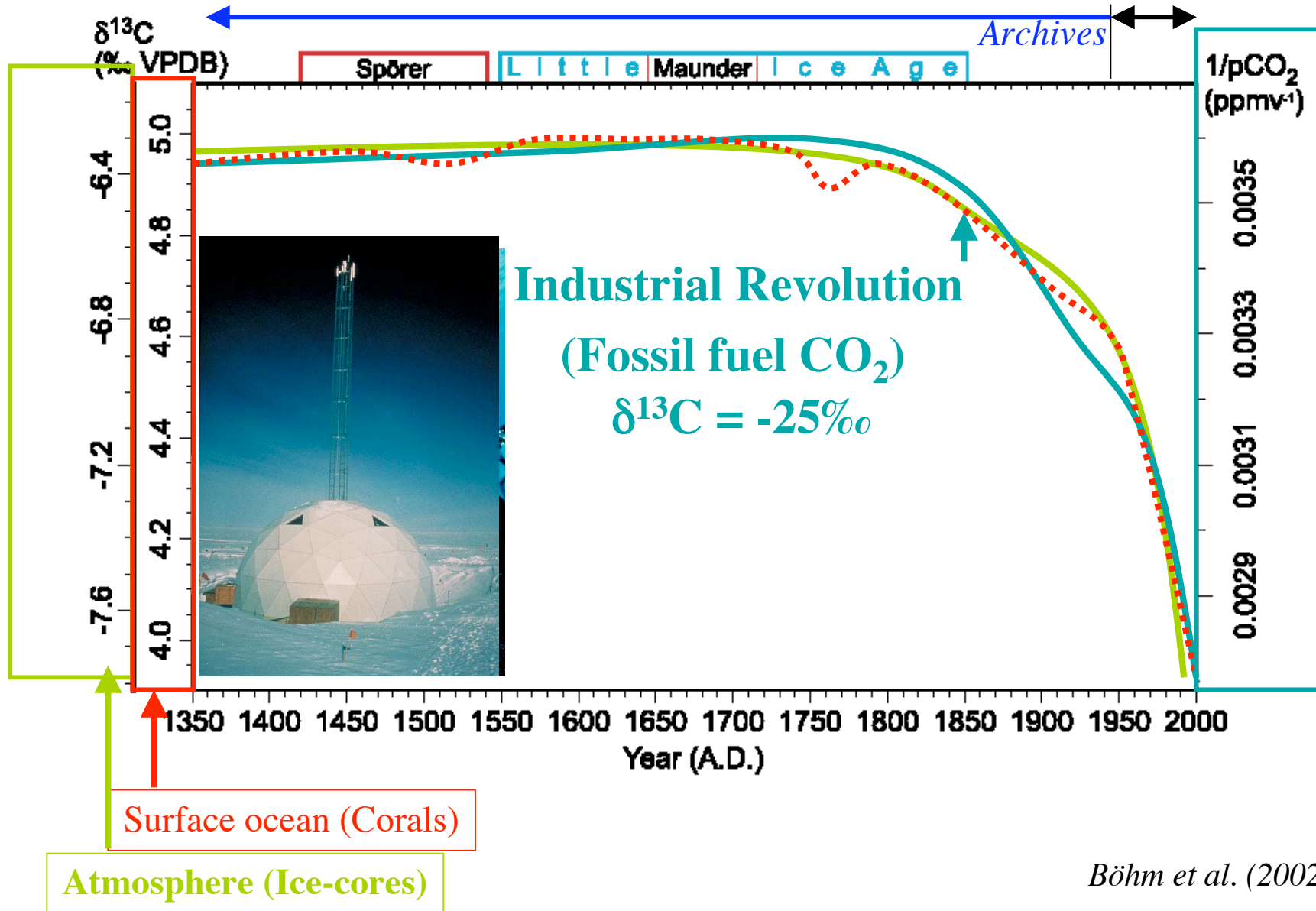
Ocean Uptake of Anthropogenic CO₂



Carbon Isotope Signature of Fossil Fuel C



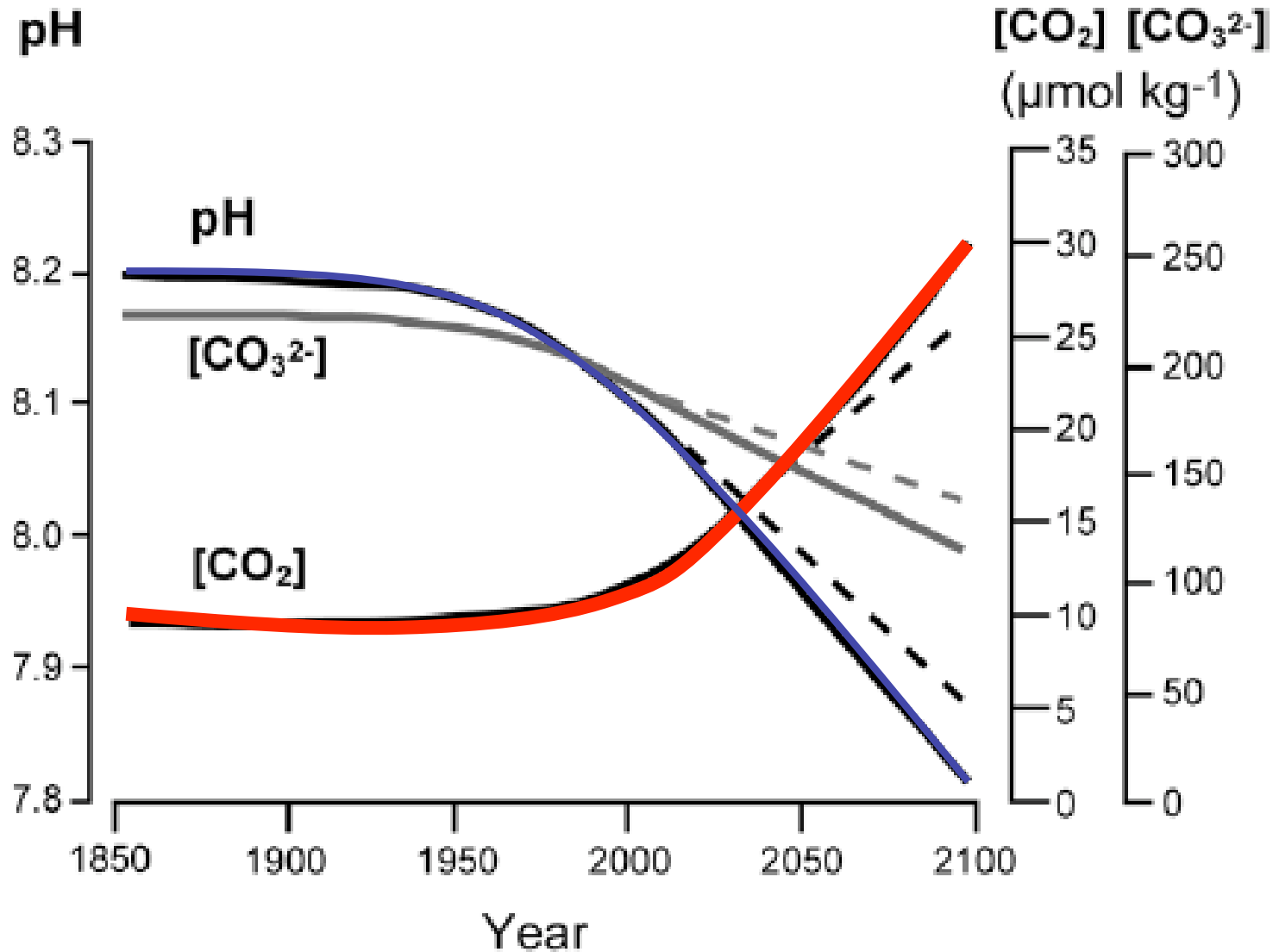
Atmosphere pCO₂/Surface Ocean δ¹³C Last 650 y



Böhm et al. (2002)

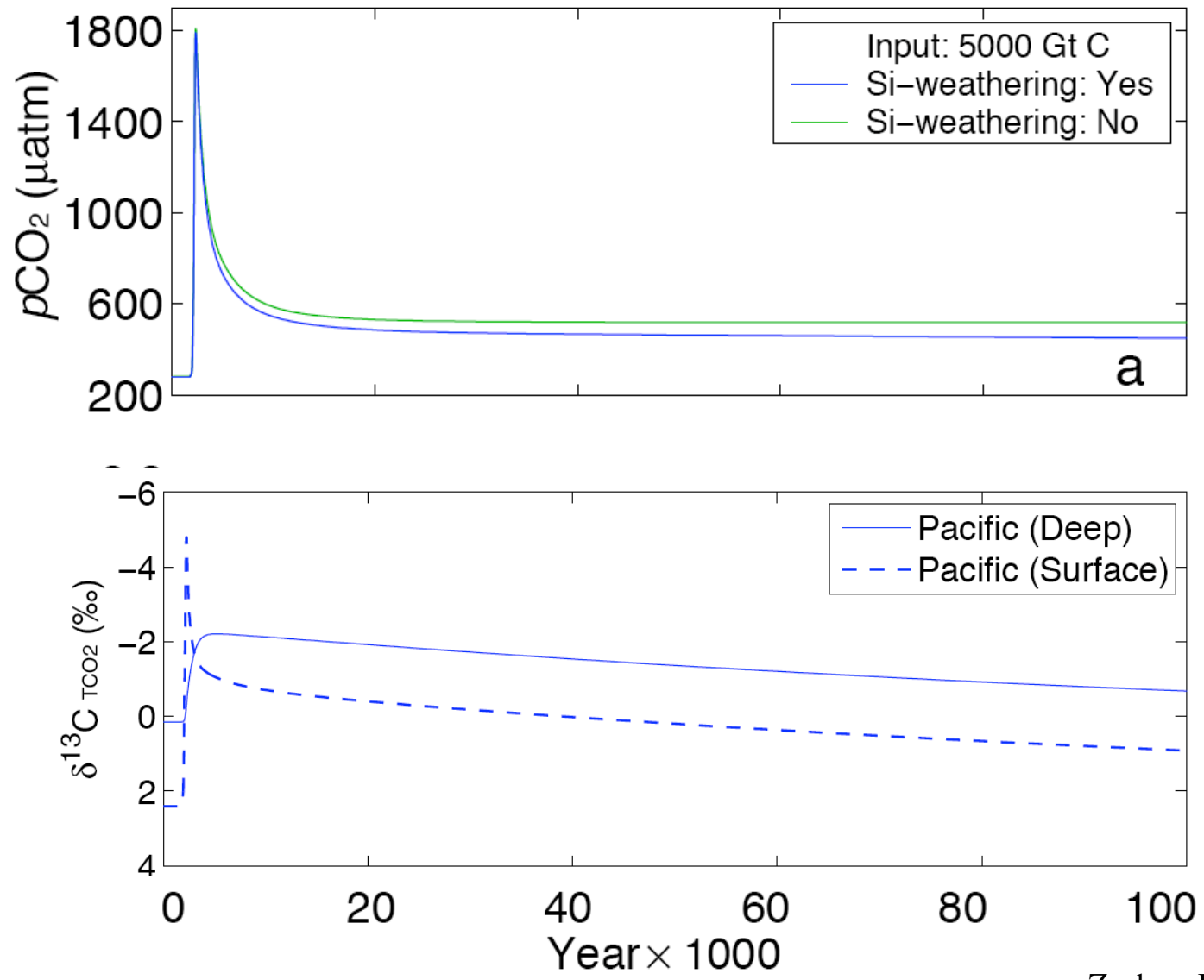
pCO₂, Ocean pH & Carbonate Ion

Next 100 years



Atmosphere pCO₂/Ocean pH & δ¹³C

Next 100,000 years



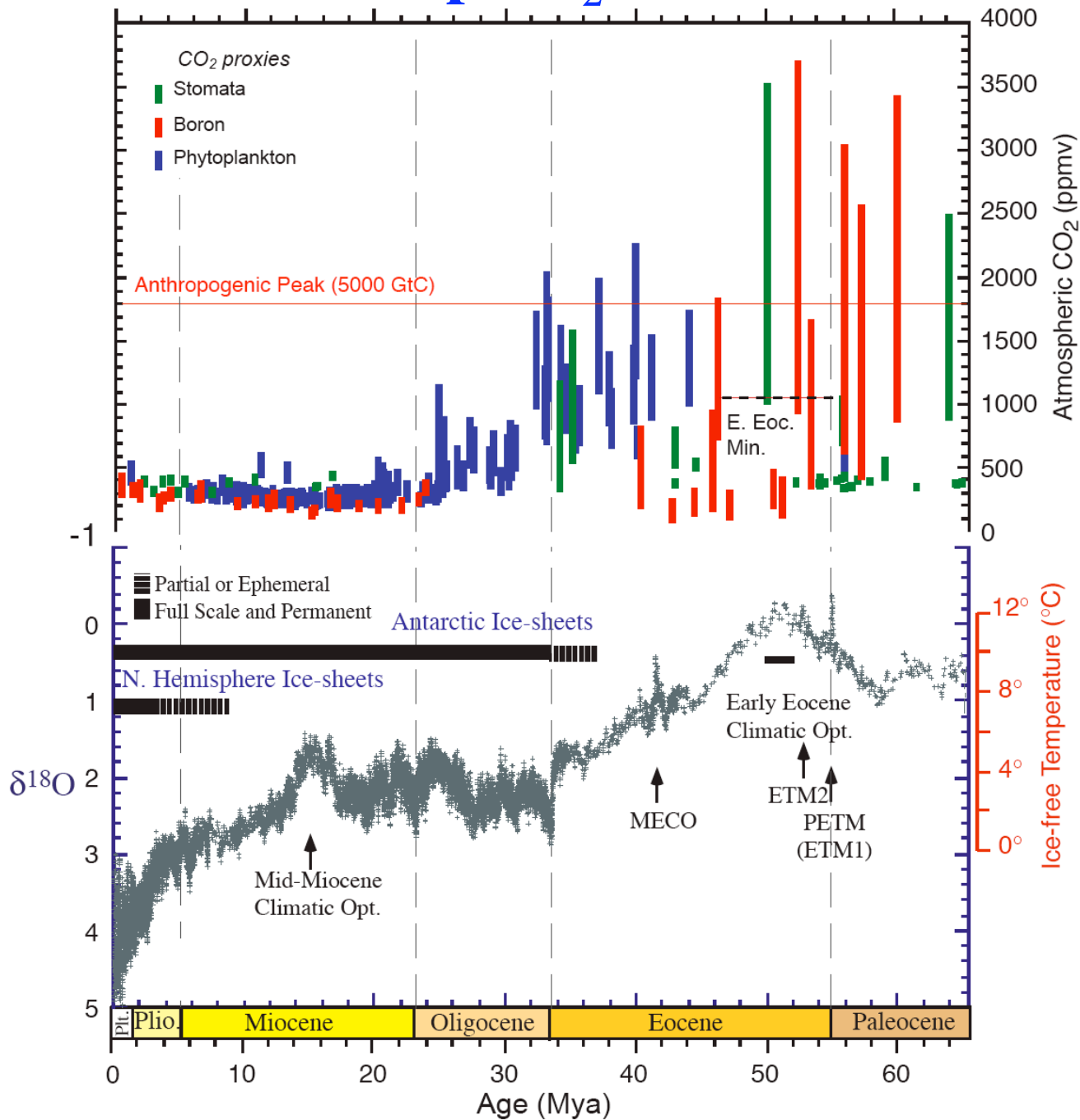
Zachos, Dickens, Zeebe,
(In press) *Nature*

How can we assess numerical predictions of short- and long-term changes in the carbon cycle?

Past greenhouse events with similar magnitude/rate of change in CO₂



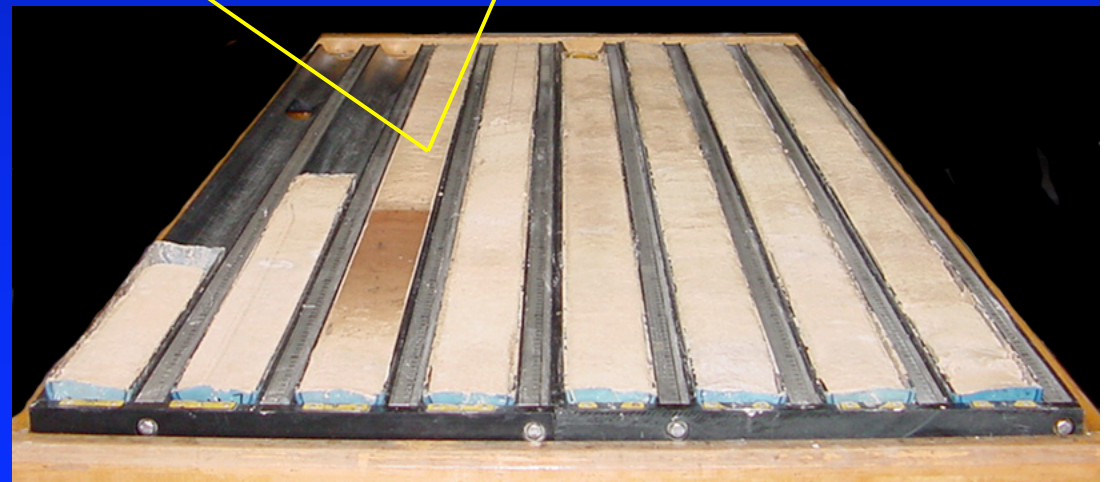
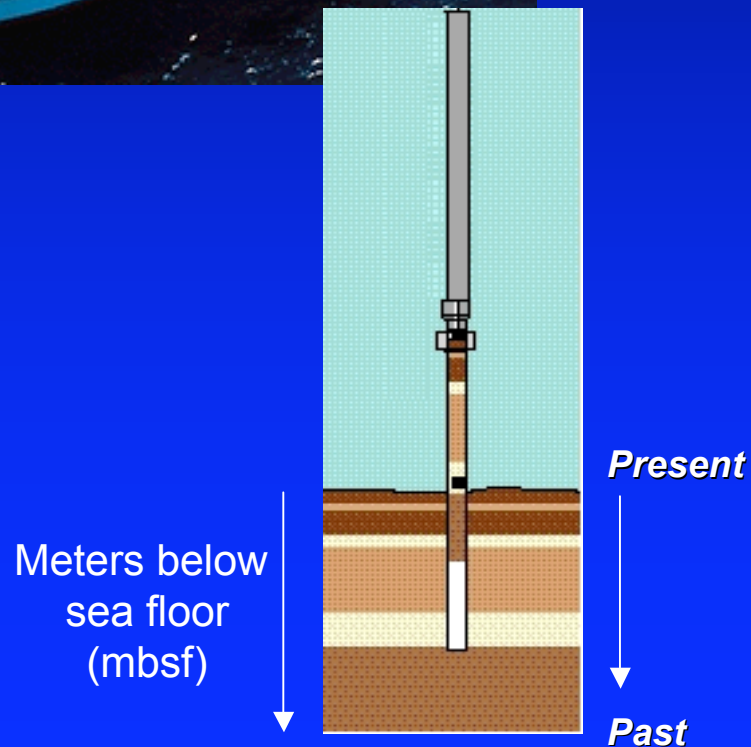
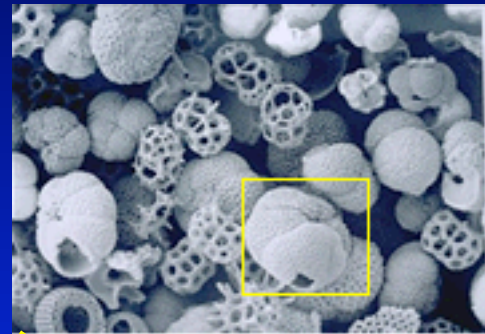
Cenozoic pCO₂ & Climate



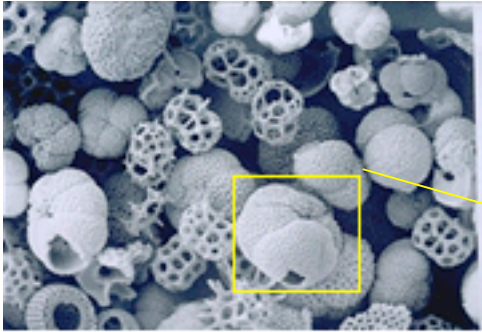
IPCC (2007);
Zachos et al.,
(In press) *Nature*

Deep Sea Sediments: Archive of Ocean History

Microfossil shell chemistry provides information on
ocean temperature & carbon chemistry



Reconstructing Ocean History with Stable Isotopes of Microfossils in Deep-sea Sediments



Foraminifera shells
~ calcite (CaCO_3)

Mass Spectrometers

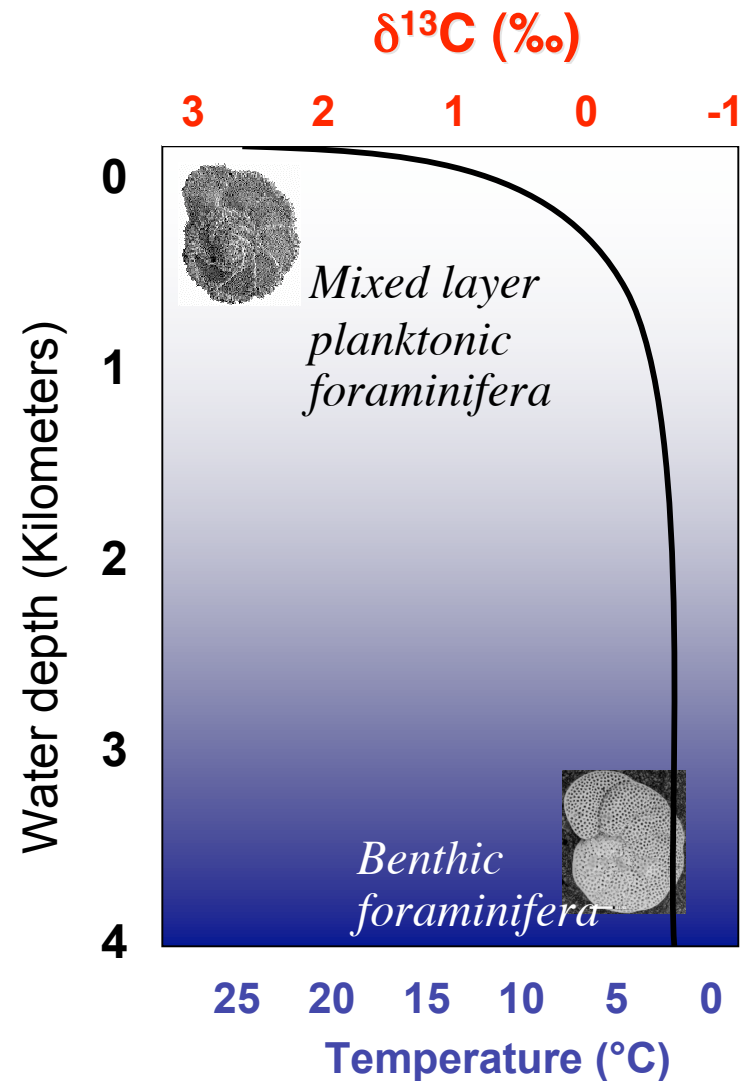


Isotopes: $^{13}\text{C}/^{12}\text{C}$
Notation: $\delta^{13}\text{C}$ (‰)

$^{18}\text{O}/^{16}\text{O}$
 $\delta^{18}\text{O}$ (‰)

Δ Ocean Temperature
As T increases, $\delta^{18}\text{O}$
decreases 1‰ ~ 4°C

Dissolved inorganic carbon (DIC) of seawater. Mean $\delta^{13}\text{C}_{\text{DIC}}$ of the ocean varies with changes in the input and output of reduced and oxidized carbon

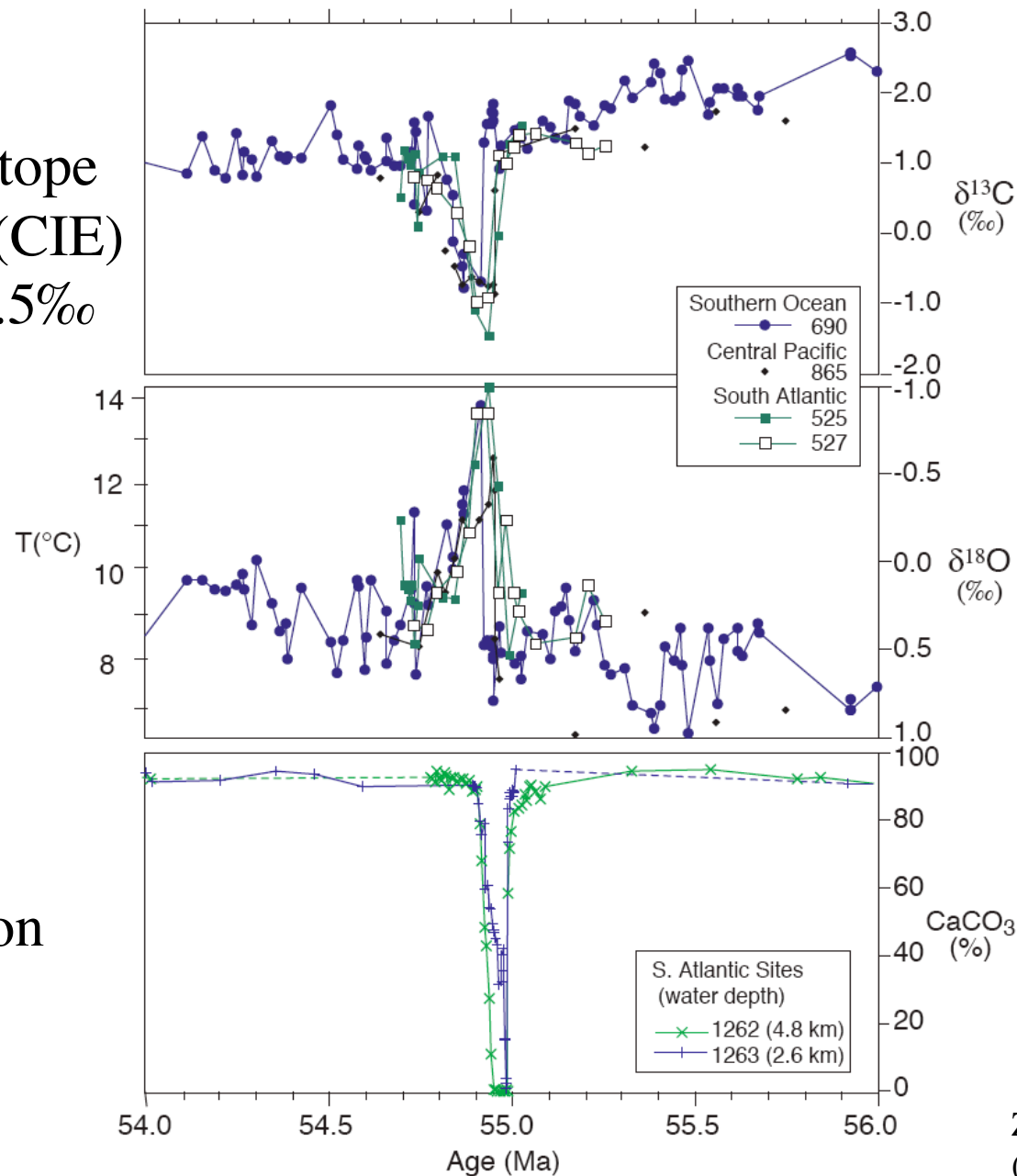


Paleocene-Eocene Thermal Maximum (PETM)

- Carbon Isotope Excursion (CIE)
- $\Delta\delta^{13}\text{C} = -2.5\text{‰}$

- Warming
- $\Delta T = 5^\circ\text{C}$

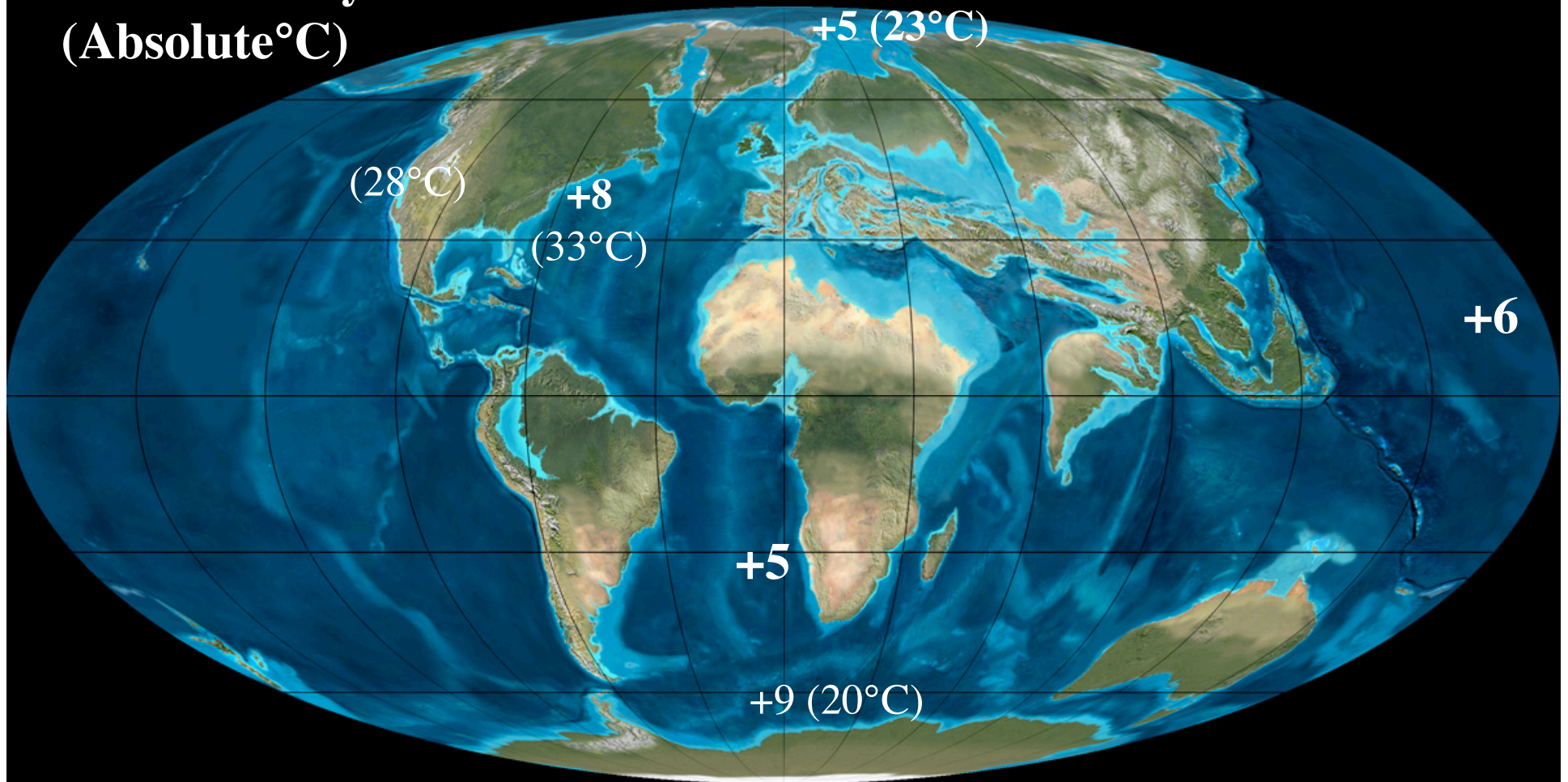
- Ocean Acidification



Zachos, Dickens, Zeebe,
(In press) *Nature*

Paleocene-Eocene Thermal Maximum (PETM) ~ 55 Mya

SST Anomaly
(Absolute °C)



Temperature anomalies estimated from $\delta^{18}\text{O}$, Mg/Ca, and TEX_{86} Proxies

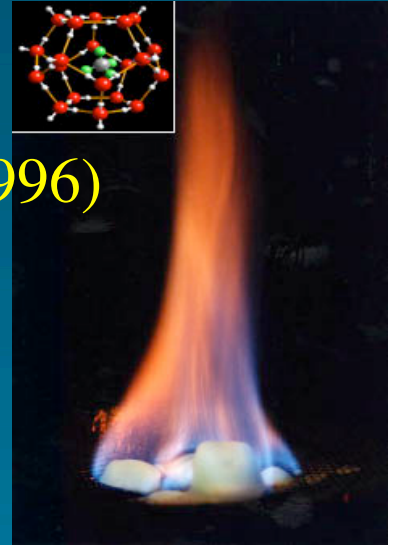
Kennett & Stott, 1991; Zachos et al., 2003; 2006; Thomas et al., 2002; Sluijs et al., 2006; John et al., submit.

Some Climatic/Environmental Consequences of the PETM

- Increased aridity in low-latitudes
- Increased precipitation in high-latitudes
- Increased seasonality in precipitation
- Increased frequency of extreme weather events & wildfires
- Changes in the diversity/abundances of fauna & flora
 - Migration, extinction (minor), reduced diversity
 - biogeographic boundaries shift poleward
- Sea level rise of 15 meters
 - 3-5 meters thermal expansion



Primary Source of Carbon?



Decomposition of Methane Hydrate - (Dickens et al., 1996)

- Bacterial, $\delta^{13}\text{C} = -60\text{‰}$.
- ~2000-10000 Pg C (modern reservoir)

Mantle Plume/Mid-ocean Ridge Volcanism - CH_4/CO_2 (Svensen et al., 2004)

- Thermal Corg decomposition, $\delta^{13}\text{C} = -7$ to -25‰
- Emission rate? 0.1-0.5 Pg C/y

Dessication & Oxidation of Corg (soils/sediments)

- Forest peats/bogs/swamps/other? - $\delta^{13}\text{C} = -20$ to -25‰
- Collectively >5000 Pg C

Positive Feedbacks?

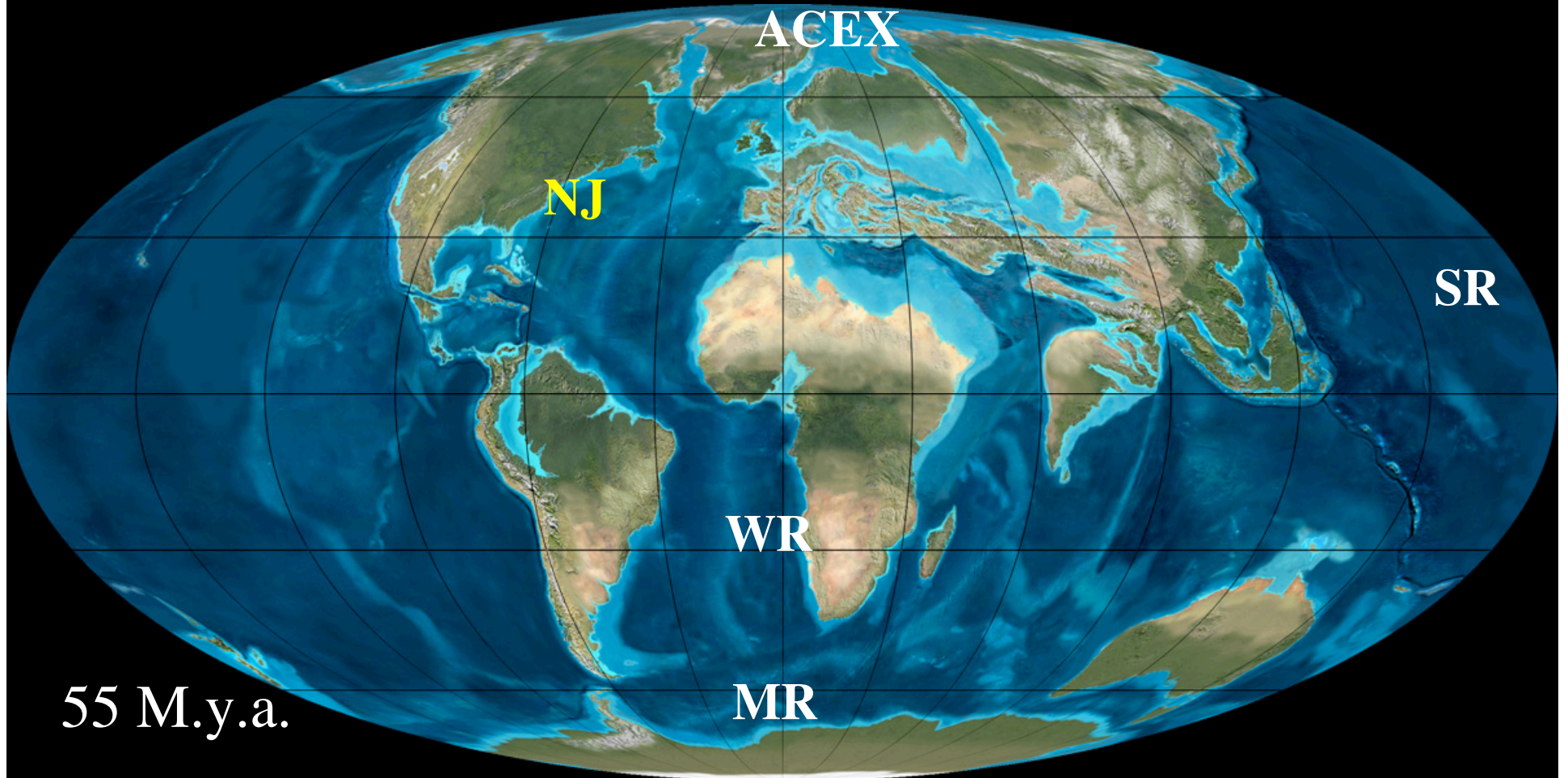
e.g., gradual warming of the ocean destabilizes methane hydrates

Source of the Massive Carbon Flux?

- Single or multiple?
- Rate of release?
 - ✓ Higher-fidelity records
 - ✓ Single shell strategy
- Mass of carbon?
 - ✓ Carbon isotope excursion (CIE)
 - ✓ Carbonate saturation changes

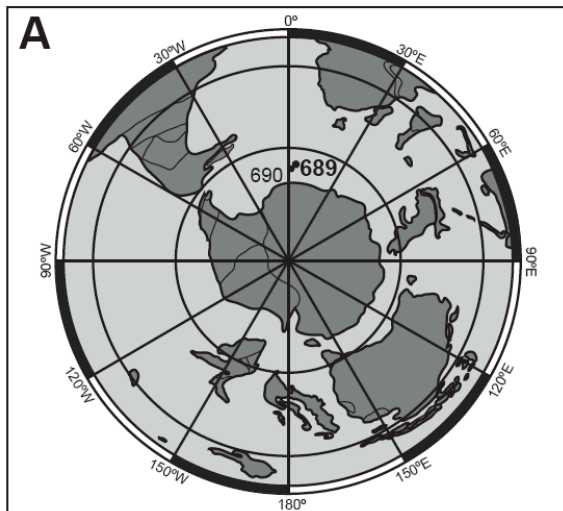


Key PE Boundary Sections

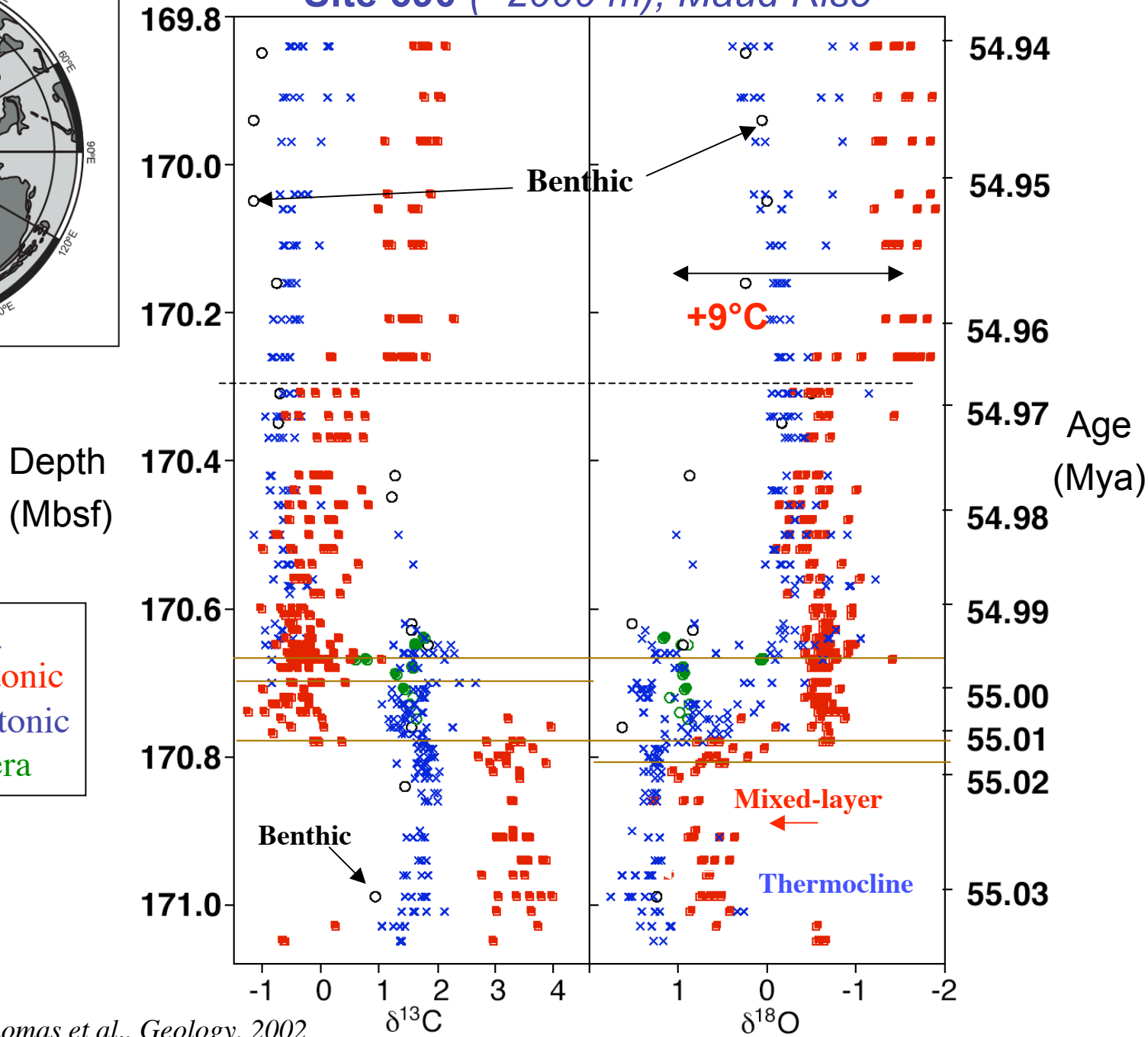


55 M.y.a.

- ACEX - Arctic coring expedition, Lomonosov Ridge
- NJ - New Jersey Margin, *Bass River* and *Wilson Lake*
- WR- Walvis Ridge (Leg 208), Sites 1262-1267
- SR - Shatsky Rise (Leg 198) Sites 1209, 1210, 1211
- MR - Maud Rise (Leg 113) Sites 689 and 690

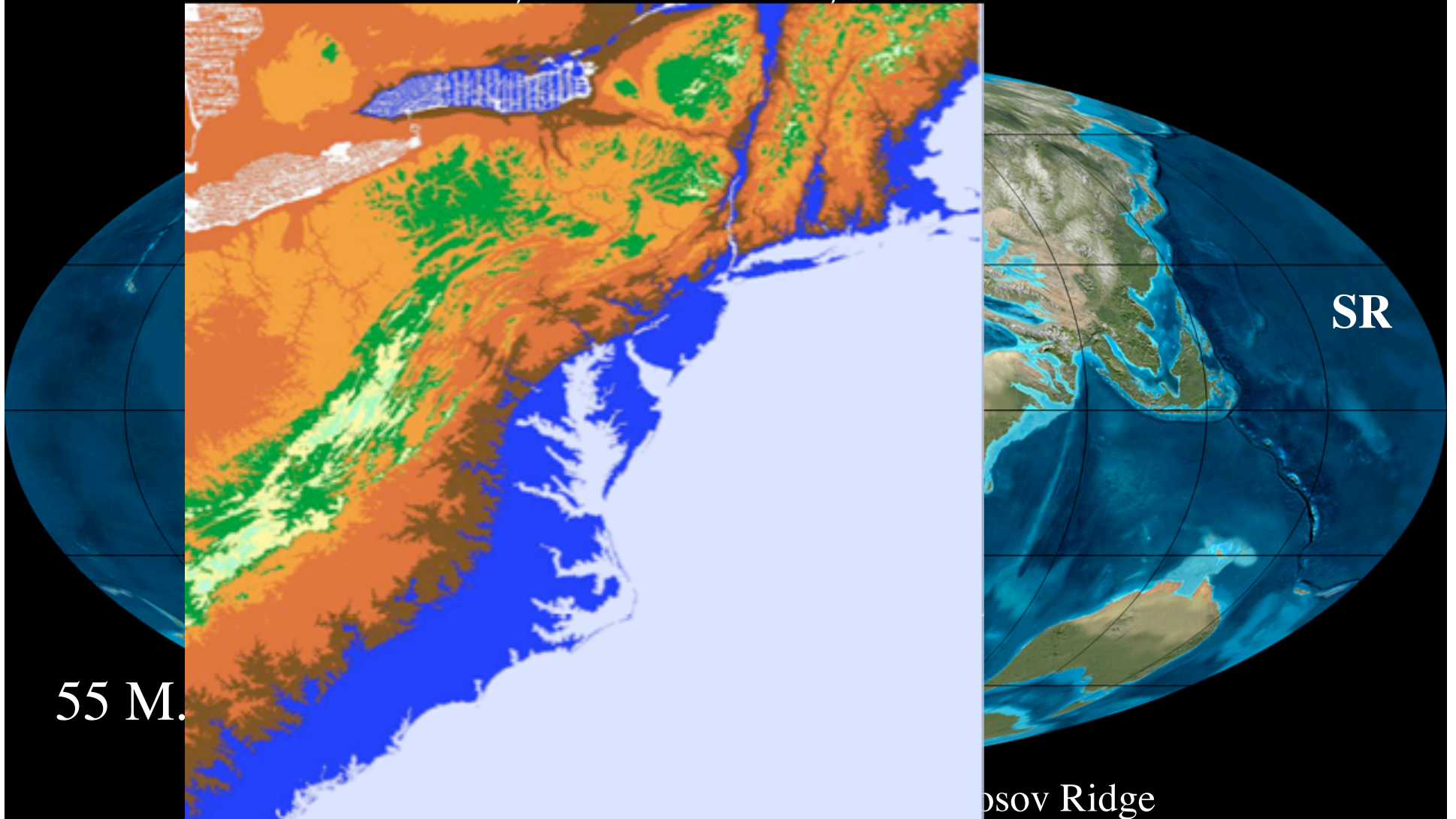


Site 690 (~2000 m), Maud Rise



Source: Thomas et al., *Geology*, 2002

Key PE Boundary Sections



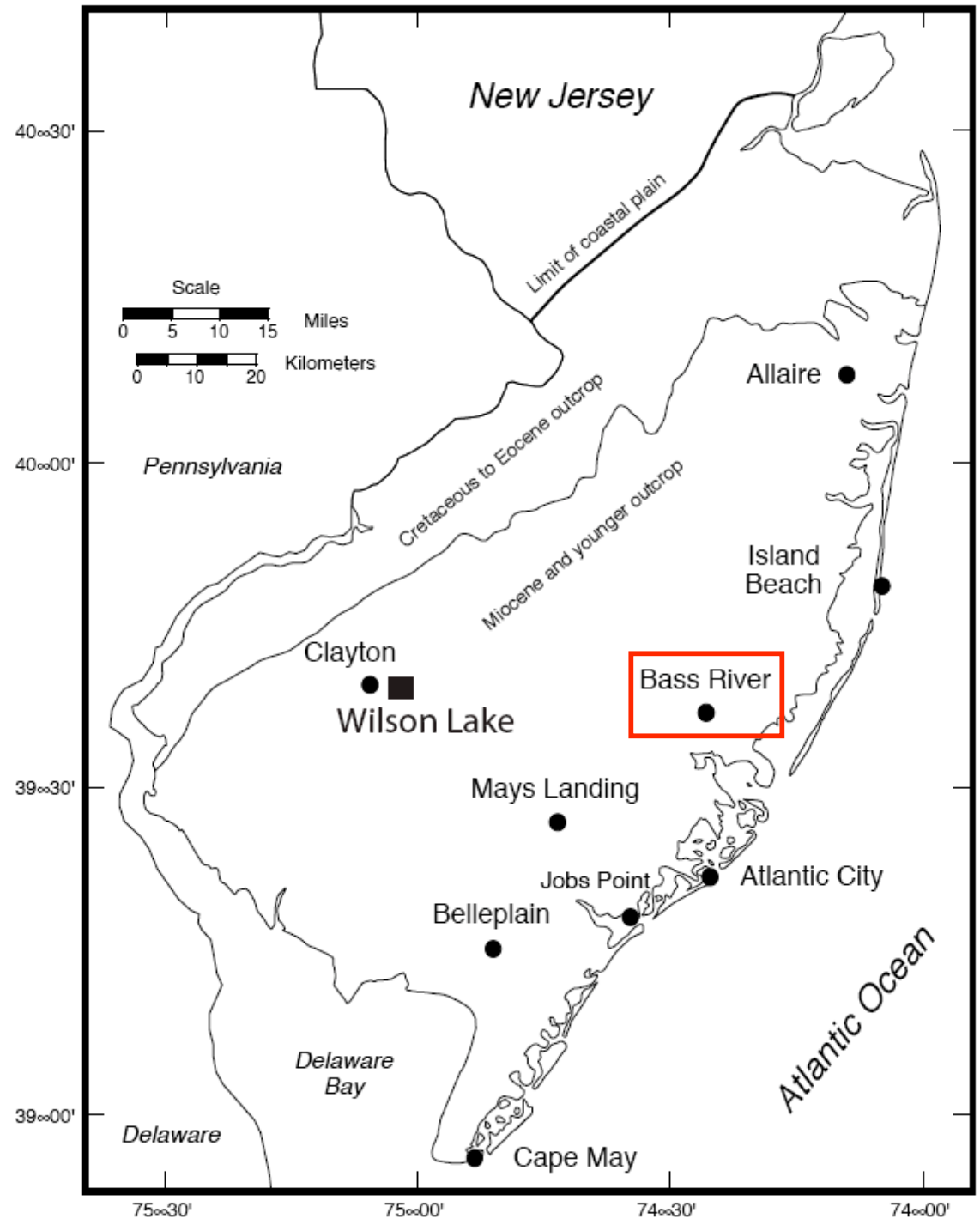
55 M.

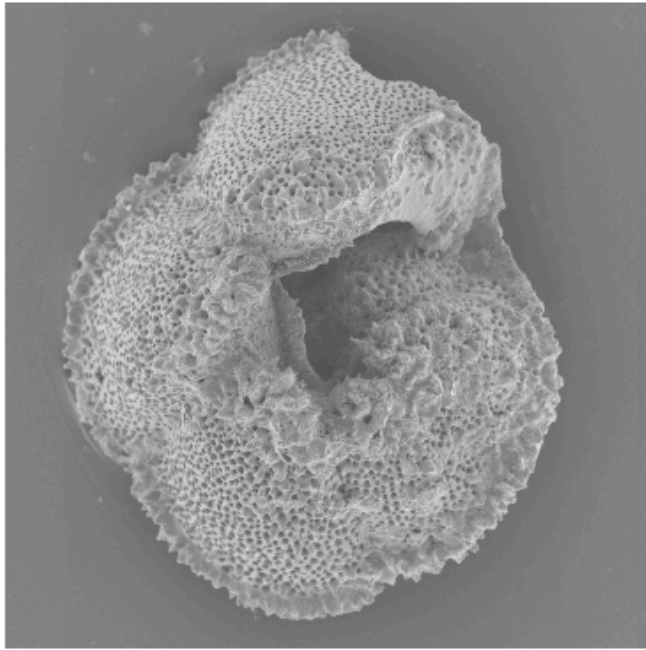
SR

- NJ - New Jersey Margin, *Bass River* and *Wilson Lake*
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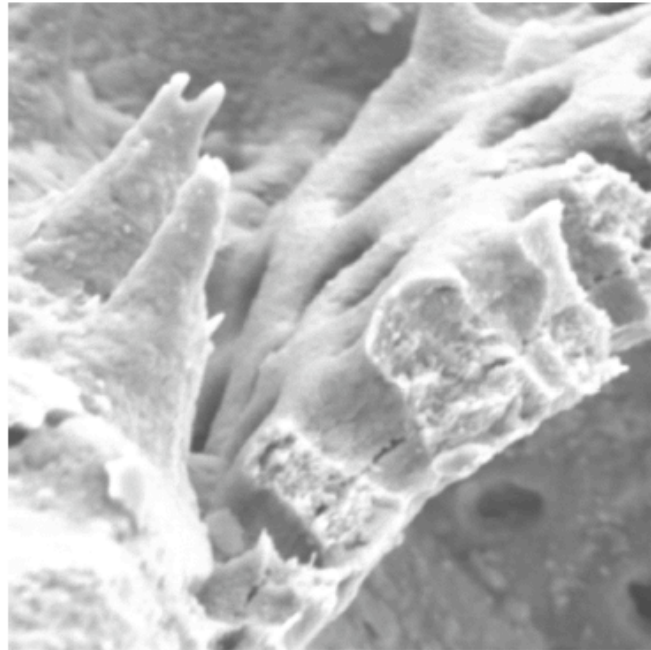
Bass River, NJ

- Mid-Shelf Environment;
unconsolidated silici-
clastic silts & clay
 - < 200 meters
 - *Kaolinite* rich
 - P-E boundary is
conformable
 - *Dinoflagellate* blooms
 - Foraminifera - scarce,
but well preserved
- High sedimentation rates
 - 5-10x deep sea



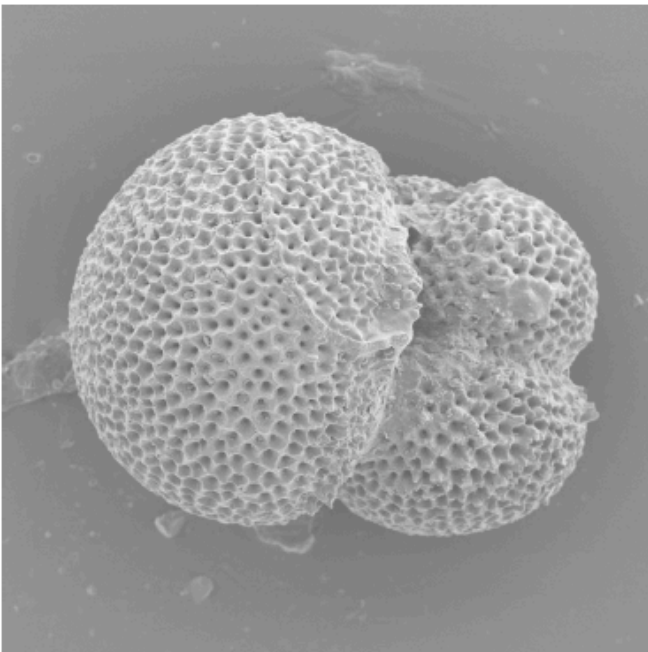


A.



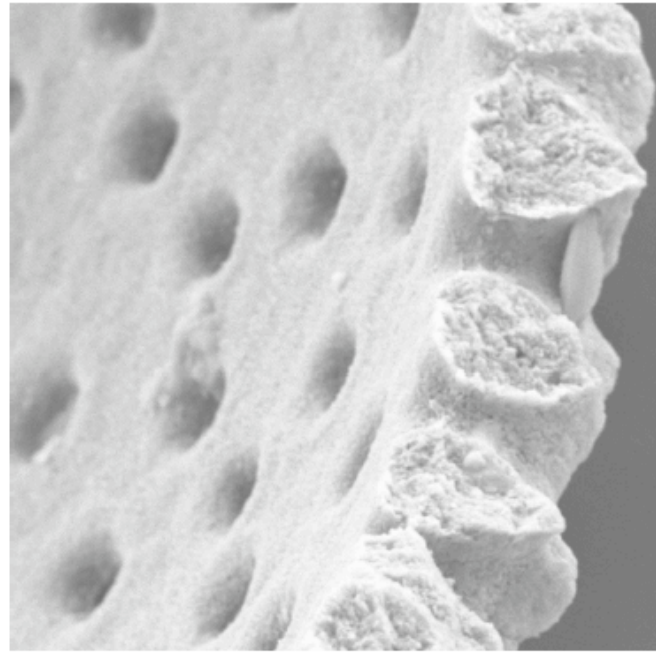
B.

6 μ m 4100X



C.

80 μ m 280X

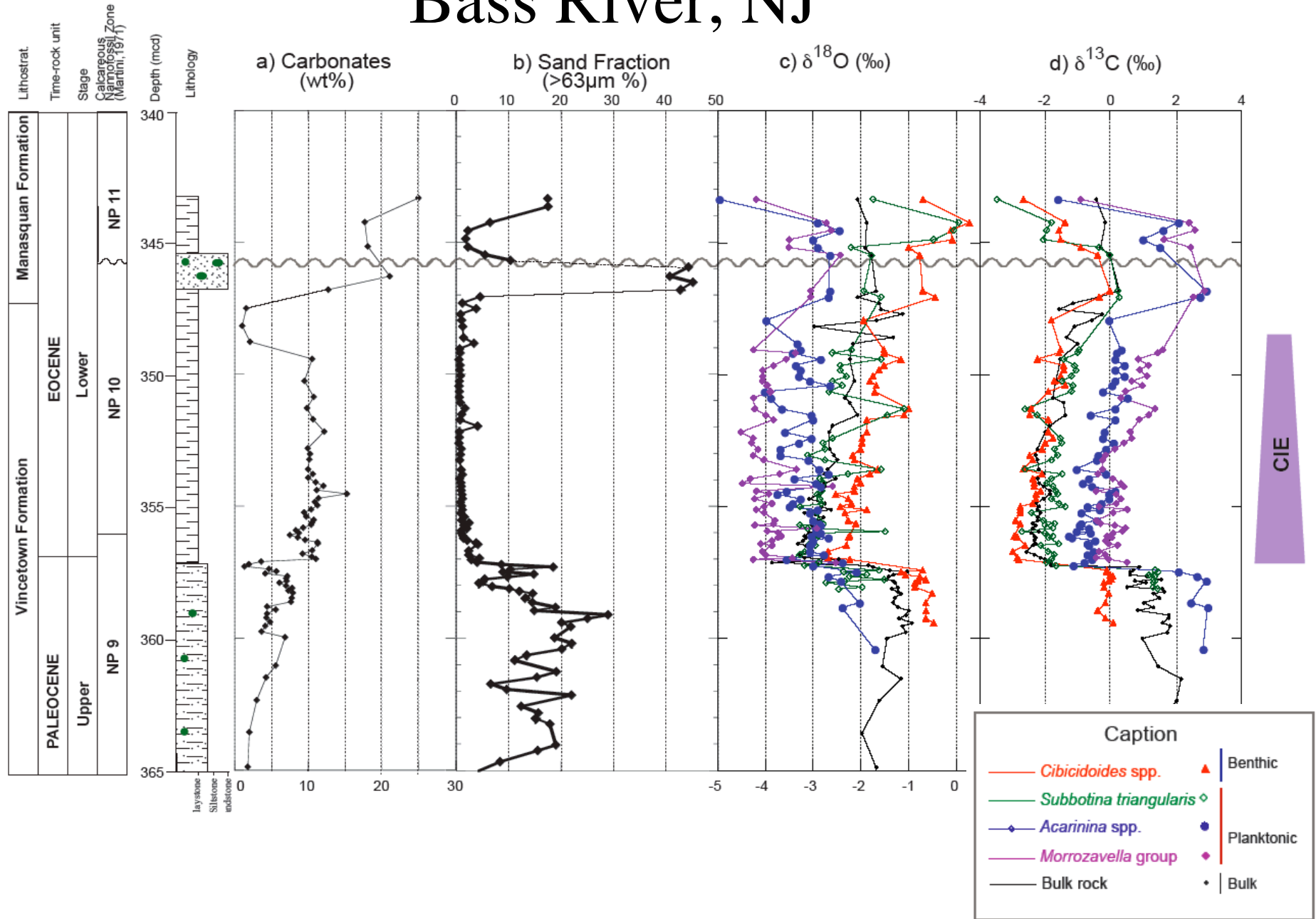


D.

6 μ m 4000X

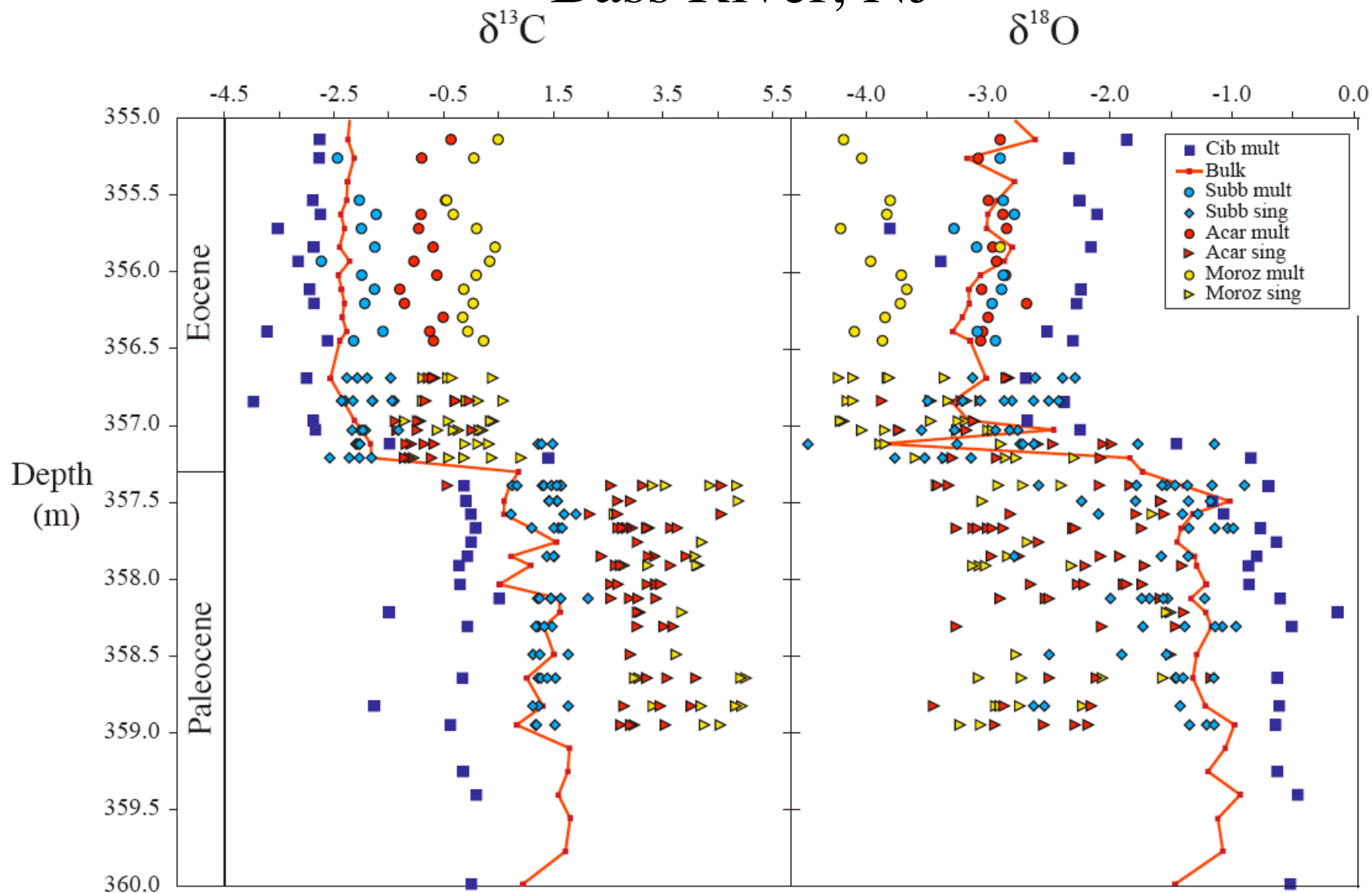
Zachos et al., 2006

Bass River, NJ



John et al., submitted

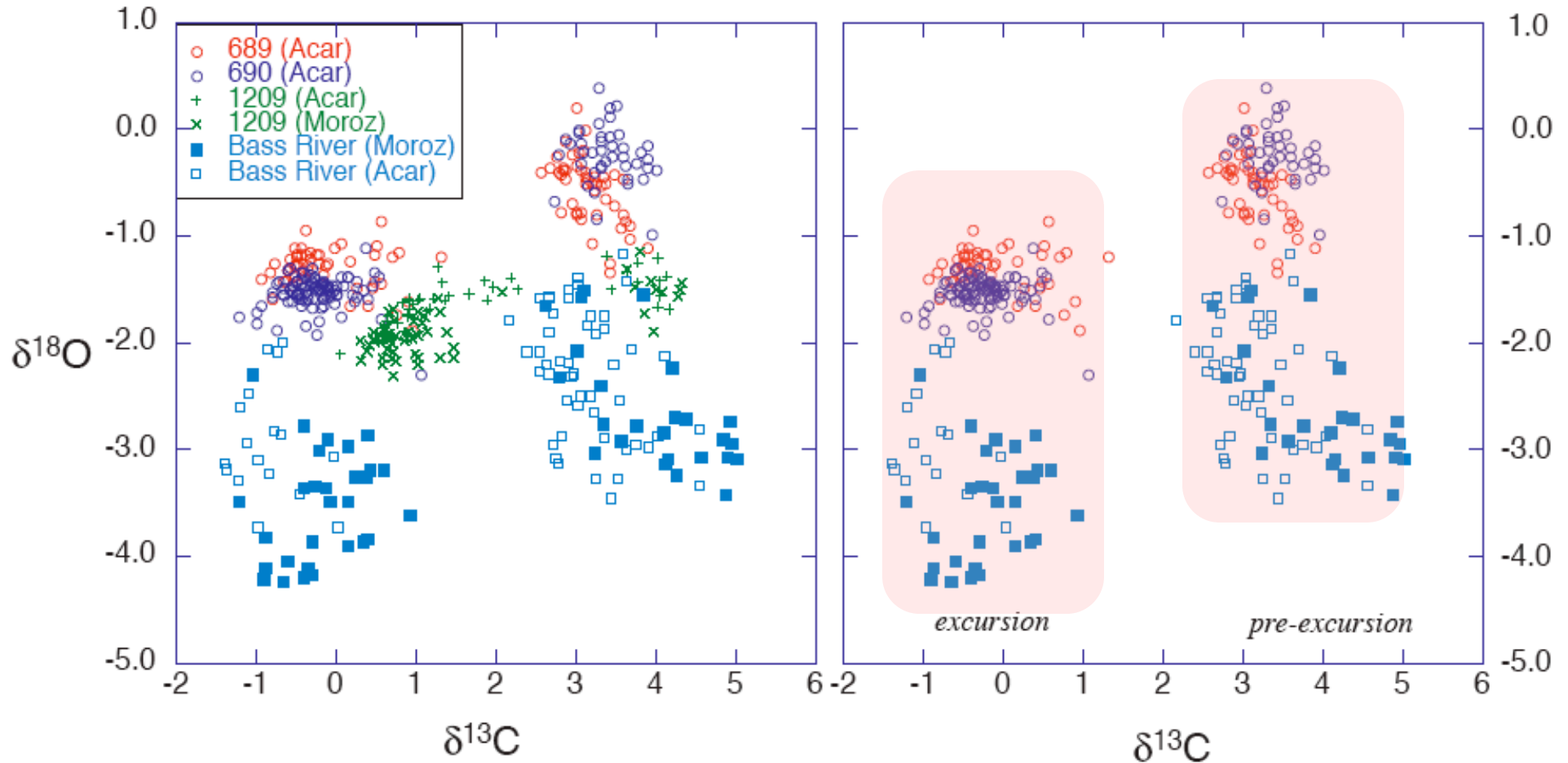
Bass River, NJ



- Multiple/Single Shell Isotope Data

Zachos et al. (2007) Proc.Royal Soc.A

P-E Single Shell C & O-Isotope Data (Southern Ocean, N. Pacific, N. Atlantic Shelf)



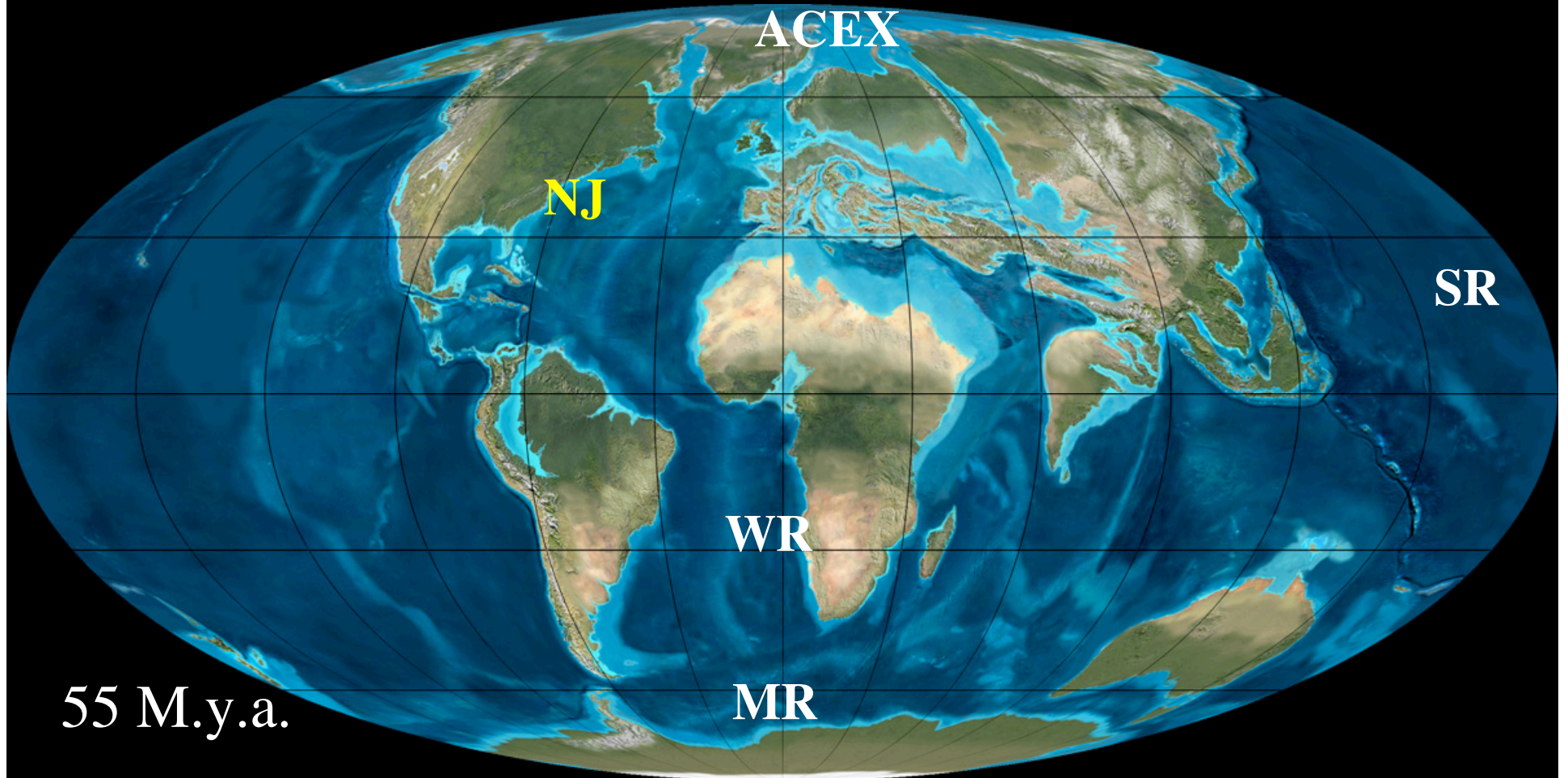
- CIE $\sim 4.0\text{‰}$ (surface ocean)
- Bi-Modal \sim No transitional values

Source of the Massive Carbon Flux?

- Single or multiple?
- Rate of release?
 - ✓ Higher-fidelity records
 - ✓ Single shell method
- Mass of carbon?
 - ✓ Carbon isotope excursion (CIE)
 - ✓ Carbonate saturation changes



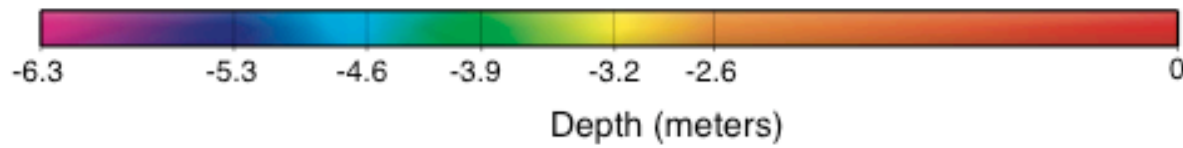
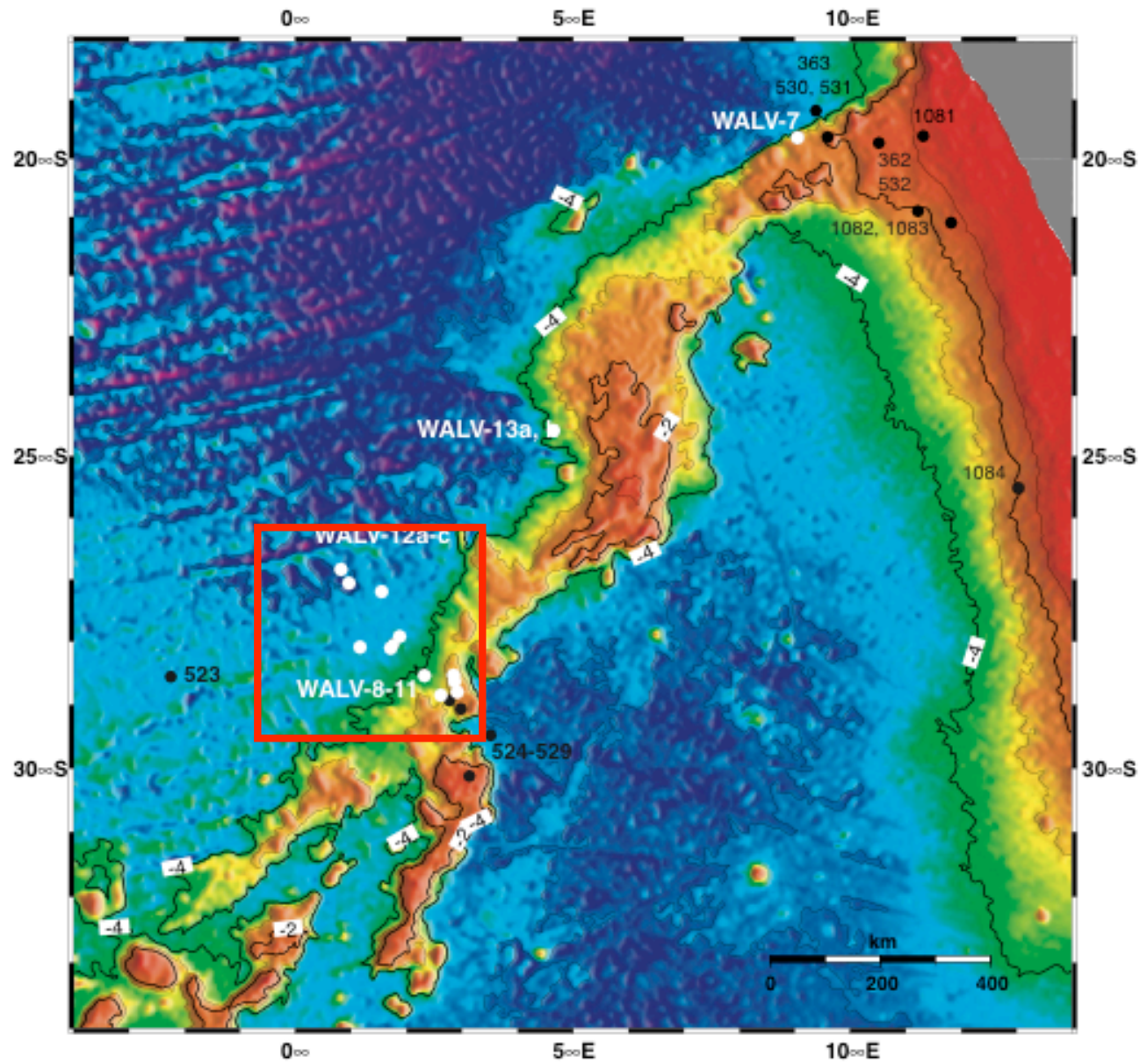
Key PE Boundary Sections



55 M.y.a.

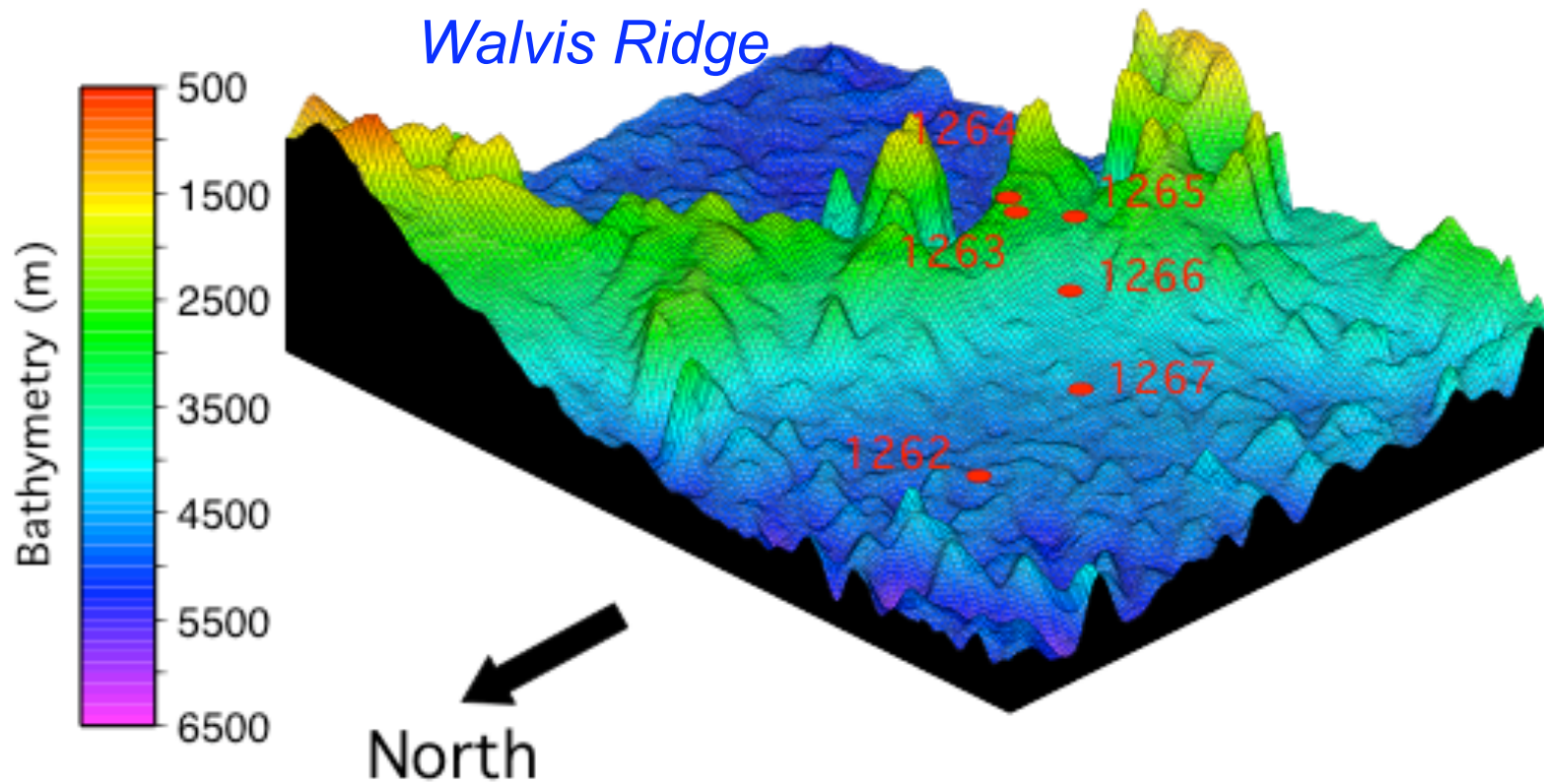
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Walvis Ridge

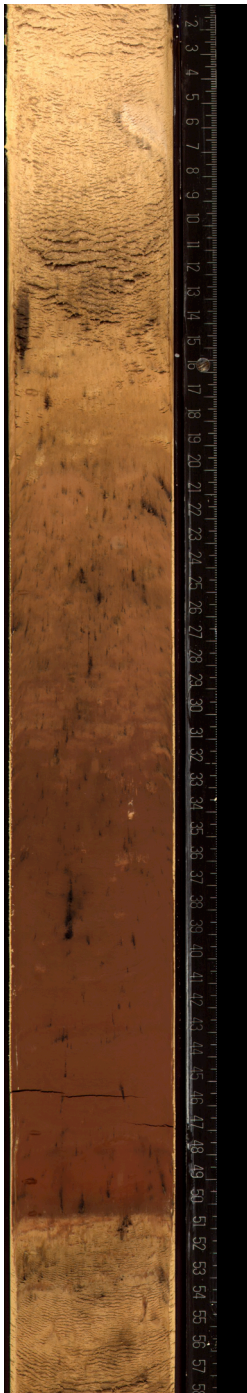


Spiess et al., 2003

ODP Leg 208 Sites

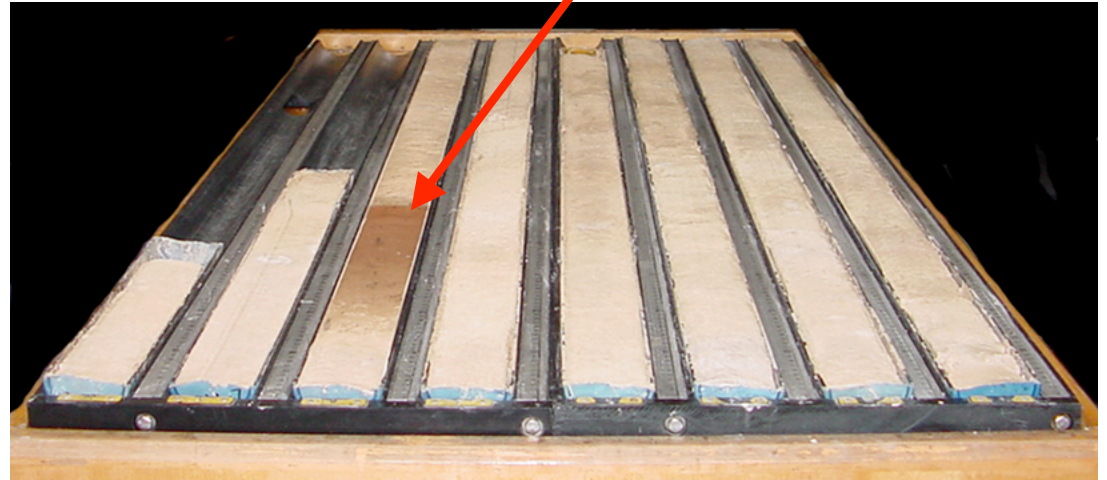


Zachos, Kroon, Blum et al. (2004)



nannofossil ooze

Carbonate Dissolution:
Ocean Acidification

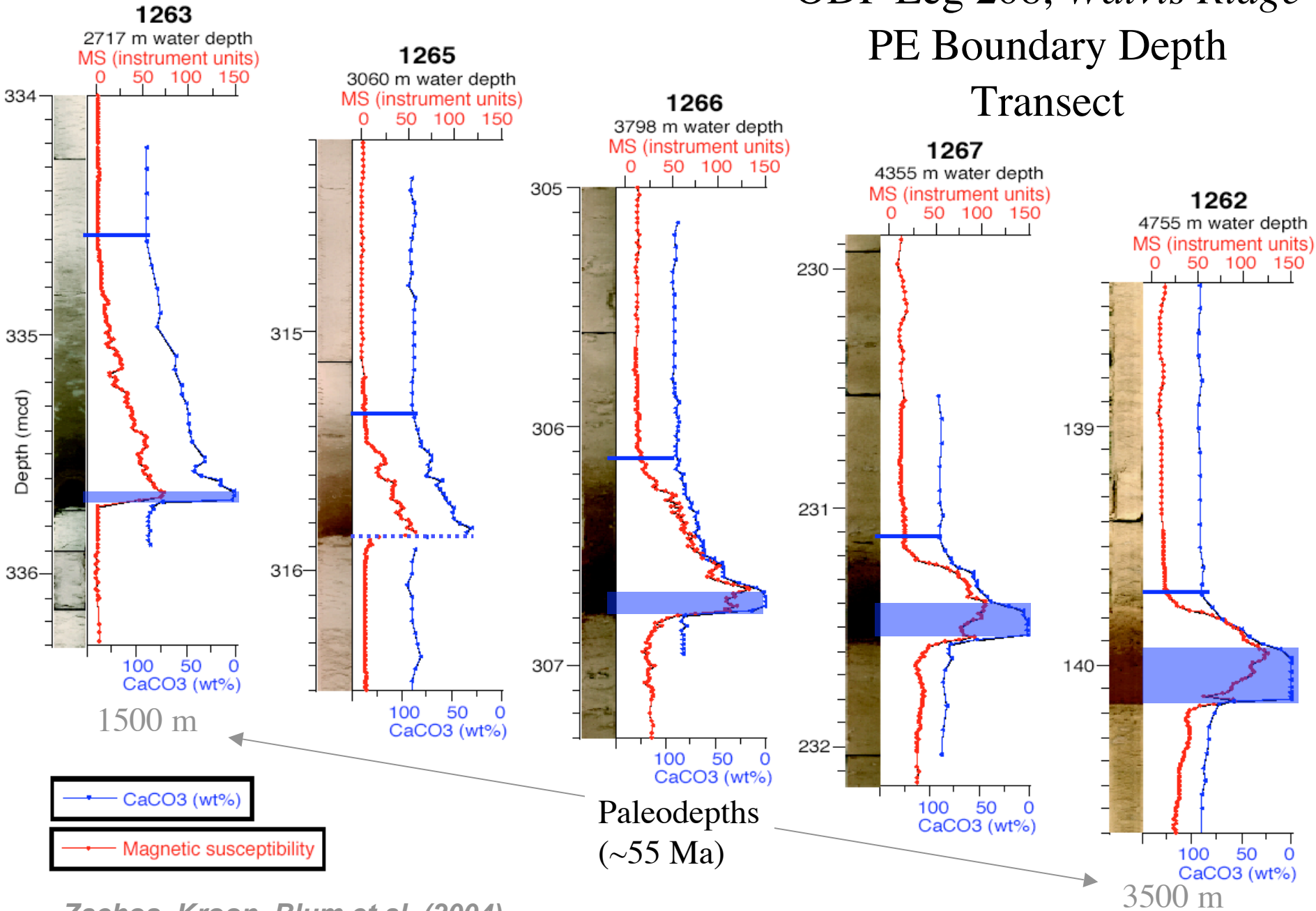


dusky-red clay

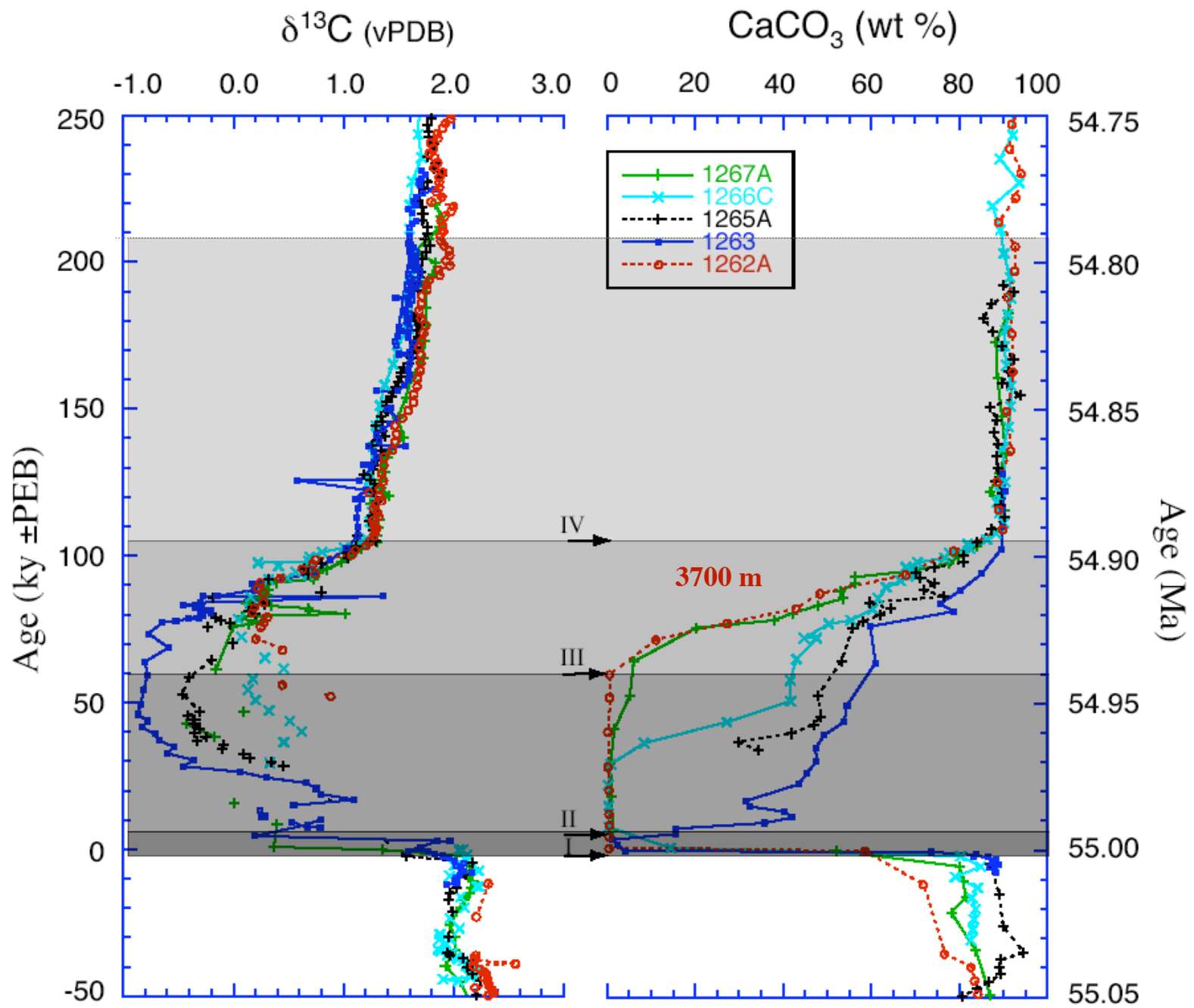
nannofossil ooze

Zachos, Kroon, Blum et al. (2004)

ODP Leg 208, *Walvis Ridge* PE Boundary Depth Transect

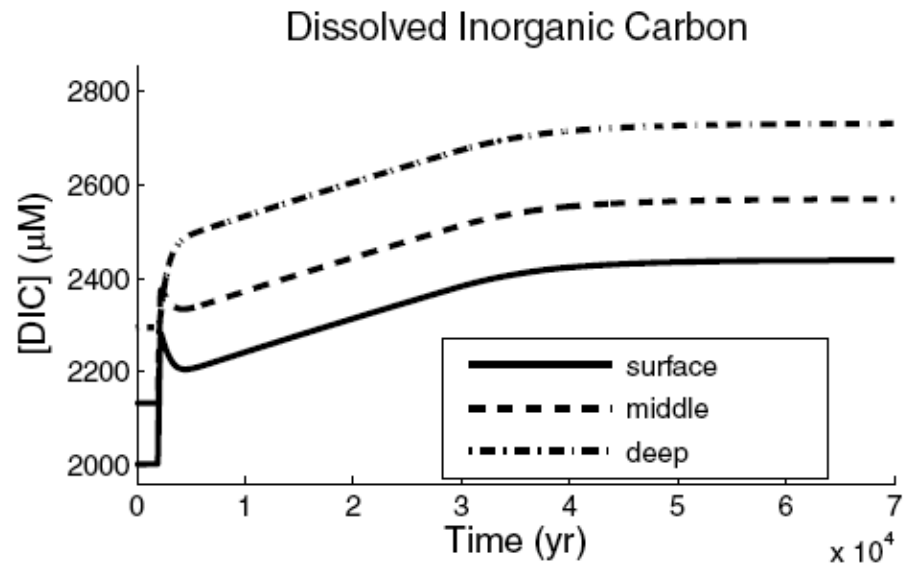
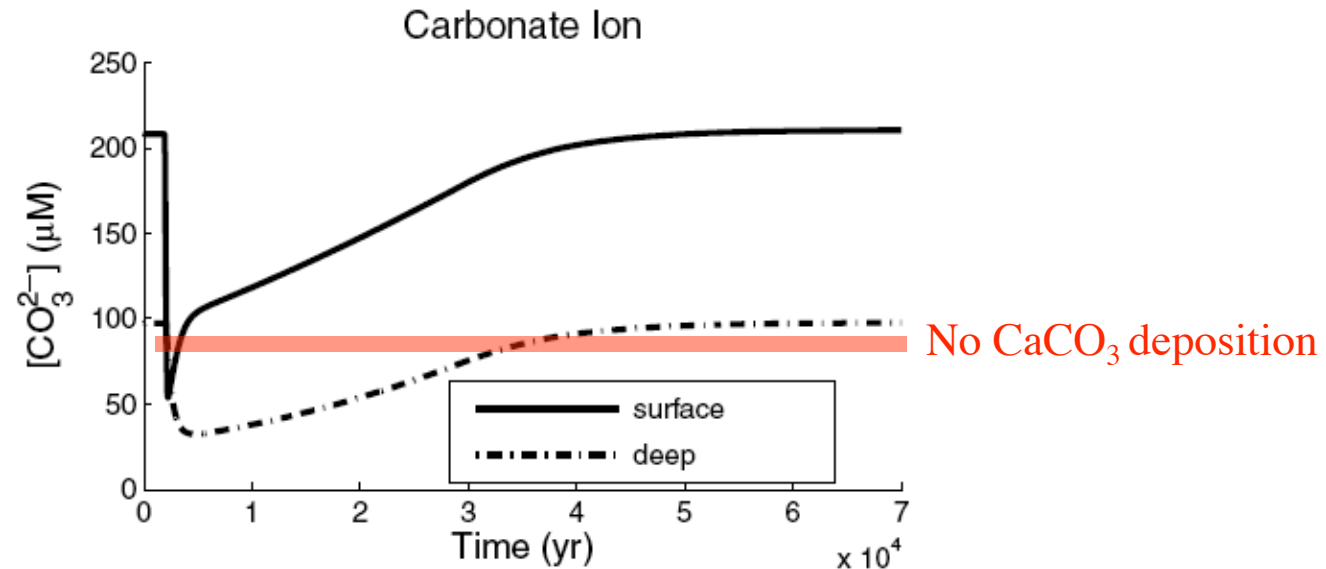


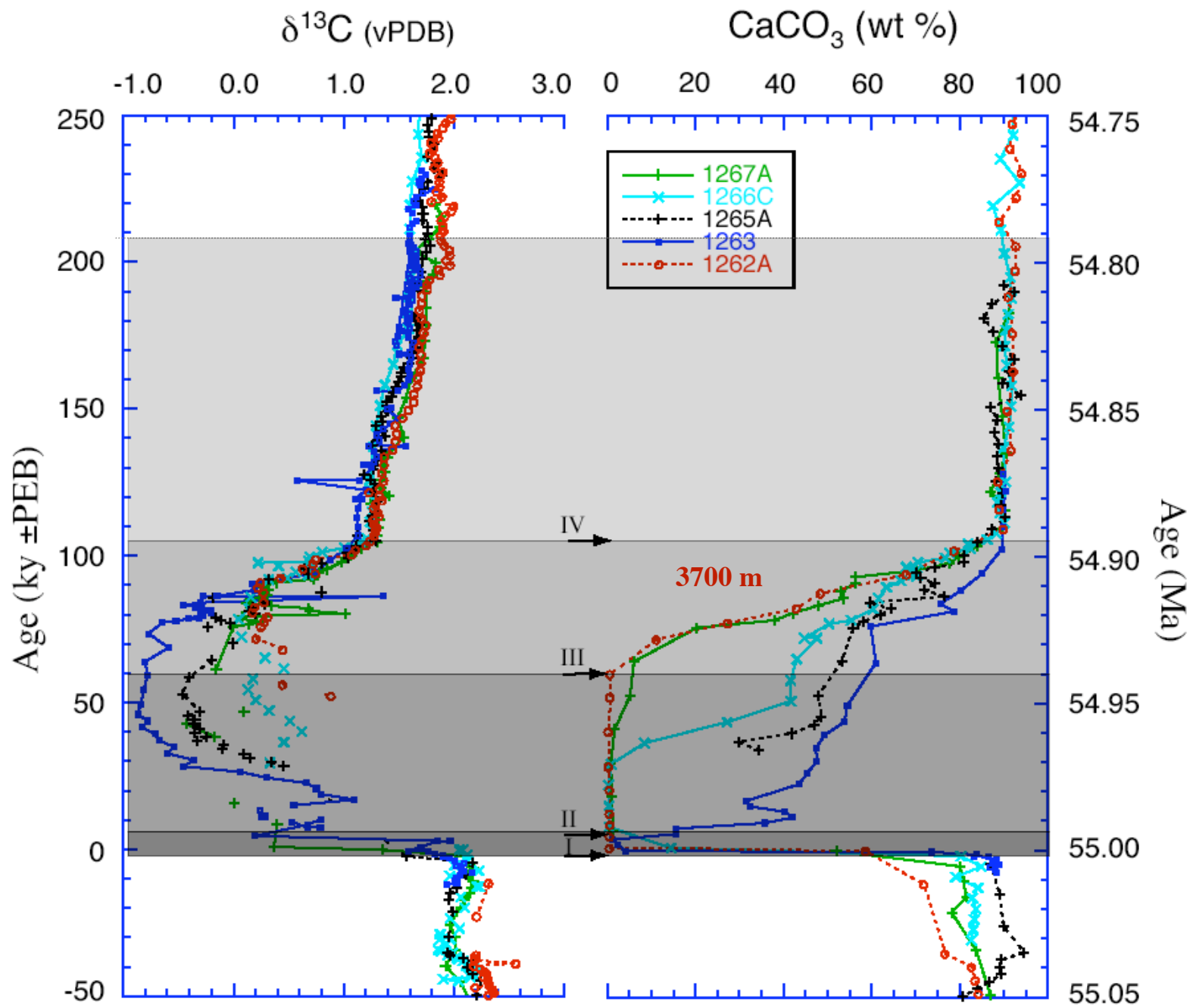
Zachos, Kroon, Blum et al. (2004)



Zachos et al. *Science*, 2005

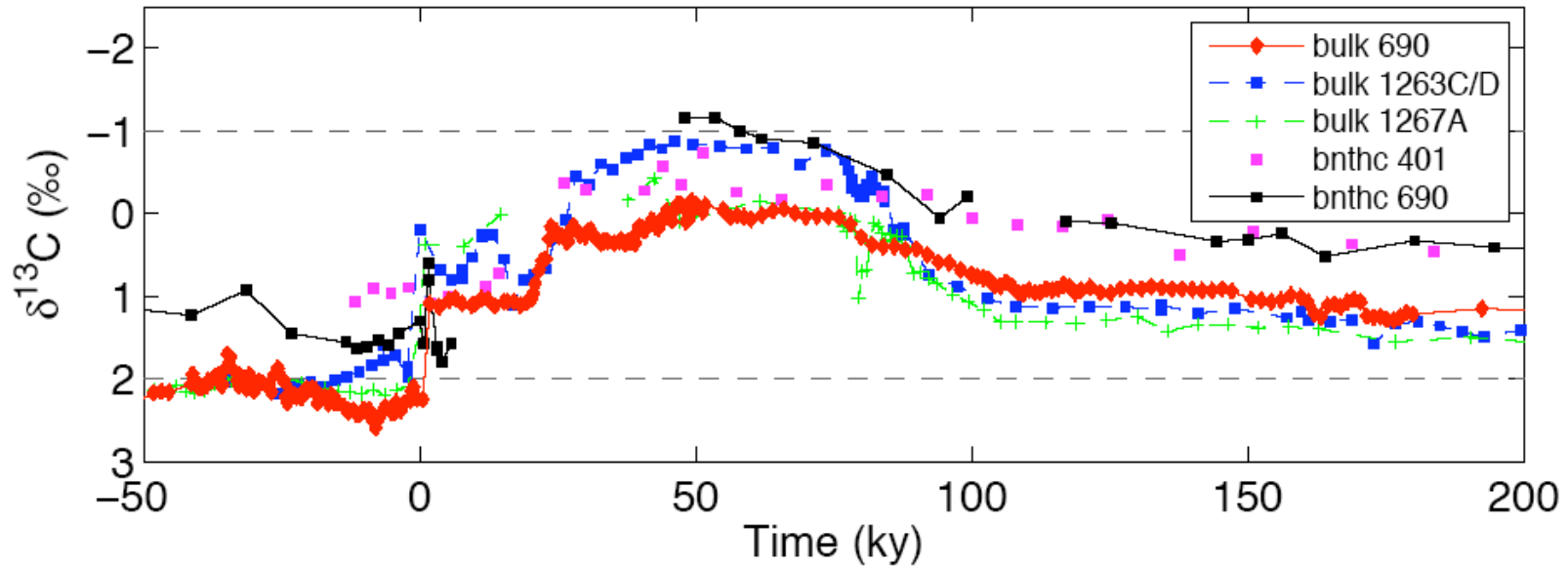
Ocean Acidification and Chemical Erosion (5000 Gt C)





Zachos et al. *Science*, 2005

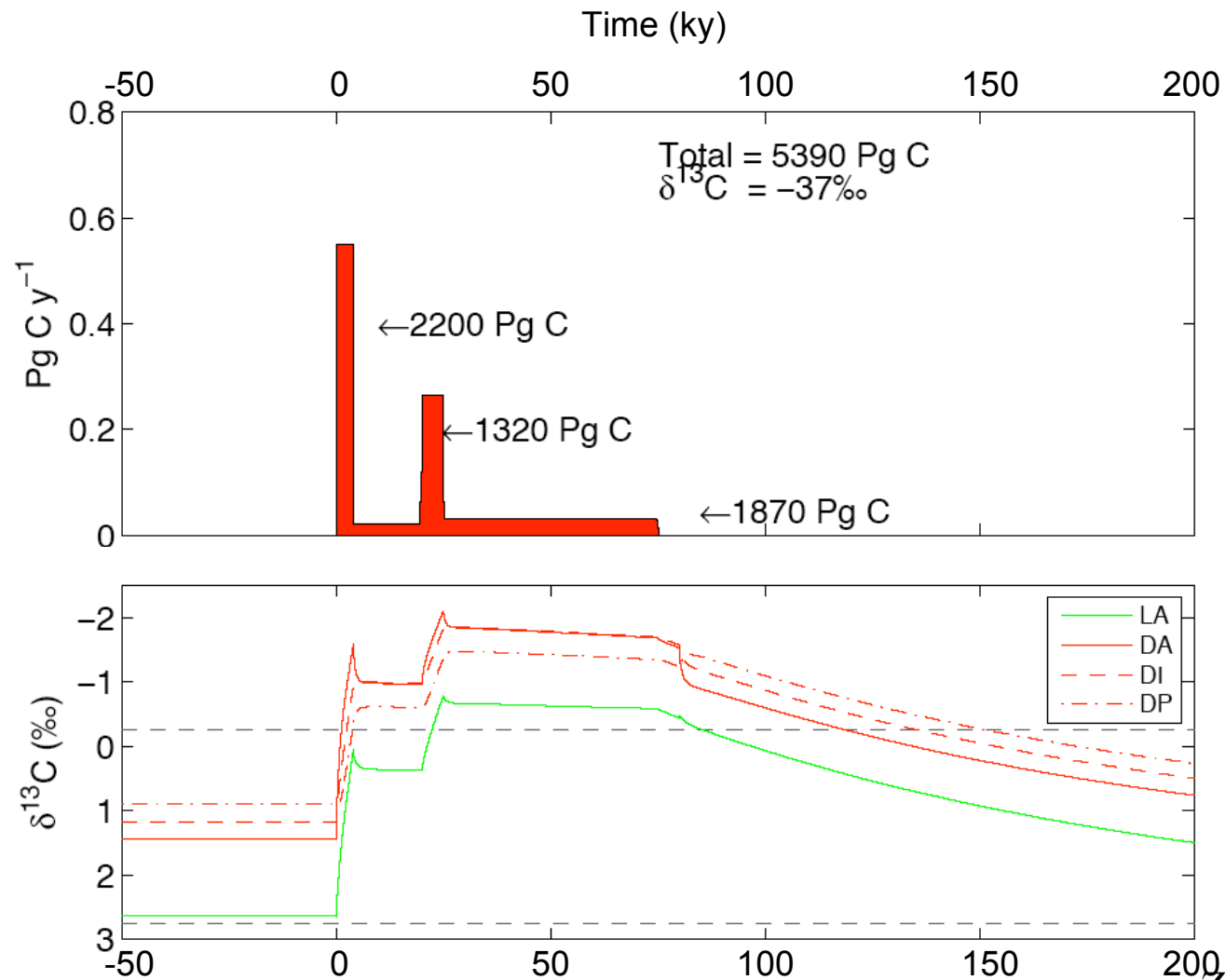
PE Carbon Isotope Excursion



- Bulk Sediment C-Isotope Records
- Structure of CIE?
- Dissolution/Reworking

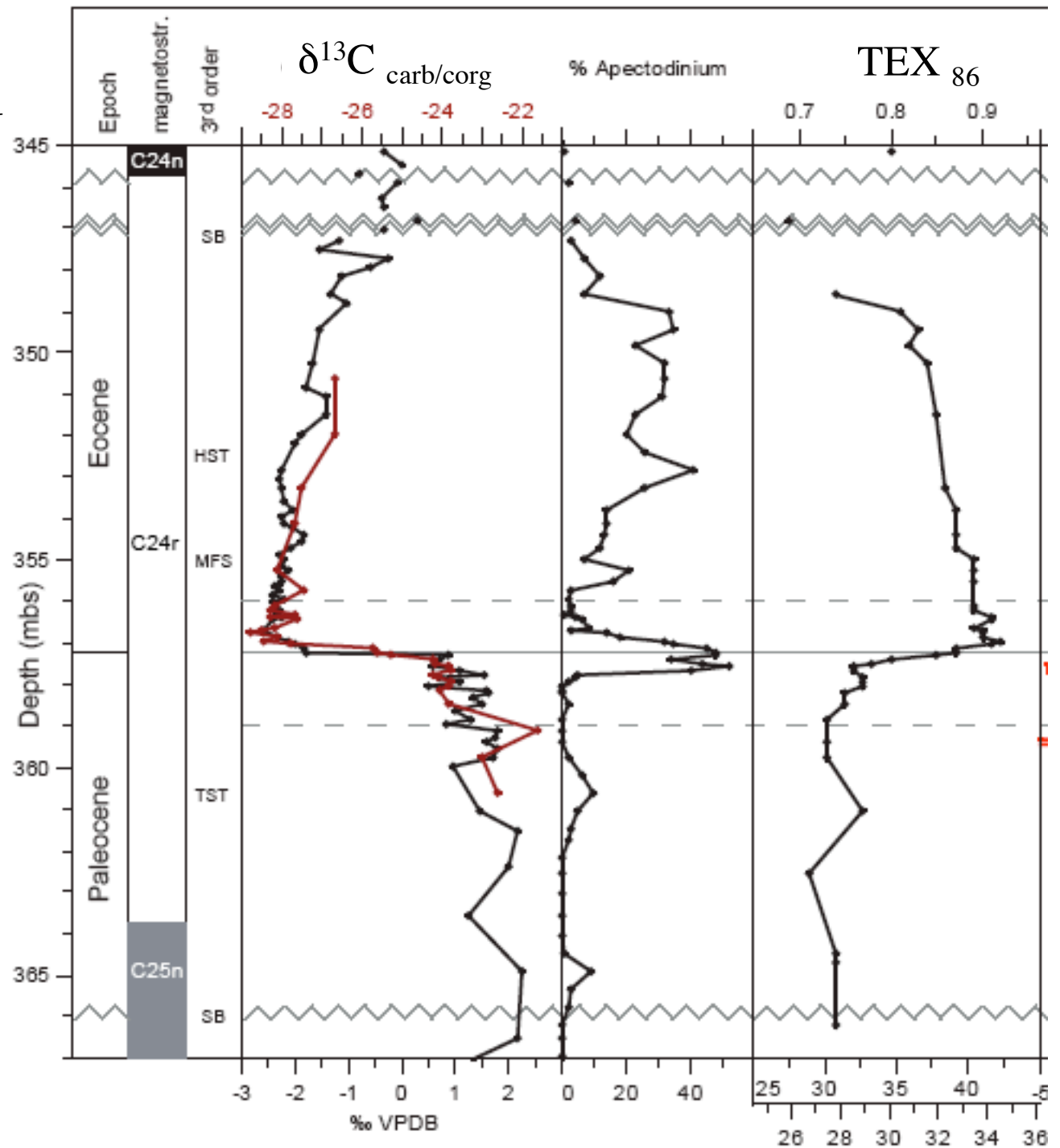
Simulating Massive Carbon Input

- Carbon Cycle Box Model (Walker & Kasting, 1993)
- Ocean/Atmosphere reservoirs

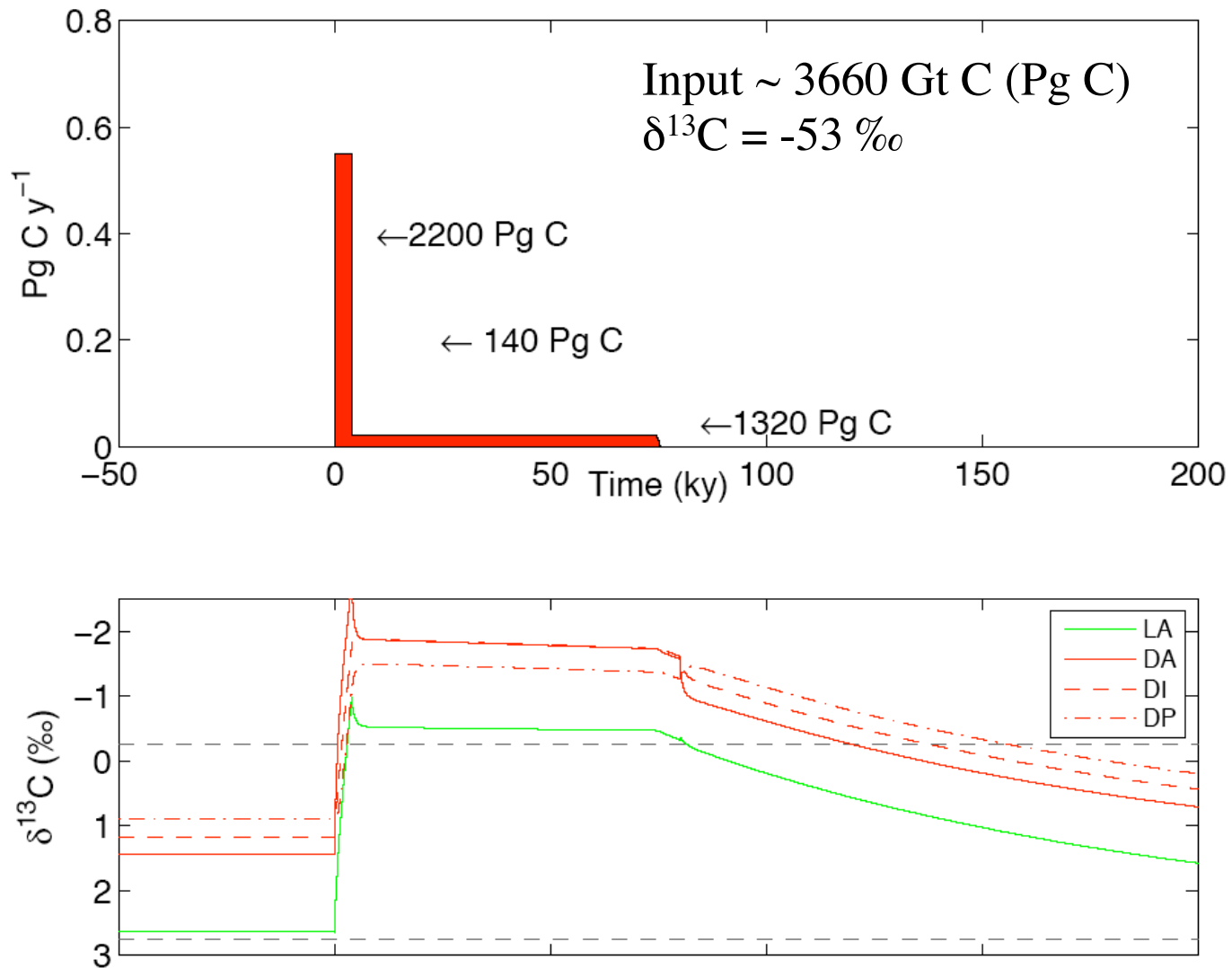


Bass River, NJ

- Siliciclastic Shelf Facies
- Multiple SST Proxies
 - Oxygen Isotopes
 - **TEX₈₆**

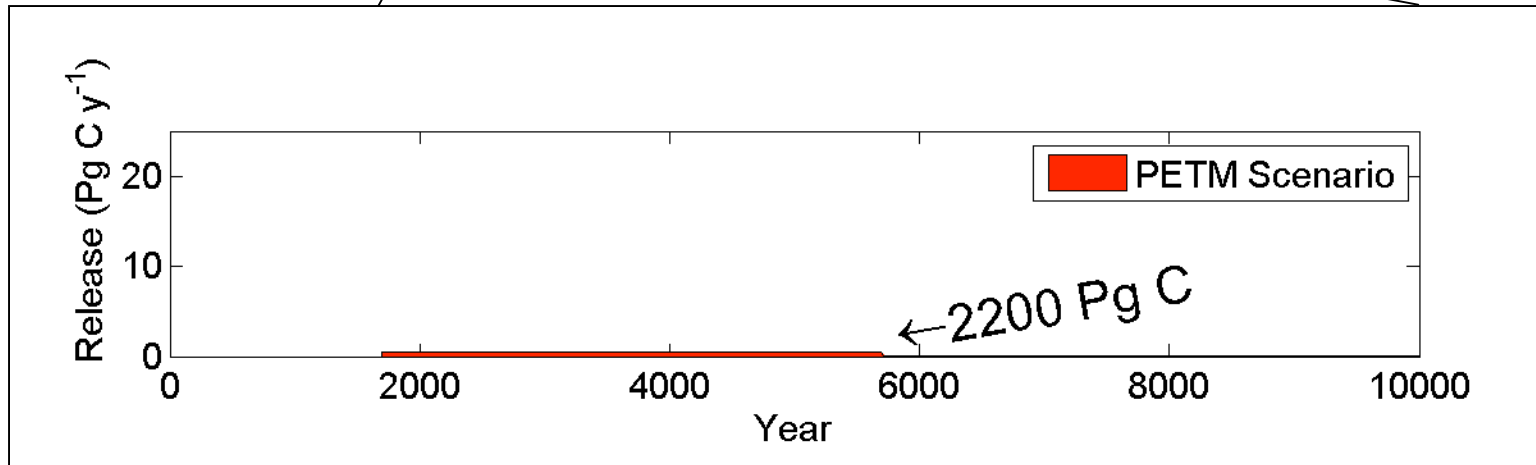
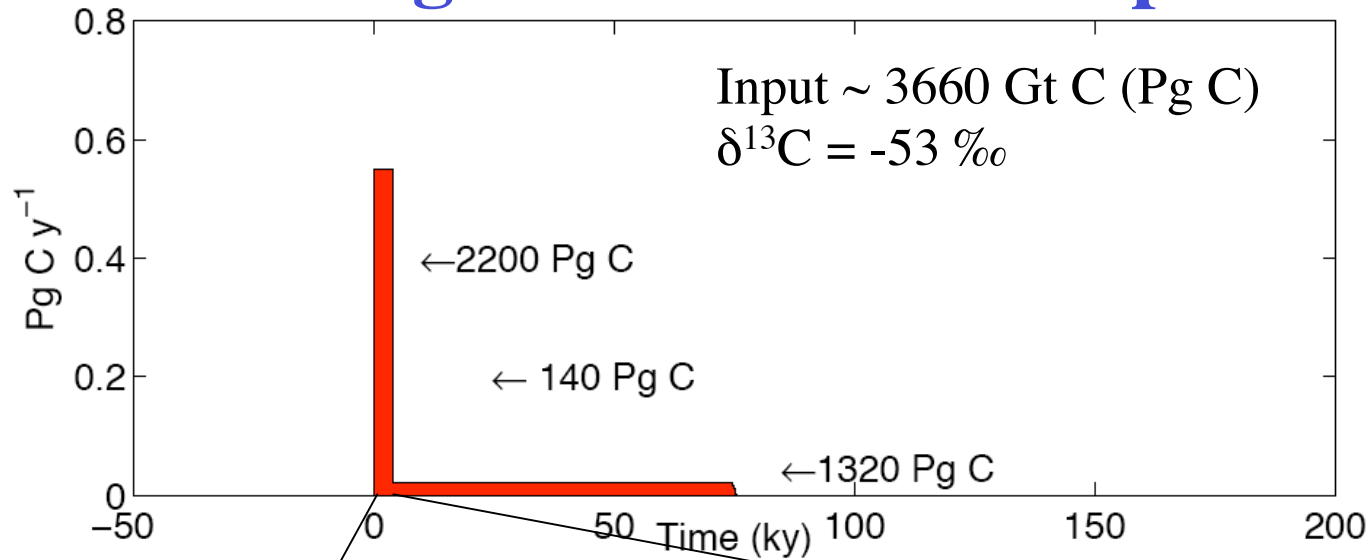


Modeling massive carbon input

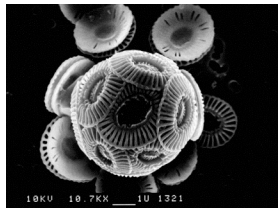
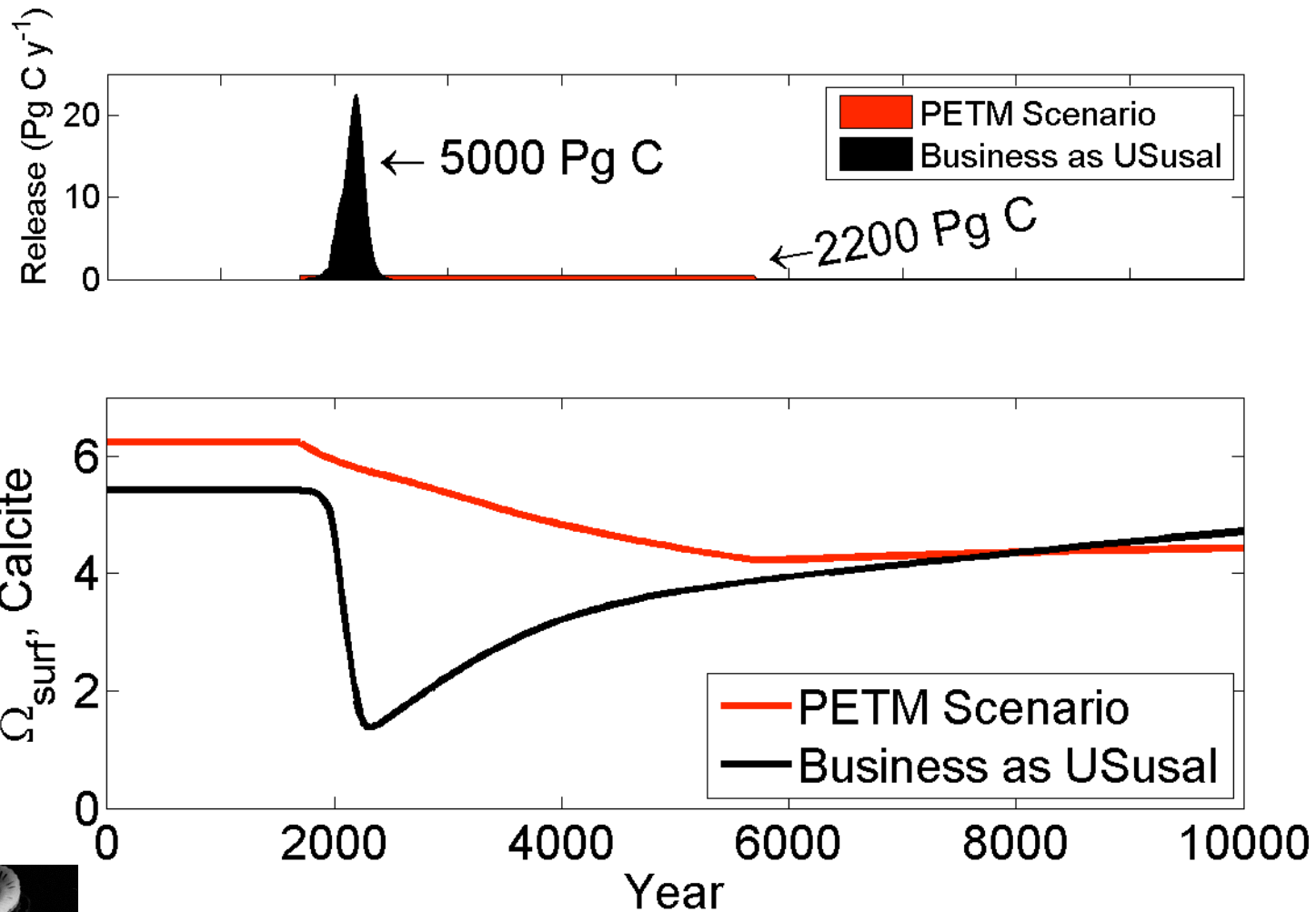


Zeebe et al, in prep

Modeling massive carbon input



Ocean Acidification PETM vs. Future



Zeebe et al, in prep

Future Impact on Marine Calcifiers?

NATURE | VOL 407 | 21 SEPTEMBER 2000 | www.nature.com

letters to nature

Reduced calcification of marine plankton in response to increased atmospheric CO₂

Ulf Riebesell^{*}, Ingrid Zondervan^{*}, Björn Rost^{*}, Philippe D. Tortell[†],
Richard E. Zeebe^{*‡} & François M. M. Morel[†]

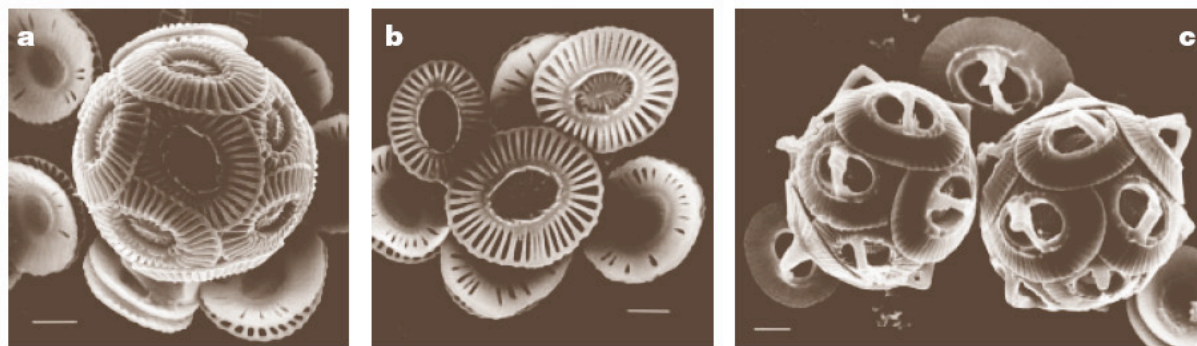
^{*} Alfred Wegener Institute for Polar and Marine Research, P.O. Box 120161,
D-27515 Bremerhaven, Germany

[†] Department of Geosciences & Department of Ecol
Princeton University, Princeton, New Jersey 08544,

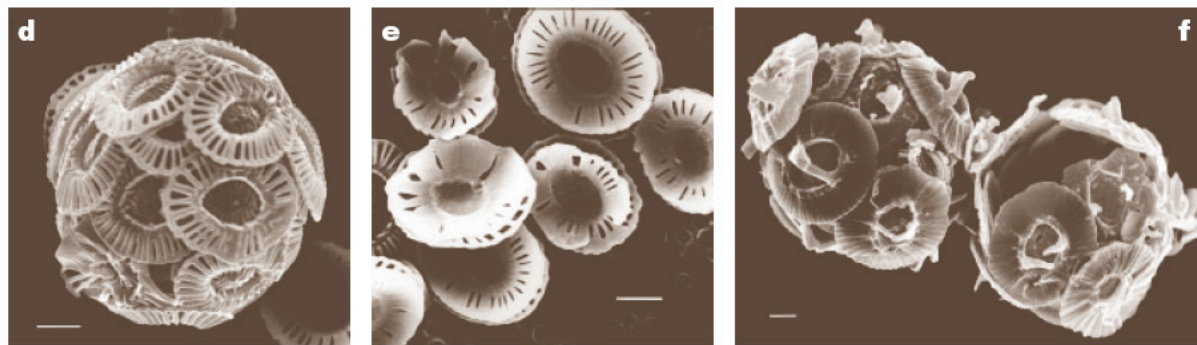
[‡] Lamont-Doherty Earth Observatory, Columbia U
New York 10964, USA

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300 ppmv →

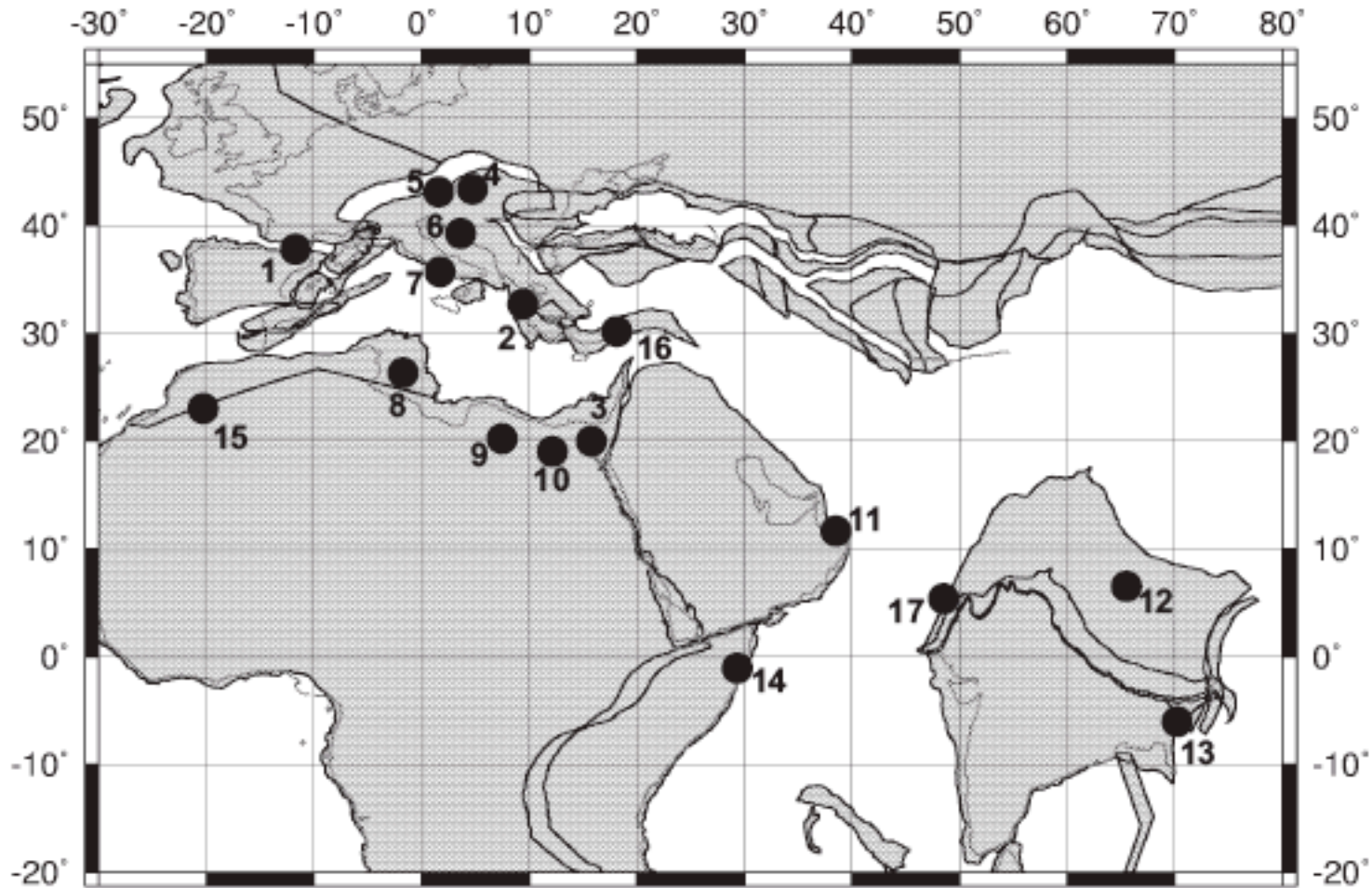


800 ppmv →



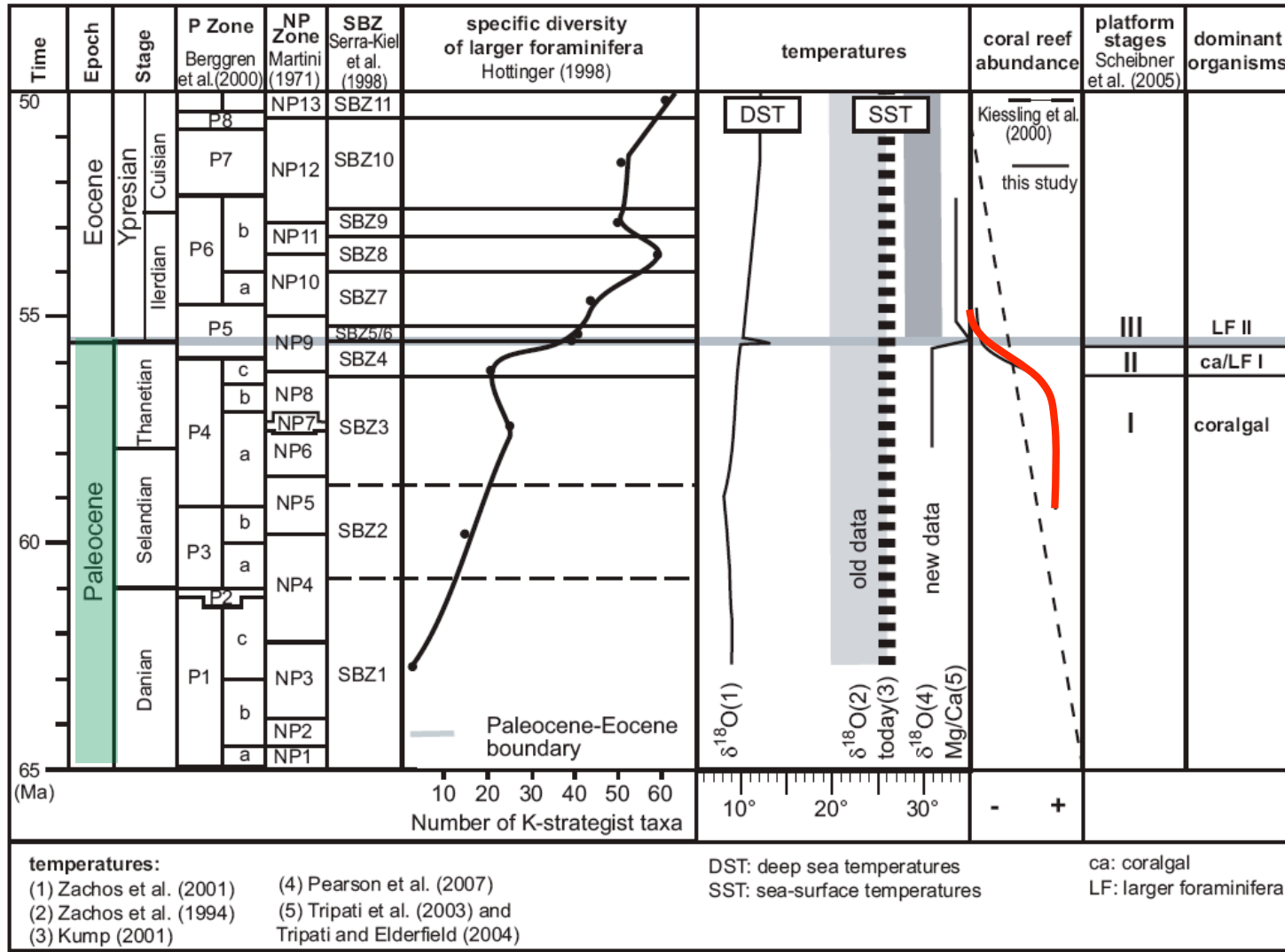
Impact of PETM on Calcifiers?

Early Paleogene Carbonate Platforms



Scheibner and Speijer, 2007, Earth

Decline in Diversity of Larger Foraminifer & Corals



Summary

- CIE Magnitude $\sim 4.0\text{‰}$
- Mass of carbon released during the PETM $\gg 4500$ Gt
- Multiple sources are required
 - ✓ Volcanic (N. Atlantic)
 - ✓ Methane hydrates (feedback)
 - ✓ Terrestrial?
- Decreased pH/warming triggered end Paleocene decline in coral diversity

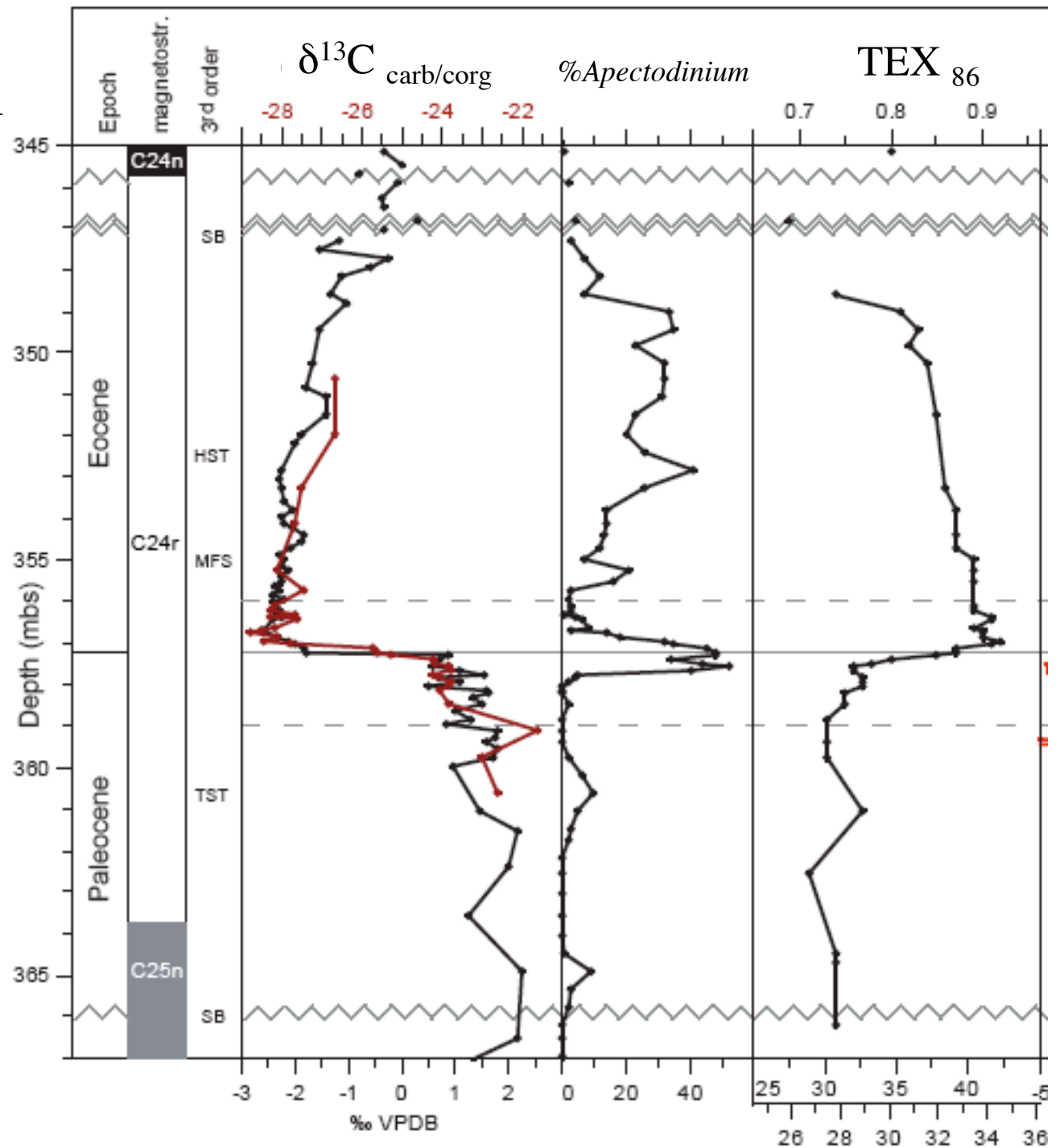
Implications for the future?

- Unabated CO₂ emissions will lead to a severe drop in pH of the surface ocean
- **Positive feedbacks** will likely accelerate the rise in pCO₂
 - ✓ Reduced vertical mixing
 - ✓ Saturation state of the ocean surface
- Will **methane hydrates** dissociate?
 - ✓ Depends on magnitude of warming & propagation of heat into the upper ocean

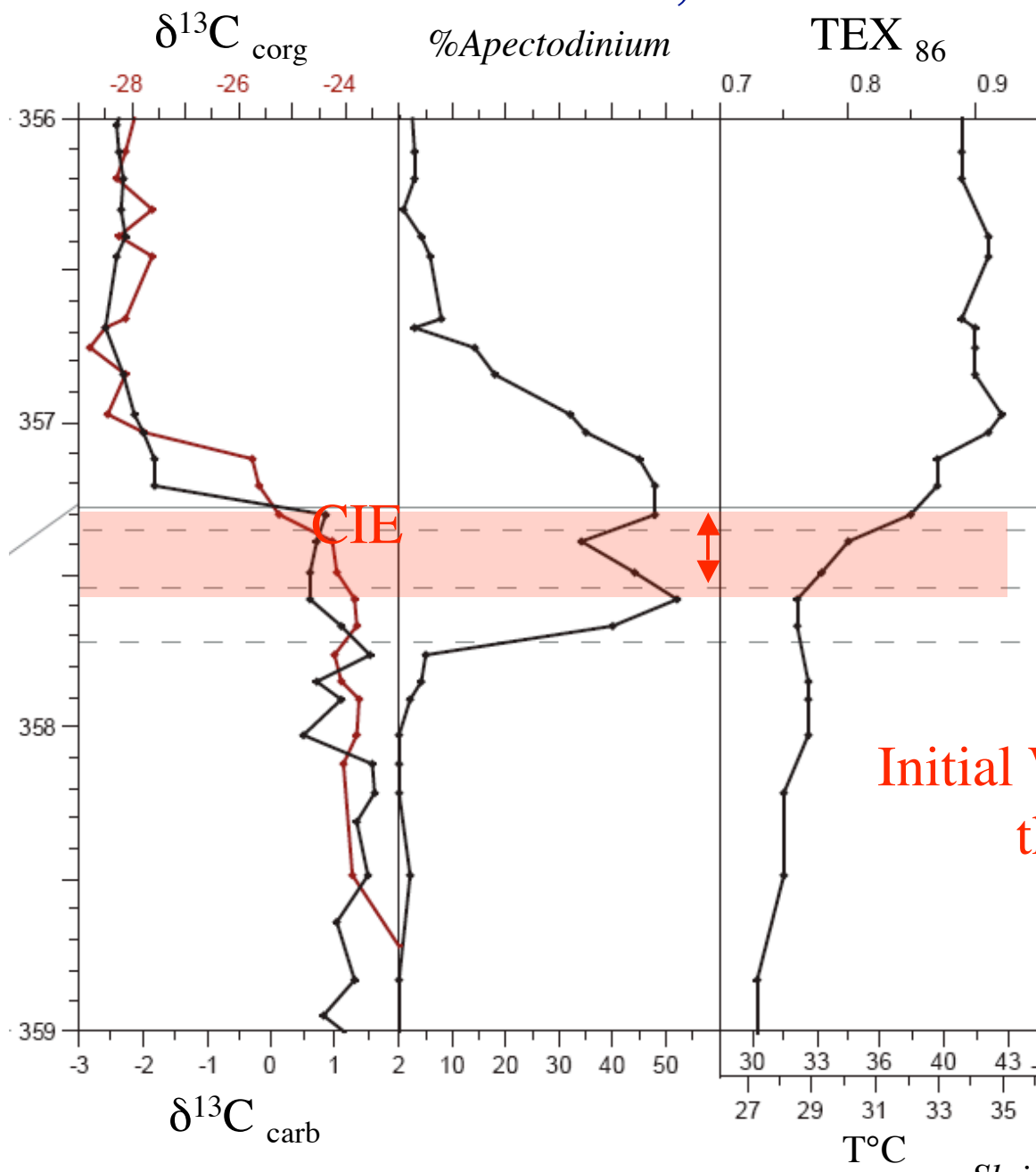


Bass River, NJ

- Siliciclastic Shelf Sequence
- Multiple SST Proxies
 - Oxygen Isotopes
 - **TEX₈₆**



Bass River, NJ

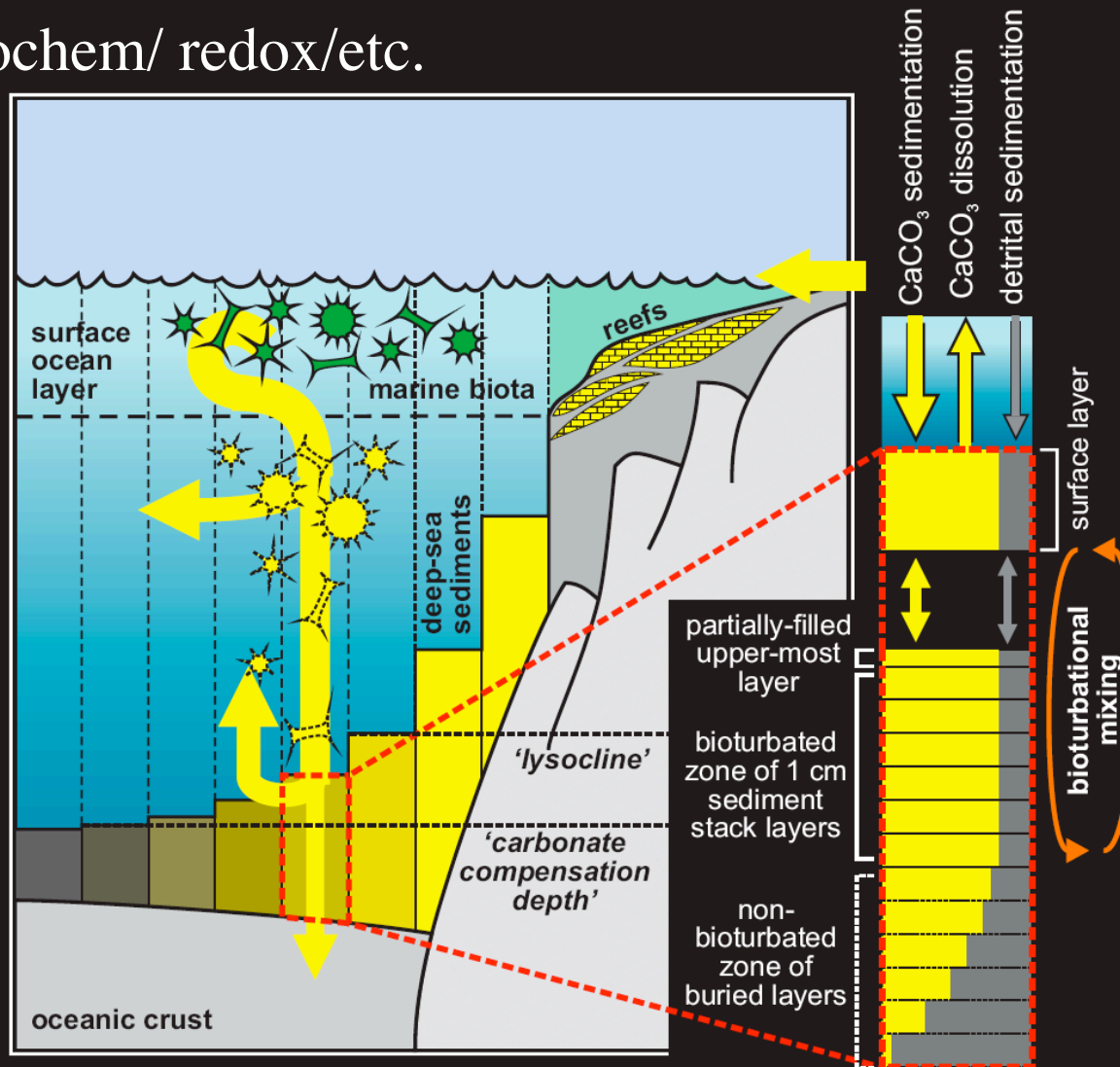


Initial Warming Leads
the CIE!?!

Earth System Models

Ridgwell, 2006; Panchuk, Kump, Ridgwell, in review

- GENIE -3-d w/ circulation
- biogeochem/ redox/etc.

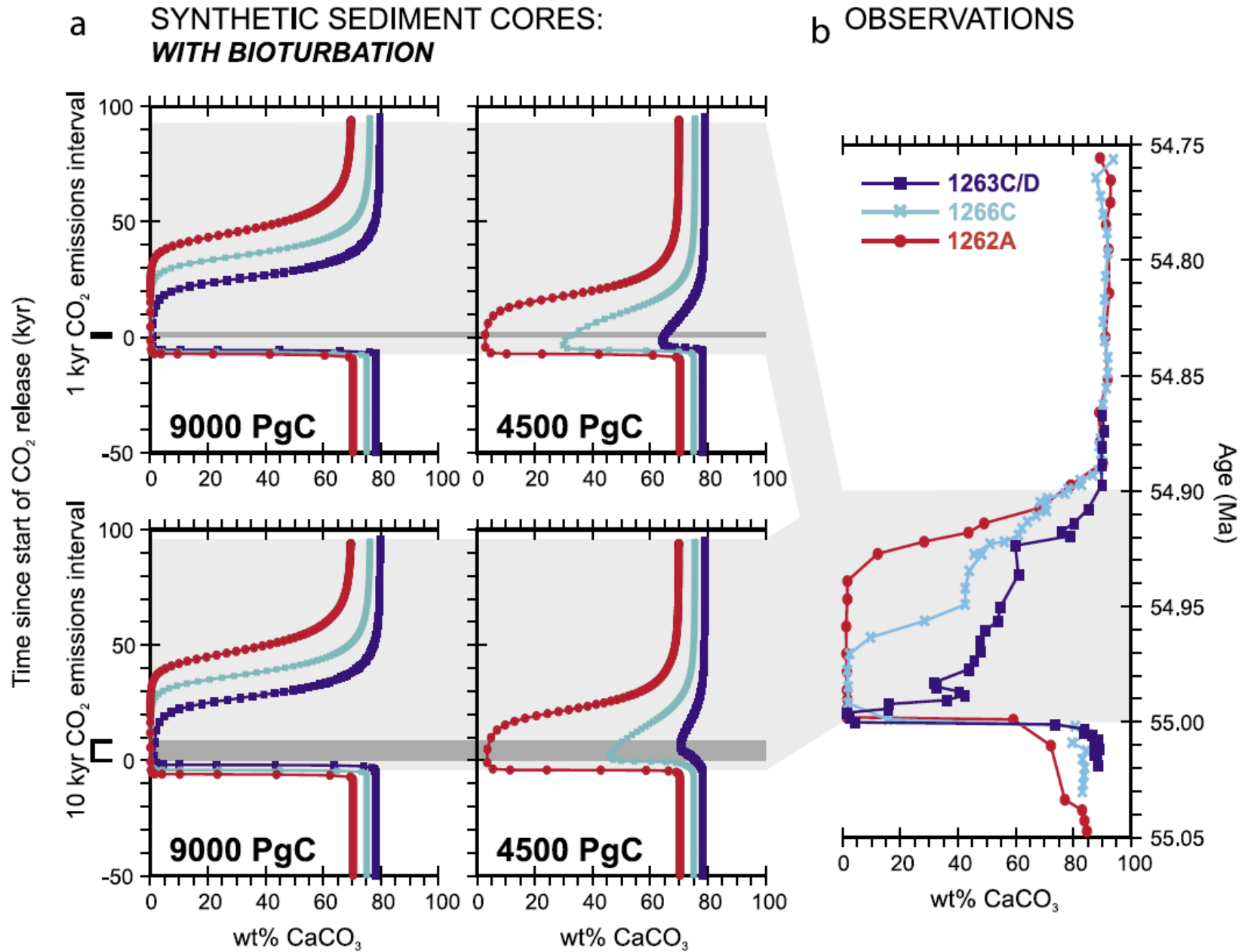


synthetic sediment cores

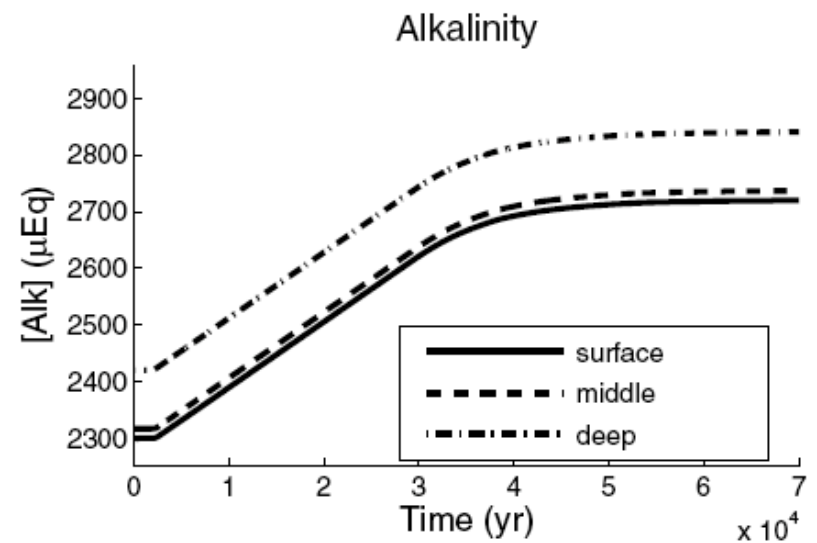
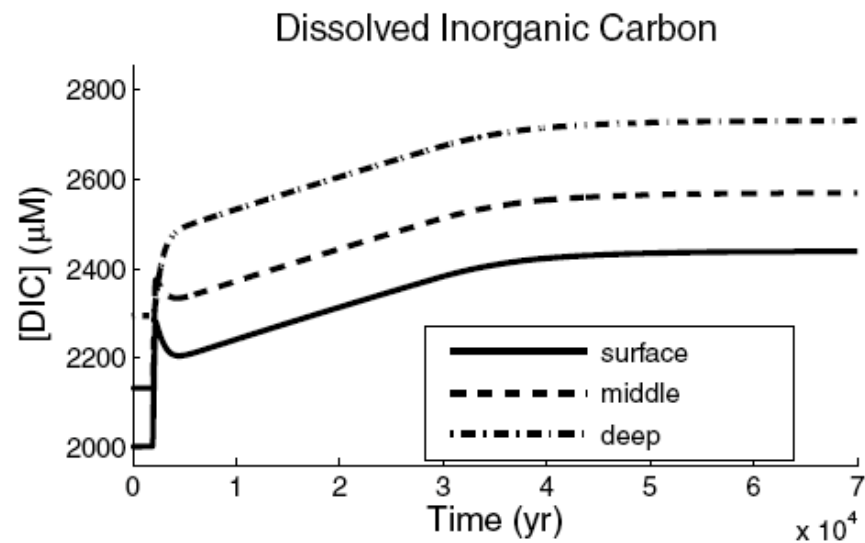
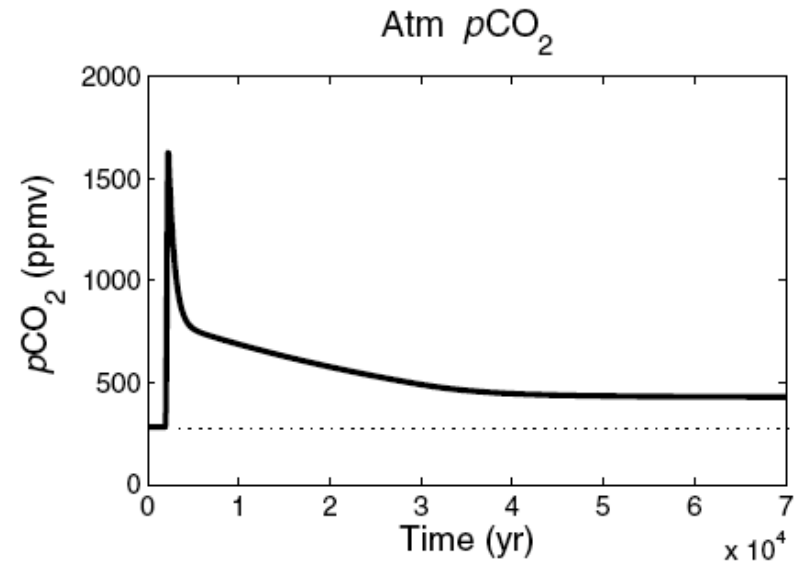
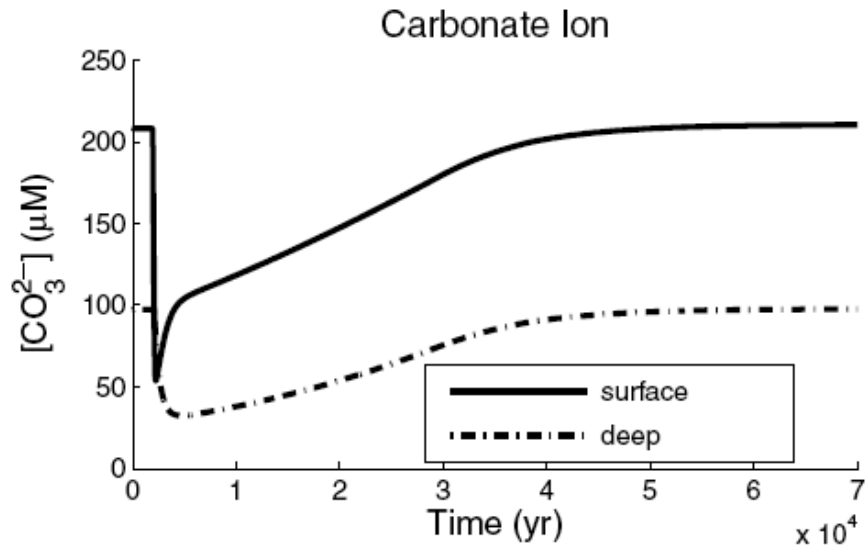
The rate of production at the surface and fate in the ocean interior of organic and inorganic (carbonate) carbon are calculated in the model (based on PO_4 availability).

CaCO₃ preservation in deep-sea sediments is predicted, and historical composition recorded as a function of past changes in accumulation/erosion and bioturbational mixing: generating synthetic sediment cores.

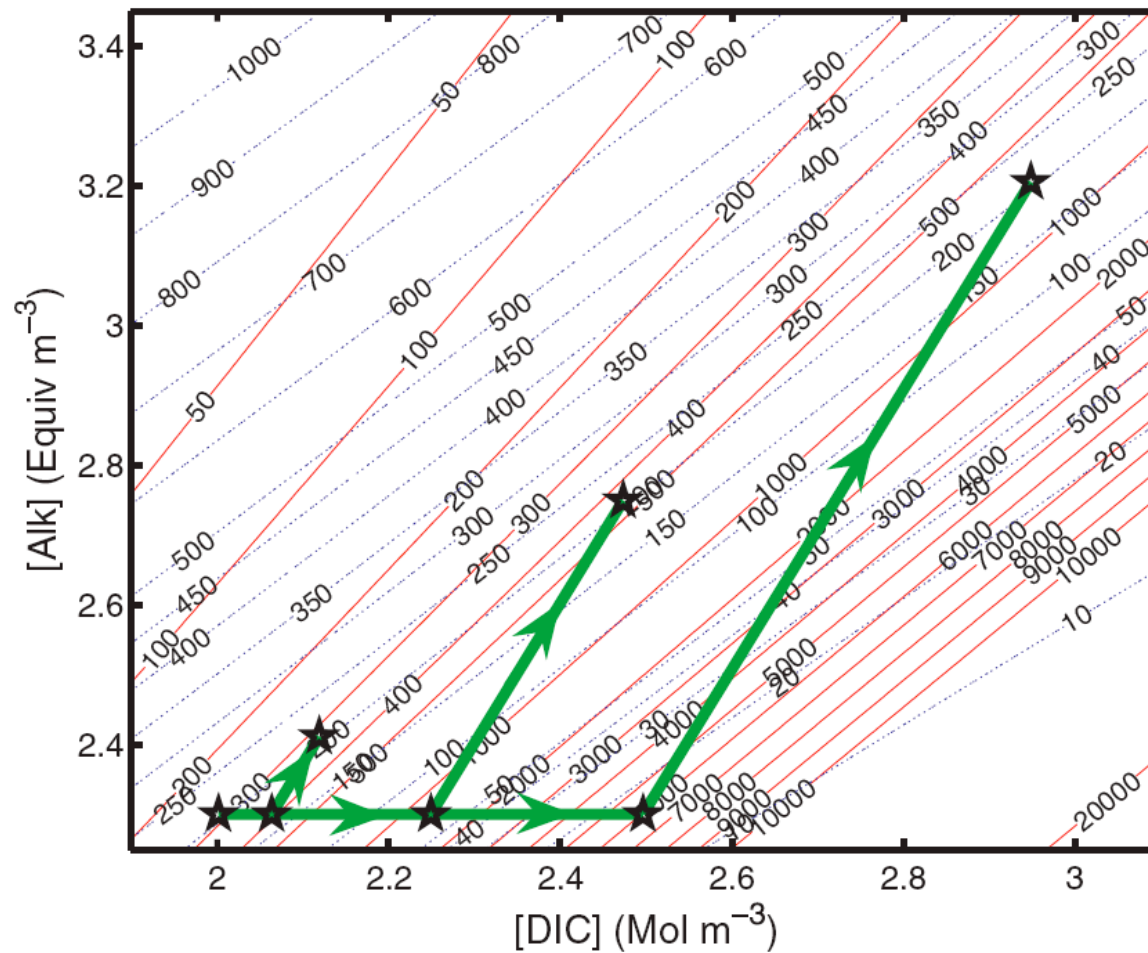
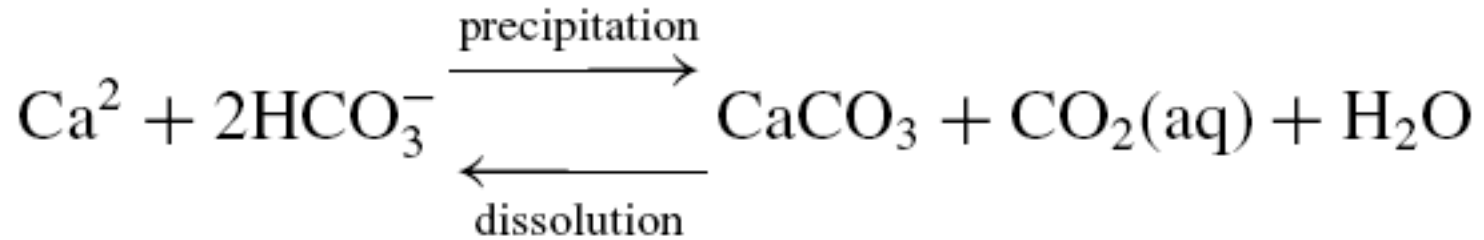
Earth System Models



Long Term CO₂ Legacy



Long Term CO₂ Legacy



Tyrrell et al., 2007