

DR. ZACHOS: YEAH, I DON'T RECALL EVEN  
25 SUBMITTING THE TITLE, BUT THAT'S IN THE PROGRAM.

0446

1 (LAUGHTER)

2 BUT THERE'S ACTUALLY A VERY SHORT ANSWER TO  
3 THAT QUESTION: IS THE PETM AN ANALOG FOR WHAT IS  
4 HAPPENING NOW AND IN THE FUTURE? AND I THINK THE  
5 SHORT ANSWER IS NO; AND I THINK WHAT YOU'LL SEE IS  
6 THAT WHAT THE PETM DOES FOR US, IT ALLOWS US TO TEST  
7 IDEAS ABOUT OUR THEORY ABOUT THE CARBON CYCLE, HOW  
8 THE CARBON CYCLE WORKS. AND THIS WILL BECOME  
9 APPARENT AS YOU MOVE THROUGH THIS PRESENTATION.

10 I START OUT WITH THIS SLIDE, WHICH SCOTT  
11 ALREADY TALKED ABOUT, THE C4M COMPARISON, THE COUPLED  
12 CLIMATE CARBON CYCLE MODEL COMPARISON; AND THE MAIN  
13 POINT OF THIS IS SIMPLE: AS WE MOVE FURTHER AND  
14 FURTHER INTO THE FUTURE AWAY FROM THE CURRENT  
15 CONDITIONS, ONCE WE START TO THINK ABOUT THE SYSTEM  
16 AS A COUPLED SYSTEM AND INVOLVE BIOGEOCHEMISTRY, TIE  
17 THAT IN WITH THE PHYSICS, THIS IS WHERE WE START TO  
18 FIND THAT OUR KNOWLEDGE IS SOMEWHAT LIMITED ABOUT HOW  
19 THESE SYSTEMS SHOULD FUNCTION, AND SO THIS IS A GOOD  
20 EXAMPLE HERE, IT'S ABOUT A 200-PPM SEPARATION BETWEEN  
21 THESE MODELS. PART OF IT HAS TO DO WITH CARBON  
22 UPTAKE ON LAND, AND PART OF IT HAS TO DO WITH  
23 DIFFERENCES IN CARBON UPTAKE BY THE OCEAN. AND SCOTT  
24 HAS ALREADY TALKED ABOUT WHY THAT IS.

25 CLEARLY, THERE ARE FEEDBACKS IN THE SYSTEM  
0447

1 THAT WE NEED TO UNDERSTAND, AND THERE ARE POSITIVE  
2 FEEDBACKS AND NEGATIVE FEEDBACKS; AND FOR THE MOST  
3 PART, I THINK WE WORRY ABOUT THE POSITIVE FEEDBACKS,  
4 AMPLIFIERS.

5 THE ONE EXAMPLE THAT SCOTT TALKED ABOUT WAS  
6 THE OCEAN CIRCULATION, INCREASED STRATIFICATION,  
7 REDUCED VERTICAL MIXING, WOULD TEND TO ALLOW MORE  
8 CARBON TO BUILD UP IN THE ATMOSPHERE, SLOW THE UPTAKE  
9 OF CARBON BY THE OCEAN. AND THIS IS SOMETHING THAT  
10 WE PROBABLY HAVE TO WORRY ABOUT ON THE DECADAL TO  
11 CENTENNIAL TIME SCALE. IT IS A FEEDBACK THAT IS  
12 GOING TO OPERATE FAIRLY QUICKLY.

13 ON LONGER TIME SCALES, SCOTT ALSO MENTIONED  
14 ANOTHER POTENTIAL FEEDBACK, METHANE HYDRATES, A  
15 FAIRLY LARGE RESERVOIR OF CARBON, WHICH COULD  
16 POTENTIALLY BECOME INVOLVED IN RELEASING MORE CARBON  
17 INTO THE ATMOSPHERE ON TOP OF THE ANTHROPOGENIC  
18 CARBON. THIS IS SOMETHING, AS I SAID, THAT MAY  
19 HAPPEN OVER A MILLENNIUM. IT IS PROBABLY NOT  
20 IMMEDIATE FEEDBACK THAT WE HAVE TO WORRY ABOUT.

21 OF COURSE, THERE ARE NEGATIVE FEEDBACKS IN  
22 THIS SYSTEM; AND THESE NEGATIVE FEEDBACKS HELP  
23 EVENTUALLY RESTORE SOME SORT OF EQUILIBRIUM OR STEADY  
24 STATE TO THE SYSTEM.

25 THE REAL ISSUE IS THAT THE POSITIVE

0448

1 FEEDBACKS AND NEGATIVE FEEDBACKS DON'T NECESSARILY  
2 OPERATE AT THE SAME RATES, AND IT'S LOOKING LIKELY

3 THAT THE POSITIVE FEEDBACKS, AT LEAST SOME OF THEM  
4 THAT WE'RE CONCERNED ABOUT, CAN OPERATE AT VERY RAPID  
5 RATES RELATIVE TO THE NEGATIVE FEEDBACKS; AND THIS IS  
6 WHERE A RAPID OR AN ABRUPT CHANGE IN THE SYSTEM  
7 BECOMES SOMEWHAT MORE IMPORTANT.

8 SO RETURNING TO THIS FIGURE JUST SHOWING  
9 THE UPTAKE OF ANTHROPOGENIC CARBON BY THE OCEAN, WHAT  
10 WE DO KNOW IS THAT SOMETHING LIKE 380 GIGATONS OF  
11 CARBON HAVE BEEN RELEASED IN THE ATMOSPHERE, AND  
12 ABOUT 120 GIGATONS OR SO HAVE BEEN ABSORBED BY THE  
13 OCEAN ALREADY. WE KNOW THIS FROM THE WORK OF FEELY  
14 AND HIS COLLEAGUES.

15 NOW, IN ADDITION TO THE DIRECT  
16 MEASUREMENTS, THERE'S SORT OF INDIRECT EVIDENCE THAT  
17 THE OCEAN HAS BEEN TAKING UP ANTHROPOGENIC CO<sub>2</sub>, AND  
18 THIS IS IMPORTANT FOR HOW WE LOOK AT CO<sub>2</sub> CHANGES IN  
19 THE PAST. OBVIOUSLY, FOSSIL FUEL CARBON HAS A VERY  
20 DISTINCT FINGERPRINT, AN ISOTOPIC FINGERPRINT, A LOW  
21 RATIO OF C-13 TO C-12; AND WE'RE ADDING THIS CARBON  
22 TO THE CARBON IN THE OCEAN AND THE ATMOSPHERE. AND  
23 AS A CONSEQUENCE, THE DELTA C-13 OF DISSOLVED CARBON  
24 IN THE OCEAN HAS BEEN DECLINING. WE CAN SEE THIS  
25 VERY CLEARLY.

0449

1 THIS IS A NICE FIGURE THAT BIHM, ET AL, PUT  
2 TOGETHER A NUMBER OF YEARS AGO, AND I HAVE JUST SORT  
3 OF PULLED IT APART. AND THAT'S JUST SHOWING THE RISE  
4 IN CO<sub>2</sub>, AND ONE POINT IS THAT, OBVIOUSLY, WE'VE BEEN  
5 MAKING DIRECT MEASUREMENTS FOR ABOUT 50 YEARS; BUT  
6 BEYOND THAT, WE'RE RELYING ON ARCHIVES TO RECONSTRUCT  
7 CO<sub>2</sub>; AND THAT'S, OF COURSE, THE ICE CORE RECORD. WITH  
8 THE CO<sub>2</sub>, WE HAVE AN ICE CORE, WE CAN ACTUALLY MEASURE  
9 THE CARBON ISOTOPE COMPOSITION OF THAT CO<sub>2</sub>. THIS IS  
10 SHOWING HOW THE DELTA C-13 OF THE ATMOSPHERE HAS  
11 DECLINED, STARTING ROUGHLY AT THE SAME TIME CO<sub>2</sub> LEVELS  
12 STARTED TO INCREASE. AND THEN WE KNOW THAT THE  
13 EXCHANGE OF CARBON BETWEEN THE ATMOSPHERE AND THE  
14 SURFACE OCEAN IS FAIRLY RAPID AND THAT THIS SIGNAL  
15 SHOULD BE TRANSFERRED INTO THE OCEAN; AND THE WAY WE  
16 CAN RECONSTRUCT THAT IS BY LOOKING AT THE CARBON  
17 ISOTOPIC COMPOSITION OF A GEOLOGIC ARCHIVE OF CORALS  
18 WHICH ARE RECORDING THE RATIO OF C-13 TO C-12 IN THE  
19 OCEAN. SO, CLEARLY, THE OCEAN DELTA C-13 HAS BEEN  
20 CHANGING ALONG WITH THAT IN THE ATMOSPHERE, AS WE  
21 WOULD EXPECT. AND THE OVERALL CHANGE HAS BEEN CLOSE  
22 TO ABOUT 1-AND-A-HALF-PER-MILLION DECREASE IN THE  
23 RATIO OF C-13 TO C-12.

24 NOW, IN ADDITION TO THE EFFECTS, OF COURSE,  
25 OF THE SIGNAL ON CARBON ISOTOPES, DICK ALREADY TALKED

0450

1 ABOUT THE IMPACT OF CO<sub>2</sub> ABSORPTION ON OCEAN PH AND  
2 CARBONATE CHEMISTRY. SO THIS IS JUST THE PROJECTION  
3 OUT TO THE FUTURE, 100 YEARS; AND YES, THE PH WILL  
4 DROP BY ABOUT .4 PH UNITS OVER THIS PERIOD OF TIME.

5 THIS PROCESS OF OCEAN ACIDIFICATION WILL  
6 CONTINUE WELL INTO THE FUTURE, AND WE HAVE JUST  
7 CARRIED IT OUT BASICALLY FULL COURSE WHERE WE ASSUME

8 5,000 GIGATONS OF CARBON ARE RELEASED INTO THE  
9 ATMOSPHERE OVER SEVERAL CENTURIES, AND THIS IS  
10 MODELING WORK THAT WAS DONE BY RICHARD ZEEBE AT  
11 UNIVERSITY OF HAWAII. AND IT SHOWS PATTERNS SIMILAR  
12 TO OTHER MODELS LIKE THIS ONE. THIS IS A CARBON  
13 CYCLE BOX MODEL. AND BASICALLY, THE CO2 LEVELS PEAK  
14 AT AROUND 1,800 PPM AND AT ABOUT 3 OR 4 CENTURIES.  
15 AND THEN YOU CAN SEE THAT THE LEVELS RECOVER FAIRLY  
16 QUICKLY; AND AGAIN, THIS HAS MAINLY TO DO WITH  
17 ABSORPTION OF CARBON BY THE OCEANS, AND THEN  
18 EVENTUALLY THEY STABILIZE, ALTHOUGH THEY STABILIZE AT  
19 LEVELS HIGHER THAN PRE-ANTHROPOGENIC LEVELS.

20 AND IT TURNS OUT THAT THERE'S A GOOD REASON  
21 FOR THIS. IT HAS TO DO WITH THE WAY THE OCEAN IS  
22 BUFFERING THE CHANGES IN PH. IT IS DOING SO BY  
23 DISSOLVING AWAY CALCIUM CARBONATE. IT IS TRYING TO  
24 MAINTAIN A CONSTANT CARBONATE ION LEVEL. IN DOING  
25 SO, ALKALINITY RISES, AS DOES DISSOLVED INORGANIC

0451

1 CARBON. AS A CONSEQUENCE, THE EQUILIBRIUM PCO2 LEVELS  
2 ARE HIGHER, AND THEY STAY HIGHER FOR HUNDREDS OF  
3 THOUSANDS OF YEARS. THE MAIN MECHANISM FOR REMOVING  
4 THIS CARBON OR SEQUESTERING IT IS SILICATE  
5 WEATHERING.

6 THIS SHOWS THE EFFECT ON PH; AND AGAIN, IF  
7 YOU GO OUT TO 5,000 GIGATONS OF CARBON, PH WILL DROP  
8 SIGNIFICANTLY. AND THIS IS MAINLY IN THE SURFACE  
9 OCEAN. I'M NOT SHOWING THE RESPONSE IN THE DEEPSEA,  
10 WHICH IS MORE BUFFERED.

11 AND THEN, FINALLY, WHAT I WANT TO SHOW IS  
12 THE CHANGE IN THE CARBON ISOTOPE RATIOS. I ALREADY  
13 SHOWED YOU WHAT IS HAPPENING. INITIALLY, WE HAVE  
14 ALREADY SEEN 1-AND-A-HALF-PER-MIL DECREASE IN DELTA  
15 C-13 AT THE SURFACE OCEAN. IT WILL CONTINUE TO DROP  
16 UNTIL IT GETS TO ABOUT MINUS 5 PER MIL, AND THEN IT  
17 WILL START TO RECOVER. YOU CAN SEE THAT THE DEEPSEA,  
18 THE DEEP PACIFIC DELTA C-13 RATIOS AREN'T GOING TO  
19 DROP AS MUCH. THIS IS TO MAKE ANOTHER SIMPLE POINT;  
20 THAT THE RATE OF RELEASE OF CARBON, THE ABSORPTION,  
21 RATHER, IS BEING MAINLY -- THE ABSORPTION IS BEING  
22 BORNE MAINLY BY THE SURFACE OCEAN, A VERY SMALL  
23 RESERVOIR OF CARBON. THE DEEPSEA IS FAIRLY WELL  
24 BUFFERED.

25 OKAY. SO WE HAVE PREDICTIONS, NUMERICAL

0452

1 PREDICTIONS OF HOW THE CARBON CYCLE IS GOING TO  
2 RESPOND TO ANTHROPOGENIC FORCING. AND IS THERE  
3 ANYTHING WE CAN DO TO ASSESS OR VALIDATE SOME OF THE  
4 PREDICTIONS FROM THESE MODELS? AND WE'RE FAIRLY  
5 LIMITED IN WHAT WE CAN DO, BUT IT TURNS OUT THAT WE  
6 CAN GO BACK IN THE GEOLOGIC PAST AND FIND TIMES WHEN  
7 CARBON DIOXIDE LEVELS MAY HAVE CHANGED AS MUCH AS  
8 THEY'RE GOING TO CHANGE IN THE NEXT SEVERAL HUNDRED  
9 YEARS. AND SO, OBVIOUSLY, I'M TALKING ABOUT THE  
10 PETM.

11 THIS IS JUST TO GIVE YOU SOME SORT OF  
12 GEOLOGIC TIME FRAME OR CONTEXT FOR THIS PARTICULAR

13 EVENT. WHAT'S PLOTTED HERE IS OUR BEST ATTEMPTS TO  
14 RECONSTRUCT PCO2 HERE. YOU CAN SEE THAT CO2 LEVELS  
15 HAVE BEEN FAIRLY LOW FOR THE LAST 20, 25 MILLION  
16 YEARS. YOU GO FURTHER BACK INTO THE CENOZOIC, THE CO2  
17 LEVELS ARE HIGHER. AND OUR ABILITY TO RECONSTRUCT  
18 CO2, ABSOLUTE PCO2, IS NOT VERY GOOD; BUT THE MAIN  
19 POINT HERE IS THAT DURING THIS PERIOD OF HIGH CO2  
20 LEVELS, THE CLIMATE WAS WARM, AND THIS IS AN OXYGEN  
21 ISOTOPE RECORD REPRESENTING GLOBAL CLIMATE; AND WHEN  
22 CO2 LEVELS WERE LOW, THE CLIMATE WAS RELATIVELY COLD.  
23 THE OTHER POINT I WANT TO MAKE HERE IS  
24 SIMPLE: THIS IS THE ANTHROPOGENIC PCO2 PEAK AT  
25 1,800 PPM. AND PEOPLE LIKE TO SAY THE LAST TIME THAT

0453

1 WE'VE HAD CO2 LEVELS THAT HIGH IN THE PLANET, YOU  
2 KNOW, GIVEN THE ERROR BARS HERE, THE UNCERTAINTIES,  
3 REALLY, THE LAST TIME WE'VE HAD CO2 LEVELS ON THIS  
4 PLANET THAT HIGH PROBABLY WAS AROUND 50 MILLION YEARS  
5 AGO, THE EARLY EOCENE, WHICH IS ABOUT THE WARMEST  
6 PERIOD WITHIN THE LAST 65 MILLION YEARS.

7 NOW, THERE IS ONE FEATURE ON HERE. IT'S  
8 HARD TO SEE, RIGHT THERE, RIGHT AT THE P-E BOUNDARY.  
9 AND DURING THIS PERIOD WE'RE SEEING GRADUAL GLOBAL  
10 WARMING; AND SUPERIMPOSED ON THAT GLOBAL WARMING  
11 EVENT IS A VERY SHORT-LIVED, A VERY TRANSIENT  
12 EXCURSION OF GLOBAL TEMPERATURE. AND THAT IS THE  
13 PETM. I'M GOING TO SHOW YOU A MORE DETAILED RECORD  
14 OF THAT IN A SECOND. BEFORE I GO THERE, I JUST WANT  
15 TO SORT OF REMIND YOU OF HOW WE RECONSTRUCT A CLIMATE  
16 USING THESE DEEPSEA ISOTOPE RECORDS.

17 THIS IS AN ISOTOPE RECORD BASED ON THE  
18 OXYGEN ISOTOPE ANALYSES OF BENZIE FORAMINIFERA. AND  
19 OF COURSE, THESE SAMPLES ARE BEING RECOVERED FROM THE  
20 SEAFLOOR VIA OCEAN DRILLING. THIS CORE RIGHT HERE  
21 ACTUALLY SPANS THE PALEOCENE BOUNDARY. I THINK  
22 EVERYONE CAN FIGURE OUT WHERE THAT BOUNDARY MIGHT  
23 ACTUALLY BE. OF COURSE, THE PRIMARY ARCHIVE OR ONE  
24 OF THE TOOLS THAT WE USE FOR RECONSTRUCTING PAST  
25 OCEAN TEMPERATURE AND CARBON CHEMISTRY IS THE

0454

1 CHEMISTRY OF SHELLS OF MICROFOSSILS THAT WE EXTRACT  
2 FROM THOSE CORES. WE MEASURE THE CARBON ISOTOPIC  
3 COMPOSITION, THE OXYGEN ISOTOPES. MIXED-LAYER FORAMS  
4 TELL US ABOUT THE TEMPERATURE OF SURFACE WATERS, THE  
5 OXYGEN ISOTOPIC COMPOSITION OF BENZOIC FORAMINIFERA,  
6 GIVE US DEEPSEA TEMPERATURES. AND SO THAT GLOBAL  
7 COMPILATION THAT I JUST SHOWED YOU WAS A COMPILATION  
8 OF BENZOIC FORAM OXYGEN ISOTOPE RECORDS FROM  
9 SOME 50 OR 60 DEEPSEA CORES.

10 THE CARBON ISOTOPES, OF COURSE, TELL US  
11 ABOUT THE RATIO OF C-13 TO C-12 OF DISSOLVED CARBON  
12 IN SEAWATER. AND SO JUST LIKE THE CORALS, BY  
13 MEASURING THE CARBON ISOTOPIC COMPOSITION OF THE  
14 PLANKTONIC AND BENZOIC FORAMS, WE CAN RECONSTRUCT  
15 PAST CHANGES IN C-13/C-12 RATIO OF DIC.

16 SO RETURNING TO THE PALEOCENE THERMAL  
17 MAXIMUM, THIS IS SORT OF AN ENLARGED BIT ACROSS THE

18 BOUNDARY, AND THIS TOP PANEL SHOWS THE CARBON ISOTOPE  
19 VALUES OF BENZOIC FORMIN FROM CORES IN THE SOUTHERN  
20 OCEAN, PACIFIC, AND SOUTH ATLANTIC; AND EVERY CORE  
21 SHOWS THE SAME PATTERN, WHICH IS THIS VERY ABRUPT  
22 DECREASE IN C-13 AND C-12. WE HAVE KNOWN ABOUT THIS  
23 NOW SINCE 1991. AT THE TIME WHEN WE PUBLISHED THIS,  
24 THIS WAS CONSIDERED TO BE A HIGH-RESOLUTION RECORD.  
25 IT IS ACTUALLY A VERY LOW-RESOLUTION RECORD NOW. BUT

0455

1 THE POINT IS, THIS IS A VERY UNIFORM EXCURSION; AND  
2 WE HAVE RECORDED IT IN ALL MARINE SECTIONS AND ALL  
3 TERRESTRIAL SECTIONS, AS WELL. IT IS ACCOMPANIED BY  
4 THIS NEGATIVE EXCURSION IN THE OXYGEN ISOTOPES, WHICH  
5 IS REPRESENTING WARMING OF THE DEEPSEA ABOUT 5 TO 6  
6 DEGREES CENTIGRADE; AND IN BOTH OF THESE, FROM THE  
7 ONSET TO THE RECOVERY TAKES SOMETHING LIKE 150,000  
8 YEARS. VERY ABRUPT ONSET; GRADUAL RECOVERY.

9 AND THEN FINALLY, MORE RECENTLY, WHAT WE'VE  
10 DISCOVERED, THIS JUST SHOWS THE CARBONATE CONTENT OF  
11 SEDIMENTS FROM SEVERAL CORES IN THE SOUTH ATLANTIC;  
12 AND AGAIN, THIS IS A GLOBAL PATTERN. WE SEE  
13 ESSENTIALLY A GLOBAL DISSOLUTION HORIZON THAT  
14 CORRESPONDS WITH THE ISOTOPE EXCURSION. AND THIS IS  
15 CLEARLY EVIDENCE OF A CHANGE IN OCEAN PH AND OCEAN  
16 ACIDIFICATION AND WIDESPREAD CARBONATE EROSION OF  
17 CARBONATE SEDIMENTS ON THE SEAFLOOR, CHEMICAL  
18 EROSION.

19 OKAY. AS FAR AS TEMPERATURES, THE CLIMATE,  
20 I'M NOT GOING TO SAY A WHOLE LOT. I JUST WANT TO  
21 POINT OUT THAT THESE ARE JUST SOME OF THE ESTIMATES  
22 WE HAVE FOR SEA SURFACE TEMPERATURE ANOMALIES AT  
23 SEVERAL KEY LOCATIONS, AND I'VE ONLY PLOTTED THE  
24 RECORDS FOR THOSE SITES WHERE WE HAVE MORE THAN A  
25 SINGLE PROXY. WE HAVE USED MULTIPLE PROXIES TO

0456

1 RECONSTRUCT TEMPERATURE. AND BASICALLY, IT'S A  
2 FAIRLY UNIFORM WARMING GLOBALLY, ANYWHERE BETWEEN  
3 6 TO 8 DEGREES CENTIGRADE. IN A COUPLE OF CASES, WE  
4 HAVE ABSOLUTE TEMPERATURES DURING THE PEAK OF THE  
5 THERMAL MAXIMUM. SO THIS IS JUST GIVING YOU SOME  
6 INDICATION OF HOW WARM IT GOT IN PLACES LIKE  
7 ANTARCTICA, 20 DEGREES C. THIS IS A MORE RECENT  
8 RECORD SUGGESTING TEMPERATURES AS WARM AS 23 DEGREES  
9 C IN THE ARCTIC.

10 SO, CLEARLY, A FAIRLY SIGNIFICANTLY WARMER  
11 PLANET. WE DON'T ACTUALLY HAVE TEMPERATURES FOR THE  
12 TROPICS AT THE MOMENT, AND IT'S PROBABLY A GOOD  
13 REASON FOR THAT.

14 AS FAR AS THE OTHER CLIMATIC ENVIRONMENTAL  
15 PERTURBATIONS, THIS IS JUST A LIST OF SORT OF THE  
16 IMPORTANT ONES THAT WHEN YOU LOOK AT THE LIST,  
17 BASICALLY IT'S EVERYTHING YOU WOULD EXPECT WITH  
18 GLOBAL WARMING: CHANGES IN PRECIPITATION PATTERNS.  
19 WE'VE EVEN FOUND EVIDENCE NOW OF INCREASED FREQUENCY  
20 OF EXTREME WEATHER EVENTS DURING THIS PERIOD,  
21 WILDFIRES. THE EFFECTS ON BIOTA ARE SUBSTANTIAL.

22 WE DON'T GET MASSIVE EXTENSIONS. I THINK

23 THIS HAS SOMETHING TO DO WITH THE RATE AT WHICH THIS  
24 EVENT UNFOLDS, AND I'LL EXPLAIN THAT IN A SECOND.

25 BASICALLY, THE PATTERNS THAT WE SEE ARE  
0457

1 VERY CONSISTENT WITH WHAT WE WOULD EXPECT WITH GLOBAL  
2 WARMING.

3 MORE RECENTLY, WE THINK WE'VE GOT A GOOD  
4 HANDLE NOW ON SEA LEVEL DURING THE EVENT, SOMETHING  
5 LIKE A 10-TO-15-METER RISE IN SEA LEVEL, 3 TO 5  
6 METERS OF WHICH WE CAN DEFINITELY PUT ON THERMAL  
7 EXPANSION, AND THEN THE REST OF IT, THERE MUST HAVE  
8 BEEN SMALL ICE SHEETS ON ANTARCTICA. IT'S NOT  
9 INCONCEIVABLE THAT THERE WERE SMALL LANDLOCKED ICE  
10 SHEETS ON ANTARCTICA, AND IT'S LIKELY THAT THEY  
11 CONTRIBUTED TO THE REST OF THIS SEA LEVEL RISE.

12 OKAY. SO THAT'S ALL I'M GOING TO SAY ABOUT  
13 THE SORT OF CLIMATE ENVIRONMENTAL CONSEQUENCES AND  
14 START TO MOVE ON TO DISCUSSION ABOUT THE CARBON AND  
15 THE SOURCE OF THIS CARBON.

16 YOU'VE ALL HEARD ABOUT THE POTENTIAL  
17 DECOMPOSITION OF METHANE HYDRATE. THIS IS ONE IDEA  
18 THAT THERE MIGHT HAVE BEEN SOME SORT OF CATASTROPHIC  
19 EVENT THAT ALLOWED METHANE HYDRATES TO DECOMPOSE AND  
20 ADD SEVERAL THOUSAND GIGATONS OF CARBON TO THE OCEAN.

21 THERE ARE OTHER IDEAS. WE KNOW THAT AROUND  
22 THIS TIME THE NORTH ATLANTIC WAS STARTING TO OPEN,  
23 RIFTING BETWEEN GREENLAND AND EUROPE, AND THERE'S  
24 BEEN SOME SUGGESTIONS THAT THIS MIGHT HAVE INVOLVED  
25 SOME THERMAL DECOMPOSITION OF ORGANIC MATTER AND

0458  
1 SEDIMENTS IN THAT PART OF THE WORLD.

2 AND THEN ANOTHER IDEA IS THAT THERE IS  
3 EVIDENCE TO SUGGEST WE HAVE DESICCATION OXIDATION OF  
4 ORGANIC-RICH SOILS OR EVEN LAKES LARGER THAN SEAS.

5 THE MAIN THING ABOUT ALL THREE OF THESE,  
6 THAT THESE ARE REALLY THE ONLY POTENTIAL SOURCES FOR  
7 THE AMOUNT OF CARBON THAT WE'RE TALKING ABOUT. WE  
8 NEED THOUSANDS OF GIGATONS OF CARBON TO BE RELEASED  
9 IN A VERY SHORT PERIOD OF TIME, AND THIS IS IT.

10 NOW, I WILL SAY AHEAD OF TIME THAT I THINK  
11 WHAT HAPPENED IS THAT THE EVENT MIGHT HAVE BEEN  
12 TRIGGERED BY THIS, BUT THAT THESE TWO SOURCES OF  
13 CARBON CAME IN LATER ON AS FEEDBACKS, STARTING TO --  
14 ESSENTIALLY, DOUBLING UP OR TRIPLING THE AMOUNT OF  
15 CARBON THAT CAME INTO THE SYSTEM, AND I WILL EXPLAIN  
16 WHY I BELIEVE THAT IN A SECOND, AT THE END.

17 OKAY. SO WHAT WE HAVE BEEN TRYING TO DO IS  
18 TO FIGURE OUT WHETHER OR NOT WE'RE DEALING WITH  
19 SINGLE OR MULTIPLE SOURCES OF CARBON. THAT MEANS TO  
20 ME THAT WE NEED TO BE ABLE TO SAY SOMETHING ABOUT THE  
21 RATE OF RELEASE, HOW FAST WAS THIS EXCURSION, HOW  
22 FAST DID THE OCEAN CARBON CHEMISTRY CHANGE. AND  
23 THERE ARE DIFFERENT STRATEGIES THAT WE'RE USING TO  
24 TRY TO GET AT THIS. IT IS NOT THAT EASY TO DO, IN  
25 PART BECAUSE WE'RE DEALING WITH RECORDS THAT HAVE

0459  
1 BEEN TRUNCATED BY DISSOLUTION OCEAN ACIDIFICATION.

2 AND THEN IF YOU KNOW SOMETHING ABOUT THE RATE, THEN  
3 YOU JUST HAVE TO GET A HANDLE ON THE MASS OF CARBON  
4 RELEASED. WE HAVE THE CARBON ISOTOPE EXCURSION; BUT  
5 IF WE START TO ASSUME THAT THERE ARE MULTIPLE SOURCES  
6 OF CARBON ISOTOPE EXCURSION, IT DOESN'T HELP US  
7 CONSTRAIN THE SOURCE. WHAT WE CAN USE ARE CHANGES IN  
8 OCEAN CHEMISTRY. CARBONATE SATURATION STAYING IN THE  
9 OCEAN MAY ALLOW US TO CONSTRAIN THAT. SO WE ARE  
10 DEALING WITH MODELS, ERROR PROCESS BETWEEN MODELS, TO  
11 TRY TO CONSTRAIN THE MASS OF CARBON.

12 NOW, WHAT I'M GOING TO JUST SHOW YOU ARE  
13 SOME RECORDS THAT WE'VE DEVELOPED THAT SUGGEST THAT  
14 THE CARBON ISOTOPE EXCURSION ITSELF WAS FAIRLY RAPID.  
15 THERE ARE LOTS OF RECORDS THAT HAVE BEEN GENERATED;  
16 AND DEPENDING ON THE TYPE OF MATERIAL YOU MEASURE AND  
17 WHERE YOU GENERATE THESE RECORDS, THE PATTERN OF THE  
18 EXCURSION ALWAYS LOOKS DIFFERENT. AND SO WHAT WE  
19 TRIED TO DO WAS SORT IT OUT AND GET IT RIGHT AND USE  
20 STRATEGIES THAT WE THOUGHT WOULD BEST REPRESENT THE  
21 ACTUAL CHANGES IN THE CARBON ISOTOPIC COMPOSITION OF  
22 SEAWATER.

23 I'M GOING TO SHOW YOU SOME RECORDS FROM TWO  
24 LOCATIONS: ONE RECORD FROM OFF OF ANTARCTICA, A  
25 DEEPSEA RECORD; AND THEN A SHALLOW MARINE RECORD FROM  
0460

1 A NEW JERSEY MARGIN. AND THIS IS FROM AN OLD PAPER.  
2 2002, OKAY, NOT THAT OLD. BUT WHAT WE DID IN THIS  
3 CORE, THIS IS A SEDIMENT CORE, JUST LIKE THE ONE I  
4 SHOWED A PHOTOGRAPH OF EARLIER; AND WE SAMPLED EVERY  
5 CENTIMETER ALONG THIS CORE THROUGH THE EXCURSION  
6 LAYER, THROUGH THE P-E BOUNDARY; AND WHAT WE DID WAS  
7 TO ANALYZE SHELLS OF PLANKTONIC AND BENZOIC  
8 FORAMINIFERA INDIVIDUALLY RATHER THAN FROM EACH  
9 SAMPLE, GROUP 10 SHELLS. WE THOUGHT WE WOULD JUST  
10 ANALYZE A SINGLE SHELL, OR SINGLE SHELLS, AS MANY AS  
11 10 FROM EACH LEVEL. AND THE REASON FOR THAT IS WE  
12 WERE WORRIED ABOUT THE EFFECTS OF MIXING OUR  
13 OBSERVATION ON SMOOTHING THE ISOTOPIC EXCURSION. AND  
14 WHAT WE FOUND WAS PRETTY INTERESTING, WHICH WAS, IF  
15 YOU LOOK AT THE MIXED LAYER OF PLANKTONIC FORAMS, TO  
16 GET PRE-EXCURSION VALUES AND THEN YOU GET TO THIS  
17 HORIZON WHERE ALL OF A SUDDEN YOU HAVE A MIXTURE OF  
18 PRE-EXCURSION AND EXCURSION FORAMS, AND WHAT THIS  
19 WOULD SUGGEST AT FACE VALUE IS SIMPLE; THAT THE DELTA  
20 C-13 OF THE SURFACE OCEAN CHANGED BY 4 PER MIL IN A  
21 GEOLOGIC INSTANT. AND THEN IF YOU LOOK AT FORAMS  
22 THAT GO DEEPER IN THE WATER COLUMN OR BENZOIC FORAMS,  
23 IT LOOKS LIKE THE DELTA C-13 OF THOSE RESERVOIRS  
24 CHANGED LATER; THAT THERE WAS A LAG EFFECT. AND THIS  
25 WOULD ACTUALLY BE CONSISTENT WITH THE IDEA THAT

0461  
1 CARBON WAS RELEASED INTO THE ATMOSPHERE AND THEN WAS  
2 GRADUALLY SWEEPED INTO THE DEEPSEA. AND SO THIS WAS  
3 ONE OF THE INTERPRETATIONS THAT WE PUT FORWARD TO  
4 EXPLAIN THIS RECORD.

5 NOW, THE OTHER THING, OF COURSE, IS THAT  
6 THIS COULD JUST BE AN ARTIFACT. I JUST TOLD YOU THAT

7 THERE IS MASSIVE CARBONATE DISSOLUTION DURING THE  
8 EVENT, AND MAYBE THIS RECORD IS TRUNCATED; AND THAT  
9 IF WE HAD THE FULL RECORD, IT WOULD BE MORE GRADUAL.

10 SO TO GET AROUND THIS, WE DECIDED TO GET  
11 OUT OF THE DEEPSEA AND MOVE UP ONTO THE SHELVES,  
12 WHICH SHOULD HAVE STAYED OVERSATURATED WITH RESPECT  
13 TO CALCIUM CARBONATE; AND THIS IS WHERE WE'VE BEEN  
14 WORKING ON THESE SECTIONS FROM NEW JERSEY AND  
15 ELSEWHERE. THIS IS WHAT THE COASTLINE LOOKED LIKE  
16 55 MILLION YEARS AGO. HERE'S NEW JERSEY. IT'S  
17 MOSTLY UNDERWATER. SOME PEOPLE THINK THAT'S A GOOD  
18 THING.

19 AND WHAT WE LIKE ABOUT THESE SECTIONS IS  
20 THAT THEY'RE REPRESENTING MID-SHELF ENVIRONMENTS.  
21 WATER DEPTH WOULD HAVE BEEN LESS THAN ABOUT  
22 200 METERS AT THIS PARTICULAR SITE. WE HAVE OTHER  
23 SITES FURTHER TOWARDS THE COASTLINE, WHICH WOULD HAVE  
24 BEEN RIGHT ABOUT HERE.

25 THE SECTIONS HAVE MOSTLY SILICICLASTIC

0462

1 SEDIMENT, CLAY-RICH. IN THE EXCURSION LAYER, THE  
2 CLAY IS PREDOMINANTLY KAOLINITE, WHICH IS CLAY THAT  
3 IS PRODUCED MAINLY IN THE TROPICS. AND THIS IS  
4 SOMETHING THAT YOU SEE THROUGHOUT THE MID AND HIGH  
5 LATITUDES, AND EVERY ONE OF THESE COULD BE BOUNDARY  
6 SECTIONS.

7 THERE ARE SOME FORAMINIFERA. THEY'RE VERY  
8 SCARCE, BUT THEY'RE REALLY WELL PRESERVED. AND THEN  
9 THE OTHER THING THAT'S APPEALING ABOUT THESE SECTIONS  
10 IS THAT THE SEDIMENTATION RATES ARE MUCH HIGHER THAN  
11 THE DEEPSEA SECTIONS; AGAIN, BECAUSE OF THE  
12 SILICICLASTIC FLUX. AND THIS GIVES US A HIGHER  
13 FIDELITY RECORD.

14 THOSE ARE SOME PHOTOS OF SOME FORAMS. OVER  
15 HERE IS JUST THE CARBONATE CONTENT, AND THIS IS THE  
16 PERCENT SAND FRACTION. AND THE PATTERNS THAT YOU SEE  
17 IN BOTH OF THESE ARE CONSISTENT WITH RISING SEA  
18 LEVEL. SO THAT'S ALL I'LL SAY ABOUT THOSE RECORDS.

19 THIS IS THE OXYGEN ISOTOPE RECORDS. THIS  
20 IS THE CARBON ISOTOPE RECORD. AND WE FIND THE SAME  
21 MIXED-LAYER FORAMS THAT WE GET IN SOME OF THOSE OCEAN  
22 SECTIONS, BENZOIC FORAMS AND SO FORTH. AND SO  
23 THERE'S THE EXCURSION, AND IT'S FAIRLY ABRUPT, OKAY.

24 WHAT WE DID WAS TO DO THE SAME THING THAT  
25 WE DID WITH THE PELAGIC SECTIONS, WHICH IS TO LOOK AT

0463

1 INDIVIDUAL SHELLS OF FORAMINIFERA, AND SO THIS IS THE  
2 CARBON ISOTOPE RECORD. THESE ARE MIXED-LAYER FORAMS,  
3 THESE ARE BENEDICTS (PHONETIC), THESE ARE THERMOCLINE  
4 FORAMS. AND ALREADY YOU CAN SEE THAT THE PATTERN IS  
5 SOMEWHAT SIMILAR TO WHAT WE SEE IN THE DEEPSEA; THAT  
6 IS, THAT WE GET A SUDDEN JUMP FROM THESE  
7 PRE-EXCURSION VALUES TO EXCURSION VALUES. AND SO THE  
8 QUESTION IS -- SO, YOU KNOW, IS THIS TELLING US THAT  
9 THE CARBON ISOTOPE EXCURSION IS GLOBALLY RAPID?

10 I THINK SO. AND WHAT I DID WAS I JUST TOOK  
11 THE VALUES FROM THIS PARTICULAR SITE, AND I PLOTTED



12 THEM, ALONG WITH THE VALUES OF MIXED-LAYER FORAMS  
13 FROM THE SOUTHERN OCEAN SITES, AND THE IDEA BEING  
14 THAT THE CARBON ISOTOPIC COMPOSITION, THE CHANGES  
15 PRE-EXCURSION TO EXCURSION SHOULD HAVE BEEN ABOUT THE  
16 SAME, THE ABSOLUTE RATIOS PLUS THE MAGNITUDE OF THE  
17 EXCURSION.

18 SO IN THIS PANEL OVER HERE, THESE ARE THE  
19 VALUES FROM THE SITE OFF OF ANTARCTICA. AND THESE  
20 ARE THE BASS RIVER, NEW JERSEY SITE VALUES. AND  
21 BASICALLY -- AND THIS IS A CROSS PLOT OF CARBON  
22 VERSUS OXYGEN. AND YET, WE SEE THE SAME  
23 PRE-EXCURSION VALUES IN CARBON, THE SAME EXCURSION  
24 VALUES; AND VERY FEW OR NO VALUES THAT WE CONSIDER  
25 TRANSITIONAL.

0464

1 SO YOU COULD ARGUE THAT, WELL, MAYBE THERE  
2 WASN'T DISSOLUTION THERE; BUT I THINK WHEN YOU LOOK  
3 AT THESE RECORDS -- AND THERE ARE OTHERS NOW WHERE  
4 WE'RE MEASURING ORGANIC CARBON, CARBON ISOTOPE RATIO.  
5 AND EVERYTHING INDICATES THAT THE CARBON ISOTOPE  
6 EXCURSION WAS FAIRLY ABRUPT. IT WAS A VERY RAPID  
7 EVENT, GEOLOGICALLY RAPID. WHAT DOES THAT MEAN? WAS  
8 IT A HUNDRED YEARS? WAS IT 500 YEARS? OR A FEW  
9 THOUSAND YEARS? AT THIS POINT, YOU KNOW, WE'RE  
10 COMFORTABLE IN SAYING IT WAS PROBABLY LESS THAN FOUR  
11 OR FIVE THOUSAND YEARS.

12 SO, FINALLY, I JUST WANT TO TALK A LITTLE  
13 BIT ABOUT THE MASSIVE CARBON; AND HERE, AS I SAID, WE  
14 HAVE THESE TWO CONSTRAINTS. I WANT TO TALK ABOUT THE  
15 CHANGES IN OCEAN CARBONATE CHEMISTRY. AND I'M JUST  
16 GOING TO TAKE YOU TO ONE RECORD, PROBABLY OUR BEST  
17 ONE, THIS ONE OFF OF SOUTH AFRICA IN THE WALRUS  
18 (PHONETIC) RIDGE, WHERE WE DRILLED A DEPTH,  
19 TRANSECTED FIVE SITES FROM ABOUT 2 AND A HALF  
20 KILOMETERS TO ABOUT 4.8 KILOMETERS WATER DEPTH; AND  
21 AT ALL THE SITES, WHEN WE DRILLED THROUGH THE  
22 BOUNDARY, WE ENCOUNTERED THIS CLAY LAYER, OKAY, SO  
23 THIS IS CLEARLY INDICATING THAT THE OCEAN BECAME --  
24 THE DEEPSEA BECAME UNDERSATURATED WITH RESPECT TO  
25 CALCIUM CARBONATE DURING THE PETM.

0465

1 WE SEE THIS CLAY LAYER, EVEN AT THE  
2 SHALLOWEST SITE. AND SO THIS PATTERN FROM THE  
3 SHALLOWEST SITE TO THE DEEPEST SITE SIMPLY SUGGESTS  
4 THAT THE OCEAN CARBONATE COMPOSITION THAT'S SHOWN  
5 VERY QUICKLY AND THEN GRADUALLY DESCENDED. AND THIS  
6 TOOK SOMETHING LIKE -- THE INITIAL SHOALING, WE DON'T  
7 KNOW EXACTLY HOW FAST THAT OCCURRED BECAUSE BASICALLY  
8 WE'RE DISSOLVING AWAY SEDIMENTS AS WE GO. BUT THE  
9 RECOVERY TAKES SOMETHING LIKE 50,000 YEARS, WHICH IS  
10 FAIRLY CONSISTENT WITH WHAT YOU WOULD EXPECT IF YOU  
11 WERE TO SUDDENLY DUMP IN SAY SEVERAL THOUSAND  
12 GIGATONS OF CARBON INTO THE OCEAN AND ACIDIFY IT.  
13 THIS IS ABOUT HOW LONG IT WOULD TAKE FOR THE OCEAN TO  
14 RESTORE SOME REASONABLE SATURATION FOR CARBONATE TO  
15 ACCUMULATE.

16 OVER HERE IS THE BULK CARBONATE CARBON

17 ISOTOPE RECORD. THE ONE THING I WILL SAY ABOUT THIS  
18 IS SIMPLE; THAT ALL THESE SITES THE RECORD LOOKS  
19 DIFFERENT, AND I THINK THIS PATTERN IS PURELY A  
20 FUNCTION OF -- IT'S AN ARTIFACT OF DISSOLUTION AND  
21 REWORKING. THE CARBON ISOTOPE EXCURSION, AS I ARGUED  
22 EARLIER, IS PROBABLY VERY ABRUPT. AND SO THIS  
23 PATTERN WHICH HAS BEEN PICKED UP IN A LOT OF DEEPSEA  
24 SECTIONS IS AN ARTIFACT. THE SIGNIFICANCE OF THAT  
25 WILL BECOME OBVIOUS IN A SECOND.

0466

1 WHAT DOES THIS MEAN? THE CARBONATE  
2 DISSOLUTION PATTERNS ARE GLOBAL, AND WE'VE MODELED  
3 THIS IN SEVERAL DIFFERENT WAYS. WE'RE USING THE 3D  
4 OCEAN MODELS AVAILABLE, THE GE MODEL, ALSO THE  
5 HAMBURG MODEL, AND THEN WE'RE USING BOX MODELS, AS  
6 WELL. AND THE BOTTOM LINE IS THAT FOR A VERY  
7 SUSTAINED PERIOD OF TIME, SEVERAL THOUSANDS YEARS,  
8 THAT THE DEEPSEA HAD TO HAVE BEEN SIGNIFICANTLY  
9 UNDERSATURATED; AND TO DO THIS, YOU WOULD NEED AT  
10 LEAST THREE TO FOUR THOUSAND GIGATONS OF CARBON.  
11 BETWEEN ALL THE MODELED ONES THAT WE HAVE DONE, THAT  
12 SEEMS TO BE THE MINIMUM TO SUSTAIN THIS DEGREE OF  
13 UNDERSATURATION.

14 SO COMING BACK TO THIS RECORD AND AGAIN  
15 LOOKING AT THE CARBON ISOTOPE RECORD HERE, WE HAVE  
16 STARTED TO TRY TO USE THE MODELS IN COMBINATION WITH  
17 THE CARBON ISOTOPES TO COME UP WITH SORT OF A UNIQUE  
18 SET OF CARBON FLUXES IN TERMS OF MASS AND ISOTOPIC  
19 COMPOSITION; AND I TOLD YOU EARLIER THAT IF YOU  
20 PLOTTED THE BULK CARBON ISOTOPE RECORDS, CARBONATE  
21 CARBON ISOTOPE RECORDS FROM VARIOUS SITES, THE  
22 PATTERN YOU WOULD GET WOULD LOOK LIKE THIS. IT WOULD  
23 SUGGEST THAT THERE ARE MULTIPLE STEPS IN THE RECORD,  
24 AND IT WAS GRADUAL IN THE EXCURSION, AND THAT IT TOOK  
25 SOMETHING LIKE 50,000 YEARS.

0467

1 SO IF YOU TRY TO SIMULATE THIS WITH A BOX  
2 MODEL, YOU COULD DO SOMETHING LIKE THIS. YOU COULD  
3 TAKE CARBON, AND IN THIS, WE'VE RUN A NUMBER OF  
4 SIMULATIONS. THE BOTTOM LINE IS WE NEED FOR THIS  
5 SOMETHING LIKE 5,000 GIGATONS. WE'RE USING CARBON  
6 THAT HAS A DELTA C-13 THAT'S MIDWAY BETWEEN A SOURCE  
7 OF METHANE AND SAY ORGANIC CARBON; AND YOU CAN PULSE  
8 THE CARBON TWICE LIKE THIS AND THEN LEAK IT FOR  
9 AWHILE. AND THIS IS SORT OF THE CARBON ISOTOPE  
10 PATTERN THAT YOU PRODUCE FOR THE DIFFERENT OCEAN  
11 BASINS. AND SO THIS IS SOMETHING THAT WOULD FIT  
12 THESE BULK CARBONATE CARBON ISOTOPE RECORDS. BUT I  
13 JUST GOT DONE TELLING YOU THAT WE THINK THOSE RECORDS  
14 ARE WRONG; THAT THE ACTUAL ONSET OF THE EXCURSION WAS  
15 FAIRLY ABRUPT.

16 AND THIS IS ANOTHER RECORD FROM BASS RIVER.  
17 I TOLD YOU WE HAVE -- I SHOWED YOU THE FORAM, THE  
18 CARBONATE RECORDS; BUT WE ALSO HAVE ORGANIC CARBON  
19 RECORDS, AND THEY SHOW THE SAME THING. THEY WOULDN'T  
20 BE AFFECTED BY THE DISSOLUTION, THESE SILICICLASTIC  
21 SEDIMENTS; BUT AGAIN, WITH THE ORGANIC CARBON, WE

22 ALSO GET A VERY ABRUPT CARBON ISOTOPE EXCURSION.  
23 SO IF YOU MODEL -- IF YOU WANTED TO  
24 SIMULATE THAT, THEN YOU CAN -- IN THIS PARTICULAR  
25 SIMULATION, WE'RE USING 3,600 GIGATONS OF CARBON WITH

0468

1 A DELTA C-13 CLOSER TO A METHANE SOURCE. YOU PULSE  
2 IN 2,200 GIGATONS AND THEN LEAK IN THE REST FOR  
3 SEVERAL TENS OF THOUSANDS OF YEARS, AND THAT PRODUCES  
4 AN EXCURSION THAT LOOKS MORE LIKE WHAT WE SEE IN THE  
5 PLANKTON AND THE BULK CARBON ISOTOPE RECORDS.

6 SO WE'RE AT THE POINT NOW WHERE WE'RE  
7 RUNNING THESE SORTS OF SIMULATIONS USING THE BOX  
8 MODELS, BUT ALSO THESE 3D OCEAN MODELS, BECAUSE THE  
9 PATTERN OF CARBONATE DISSOLUTION SEEMS TO VARY FROM  
10 BASIN TO BASIN.

11 THE ONE THING THAT I WANTED TO DO HERE WAS  
12 START TO TALK ABOUT WHAT THIS MEANS IN TERMS OF THE  
13 OCEAN SATURATION STATE, ACIDIFICATION IF WE JUST  
14 ASSUME THAT THIS PARTICULAR SET OF EXPERIMENTS BEST  
15 REPRESENTS WHAT WE SEE IN THE DEEPSEA RECORD. WE'RE  
16 ACTUALLY ADDING THIS 2,200 GIGATONS OF CARBON OVER  
17 SEVERAL THOUSAND YEARS HERE, AS YOU CAN SEE. SO IT'S  
18 NOT QUITE AT ANTHROPOGENIC EMISSION RATES, BUT IT IS  
19 SORT OF UP THERE. AND IN DOING SO, IF YOU START TO  
20 LOOK AT THE EFFECTS THAT THAT HAS ON THE SATURATION  
21 STATE OF THE SURFACE OCEANS, AND WE'RE JUST FOCUSING  
22 ON SURFACE OCEANS, YOU CAN SEE WHAT HAPPENS HERE.  
23 YEAH, THERE IS A DROP, BUT WE DON'T GET TO A POINT  
24 WHERE THE OCEAN'S UNDERSATURATED. AND AGAIN, THIS IS  
25 SOMETHING THAT WE'VE ALWAYS BELIEVED. WE HAD THE

0469

1 SUSPICION THAT THE RATE OF RELEASE WAS SLOW ENOUGH  
2 THAT YOU DON'T GET TO UNDERSATURATION IN SURFACE  
3 OCEAN.

4 NOW, FOR COMPARISON, YOU CAN JUST TAKE THE  
5 MODERN OR THE PREDICTED OR PROJECTED ANTHROPOGENIC  
6 FLUX, BUSINESS AS USUAL, AND THEN CALCULATE THE  
7 CHANGE IN THE SURFACE SATURATION STATE. AND, OF  
8 COURSE, IT IS MUCH MORE SEVERE HERE. AND DESPITE THE  
9 FACT THAT IN TERMS OF MASS, OKAY, IT'S DOUBLE, BUT  
10 STILL -- AND, YOU KNOW, THE MAIN POINT HERE IS  
11 SIMPLE, WHICH IS THAT THE RATE OF RELEASE IS REALLY  
12 IMPORTANT. THE BUFFERING CAPACITY OF THE OCEAN IS IN  
13 THE DEEPSEA; IT'S NOT IN THE SURFACE OCEAN. AND IT  
14 TAKES SOMETHING LIKE 500 YEARS OF THE OCEAN TO TURN  
15 OVER. SO IN THE CASE OF THE PETM, WHERE THE CARBON  
16 IS BEING ADDED OVER SEVERAL OCEAN MIXING CYCLES, THE  
17 OCEAN IS ABLE TO BUFFER TO SOME EXTENT THE CHANGES IN  
18 PH WITH DISSOLUTION OF CARBONATE ON THE SEAFLOOR;  
19 WHEREAS, IN THE PRESENT-DAY SITUATION, THAT'S NOT  
20 GOING TO HAPPEN. THE RATE OF RELEASE IS MUCH FASTER  
21 THAN THE TURNOVER TIME OF THE OCEAN. SO THE SURFACE  
22 OCEAN ENDS UP BEARING THE BRUNT OF THE CHANGES IN PH  
23 AND ALSO IN TERMS OF ABSORBING OR TRYING TO ABSORB  
24 THAT CARBON. OF COURSE, THAT'S WHY THERE'S THIS  
25 CONCERN ABOUT THE EFFECT OF PH CHANGES ON CALCIFIERS.

0470

1           NOW, IT'S USUALLY AT THIS POINT IN THE TALK  
2 OVER THE YEARS WHERE PEOPLE WILL SAY, WELL, WHAT  
3 HAPPENED TO CALCIFIERS? WHAT HAPPENED TO CORALS, FOR  
4 INSTANCE, DURING THE PETM? AND UP UNTIL ABOUT A  
5 MONTH OR TWO AGO, THE ONLY THING I COULD SAY IS THAT  
6 WE DON'T KNOW; THAT WE KNOW FROM STUDIES THAT  
7 PALEONTOLOGISTS HAVE DONE OVER DECADES THAT THE  
8 DIVERSITY OF CORALS IN THE EARLY EOCENE IS LOWER THAN  
9 THE DIVERSITY OF CORALS IN THE LATE PALEOCENE PRIOR  
10 TO THE EVENT, BUT NOBODY KNEW WHY OR WHEN THAT  
11 HAPPENED. MORE RECENTLY, A PAPER WAS PUBLISHED WHERE  
12 TWO PALEONTOLOGISTS LOOKED AT DATA FROM A NUMBER OF  
13 SECTIONS SURROUNDING THE TETHYS; AND WHAT THEY HAD TO  
14 DO WASN'T EASY. THEY HAD TO THINK ABOUT -- WELL,  
15 THERE WAS STRATIGRAPHIC ISSUES; BUT ULTIMATELY WHAT  
16 THEY WERE ABLE TO DEMONSTRATE OR CLAIM WAS THIS:  
17 THAT THIS REPRESENTS CORAL DIVERSITY FROM THE  
18 PALEOCENE INTO EARLY EOCENE. AND PRIOR TO THIS STUDY  
19 WE WOULD SAY THAT DURING THIS WHOLE EARLY EOCENE  
20 DIVERSITY WAS LOW COMPARED TO THE PALEOCENE.

21           WHAT THEY'RE SAYING IN THIS STUDY IS THAT  
22 THAT CHANGE IN DIVERSITY OCCURRED VERY ABRUPTLY AND  
23 IT OCCURRED RIGHT AROUND THE PALEOCENE AND EOCENE  
24 BOUNDARY, DURING THE PETM. AND THERE IS STILL A LOT  
25 MORE WORK TO BE DONE TO SORT OF RESOLVE THE SCALE OF

0471

1 THIS CHANGE AND THE TIMING, BUT THE POINT IS SIMPLE;  
2 THAT THERE IS THIS CHANGE IN DIVERSITY. THERE IS  
3 ALSO A DROP IN THE DIVERSITY OF LARGER FORAMINIFERA,  
4 WHICH WOULD HAVE INHABITED THIS SHALLOW MARINE  
5 CARBONATE SHELF ENVIRONMENTS. AND SO IT LOOKS LIKE  
6 THAT THESE CHANGES IN DIVERSITY ARE OCCURRING RIGHT  
7 AROUND THE P-E BOUNDARY.

8           NOW, COULD YOU ASK THEN: WELL, DO YOU  
9 THINK IT HAS TO DO WITH THE CHANGE IN THE SATURATION  
10 STATE? OR MAYBE IT'S TEMPERATURE? WELL, IT'S  
11 PROBABLY BOTH. YOU CAN PICK YOUR POISON, THE  
12 COMBINATION OF THE TWO, WARMING.

13           WE BELIEVE THAT TROPICAL TEMPERATURES GOT  
14 UP TO AS HIGH AS 40 DEGREES CENTIGRADE DURING THE  
15 PEAK OF THE PETM; AND I SAY WE BELIEVE BECAUSE IN ALL  
16 THE SHALLOW MARINE -- OR IN ALL THE MARINE SECTIONS,  
17 WHEN WE GO TO LOOK FOR THE MIXED-LAYER FORAMS THAT WE  
18 NORMALLY RECONSTRUCT TEMPERATURES WITH, THEY'RE  
19 ABSENT IN THE TROPICAL SECTIONS DURING THE PETM. AND  
20 IT FINALLY DAWNED ON US WHY THEY'RE GONE. BECAUSE  
21 IT'S TOO WARM, AND THEY JUST LEAVE. AND FROM WHAT I  
22 UNDERSTAND, ON LAND, THE SAME THING IS HAPPENING;  
23 THAT DURING THE PEAK OF THE PETM, PLANT DIVERSITY  
24 BASICALLY GOES TO ALMOST NOTHING; THAT THERE IS A  
25 HUGE DROP IN BIOMASS AND DIVERSITY DURING THE PEAK OF

0472

1 THE PETM, SO THE TROPICS OVERHEAT, AND SO CORALS  
2 WOULD PROBABLY RESPOND TO THAT, AS WELL AS CHANGES IN  
3 PH.

4           SO, HERE'S MY SUMMARY: THE MAGNITUDE OF  
5 THE CIE, WE THINK IT IS FAIRLY RAPID. WE'RE CLOSING

6 IN ON A NUMBER OF SOMETHING LIKE 4,500 GIGATONS OF  
7 CARBON RELEASE DURING THE PETM. AND WHETHER IT IS  
8 4,500 OR 5,000 OR 6,000, AT THE MOMENT, IT DOESN'T  
9 MATTER. BUT WHAT'S IMPORTANT ABOUT A NUMBER THIS  
10 LARGE IS THAT IT ALMOST CERTAINLY REQUIRES MULTIPLE  
11 SOURCES OF CARBON; THAT WE CAN'T SAY THAT ANY ONE OF  
12 THESE SOURCES ALONE COULD PROVIDE THIS MUCH CARBON,  
13 WHICH IS WORRISOME BECAUSE THEN IT MEANS THAT, YOU  
14 KNOW, IF THIS WAS ONE SOURCE AND ONE OR BOTH OF THESE  
15 WERE INVOLVED, THAT THESE ARE PROBABLY COMING IN IN  
16 FEEDBACK MODE, CONSISTENT WITH THE IDEA THAT WHAT'S  
17 HAPPENING AT THE PETM IS A THRESHOLD EVENT; THAT THE  
18 SYSTEM HAS CROSSED OVER SOME SORT OF THRESHOLD, AND  
19 IT STARTS TO REALLY -- WE START TO LOSE CARBON FROM A  
20 LOT OF THESE RESERVOIRS THAT HAVE BEEN ACCUMULATING  
21 CARBON FOR A LONG TIME. OKAY.

22 SO WHAT ARE THE IMPLICATIONS? WELL,  
23 OBVIOUSLY, UNABATED CO2 EMISSIONS WILL LEAD TO A  
24 SEVERE DROP IN PH OF THE SURFACE OCEAN, AND WE ALL  
25 KNOW THIS. THAT'S NOT A SURPRISE.

0473

1 THE QUESTION IS WHETHER WE HAVE TO WORRY  
2 ABOUT POSITIVE FEEDBACKS. AND I THINK THIS IS A  
3 CONCERN. I THINK IT IS SOMETHING THAT THERE SHOULD  
4 BE A LOT OF AND OBVIOUSLY THERE IS A LOT OF EFFORT  
5 TRYING TO UNDERSTAND WHAT THOSE FEEDBACKS WILL BE AND  
6 HOW SIGNIFICANT THEY WILL BE. AND THESE ARE THE ONES  
7 THAT WE'RE LOOKING AT THE MOMENT.

8 AND THEN AS FAR AS WILL METHANE HYDRATES  
9 DISSOCIATE, WILL THEY BECOME PART OF THIS PROCESS OF  
10 RISING CO2? AND, REALLY, IT DOES DEPEND ON THE  
11 MAGNITUDE OF WARMING. THE OCEANS ARE COLD. YOU HAVE  
12 TO PROPAGATE THE HEAT INTO THE OCEAN, DEEPER INTO THE  
13 OCEAN, AND INTO SEDIMENTS; AND THAT TAKES TIME, AND  
14 WE'RE TALKING ABOUT A FEEDBACK THAT COULD COME IN IN  
15 CENTURIES AND MAYBE MILLENNIA.

16 NOW, A LOT OF PEOPLE ARE SKEPTICAL THAT  
17 HYDRATES, THAT WE HAVE TO WORRY ABOUT HYDRATES; AND  
18 WHAT I'M GOING TO SHOW YOU -- I WAS GOING -- I WAS  
19 NOT GOING TO SHOW THIS, BUT THE PAPER IS GOING TO  
20 COME OUT IN "NATURE" IN A COUPLE OF WEEKS ANYWAY, NOT  
21 THAT IT MEANS ANYTHING. BUT IF WE GO BACK TO BASS  
22 RIVER, WE HAVE MULTIPLE PROXY TEMPERATURE RECORDS,  
23 ONE OF WHICH IS THE TEX86 RECORD HERE, AND THERE IS  
24 SOME CONTROVERSY ABOUT HOW WELL THIS PROXY WORKS.  
25 BUT WE GET THE SAME ABSOLUTE TEMPERATURES THAT WE GET

0474

1 WITH THE OXYGEN ISOTOPES, AND THIS IS AN ORGANIC-BASE  
2 PROXY OF TEMPERATURE. AND IF YOU LOOK AT THE DETAILS  
3 HERE, AND THESE ACCUMULATION RATES ARE VERY HIGH, IN  
4 DETAIL WHAT WE FIND IS THAT THE CARBON ISOTOPE  
5 EXCURSION OCCURS UP AT THIS LEVEL, THIS HORIZON,  
6 WHERE WE THINK THE WARMING AND AS WELL AS SOME OTHER  
7 ENVIRONMENTAL INDICATORS OF WARMING -- THIS IS A  
8 DINOFLAGELLATE SPECIES THAT BLOOMS GLOBALLY DURING  
9 THE PETM -- THAT THESE SORT OF PRECEDE THE CIE; THAT  
10 THE WARMING INITIATES SEVERAL THOUSAND YEARS BEFORE

11 THE MAIN CARBON ISOTOPE EXCURSION. THIS IS ABOUT THE  
12 ONLY SECTION WHERE WE SHOULD BE ABLE TO SEE THIS  
13 LEAD/LAG RELATIONSHIP; AND ALL THE DEEPSEA SECTIONS,  
14 THAT WOULD BE MERGED BECAUSE OF DISSOLUTION. WE NEED  
15 THESE SORT OF HIGH-FIDELITY SECTIONS.

16 SO YOU COULD ARGUE JUST BASED ON THIS ONE  
17 RECORD -- AND THAT'S WHAT WE DID -- THAT THE WARMING  
18 DOES LEAD THE CIE; THE CIE CARBON ISOTOPE EXCURSION  
19 REPRESENTING ONE OF THOSE POSITIVE FEEDBACKS, MOST  
20 LIKELY, METHANE; THAT AT SOME POINT, METHANE STARTS  
21 TO GO, THE HYDRATES START TO DISSOCIATE, AND THEN  
22 THEY START TO AMPLIFY THE WARMING. SO THAT'S AT  
23 LEAST ONE WAY THAT WE CAN INTERPRET THIS; AND  
24 OBVIOUSLY, WE NEED MORE SECTIONS LIKE THIS TO REALLY  
25 DEMONSTRATE WHETHER OR NOT THIS IS A REAL PATTERN,

0475

1 THIS IS A GLOBAL PATTERN.

2 SO I THINK I'LL END THERE.

3 THANK YOU FOR YOUR PATIENCE.

4