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Introduction

Cloud Radiative Forcing: defined as the effect of cloudiness on the surface radiative budget *NetR* NetR = SWD - SWU + LWD - LWU

Cloud Radiative Forcing (CRF) = NetR-NetR_{clr} where

Data-set used: all the collected data (not only the BSRN quality checked)



$NetR_{clr} = SWD_{clr} - SWU_{clr} + LWD_{clr} - LWU_{clr}$

We need models to represent the four clear sky fluxes. A methodology was applied, based on data acquired from radiometers, relying on accurate cloud screening to determine clear sky condition and subsequently to define semi-empirical clear fluxes.

Methodology

 $CRF = (SWD - SWD_{clr}) - (SWU - SWU_{clr}) + (LWD - LWD_{clr}) - (LWU - LWU_{clr})$

Clear sky sector

LW

SWD_{clr}

SW

we used the well known Long & Ackerman (2000) method:

Here the average results for the SWD pyranometers **g2** (a,b), and for the



An anisotropy is present during the day



A semi-empirical model based on the temperature of the top of inversion was adopted, Busetto et al. (2013)

$$LWD_{clr} = \varepsilon \sigma T_{400}^{4}$$

 $T_{400}(\sim constant with time t)$ is derived from $T_a(t)$, after a LWD cloud screening (Town et al., 2007)

LWU_{clr} $F_{\uparrow,LW}^{clear} = \epsilon_{snow} \sigma T_{skin}^4 + (1 - \epsilon_{snow}) F_{\downarrow,LW}^{clear}$ Comparing T_{skin} and T_q during clear sky condition, a linear relationship was found:

method applied to [Dir+Dif] **g1** (a+,b+)

	a	a_+	b	b_+	N _{tot}
Summer	1400 ± 67	1394 ± 208	1.13 ± 0.04	1.13 ± 0.05	502
Autumn	1427 ± 93	1403 ± 87	1.14 ± 0.04	1.15 ± 0.04	331
Spring	1390 ± 116	1382 ± 115	1.12 ± 0.07	1.14 ± 0.04	215



SWD cloud effect monthly averages





 $=\frac{A+B\cos(\phi-\phi_0)+C\cos(2(\phi-\phi_0))}{A+B+C}$

An empirical model was developed, depending on both SZA and azimuthal angle

 $(D + E\cos(\theta) + F\cos^2(\theta))(a\mu^b)$

SWU cloud effect monthly averages





 $T_{a}(t)$ has been used to simulate the temperature of the inversion layer removing daily and annual behavior. Main advantage: **wv** not required

$LWD_{clr} = \epsilon (T_{400}/T_g)^4 \sigma T_g^4$

 $\epsilon = -1.41 + 0.0077T_{400} - A_y \cos(\omega_y d + \phi_y)$

LWD cloud effect monthly averages



$T_{skin} = 0.01 + 1.04T_{g} (r^{2} = 0.97)$

and assuming ϵ_{snow} =0.97, we defined the LWU_{clr} time series: uncertainties were evaluated ranging within ±5Wm⁻², considering the error propagation in the formula above

LWU cloud effect monthly averages



The cloud effect appears on average ranging between 0 and +7 Wm⁻²

Cloudiness produces a negative effect ranging between -5 and -30 Wm ⁻²		Cloudiness produces a negative effect ranging between -5 and -15 Wm ⁻²		Image: Description of the cloud effect appears on average ranging between +10 and +30 Wm ⁻²			
(Theil –Sen approach: median of all slopes [*** => p<0.001])		$\frac{Mean \pm Std [W/m^2]}{11.87 \pm 3.57}$ 10.21 ± 5.14 10.74 ± 3.29 10.47 ± 3.19 13.87 ± 3.58 16.48 ± 7.20 16.87 ± 5.93 15.45 ± 6.42 17.23 ± 7.01	 CRF 1. Positive CRF values ranging from 10-17 Wm⁻² in the considered period with a significance trend that appears to be higher that expected if compared to global averaged brightening (+2Wm⁻², Wild et al., 2016) 2. The positive trend is similar to that observed for the LWD Cloud Effect, suggesting that this is the main component affecting our CRF result 3. This can likely to be an effect of the degradation of the calibration constant of LWD that affects clear measurements and not overcast measurements (when the signal is typically around 0). 				
2008 2009 2010 2011 2012 2013 2014 year	2015	autumn (SON) winter (DJF) 0.73 [-0.66, 1.66] units/year	4.Seasonal discrepance	cy: the austral summer behaves differently			

5.The effect of blowing snow should be differentiated with respect to that of cloudiness. 6.The cloud screening based on LWD should be validate with respect to Long and Ackerman methodology

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