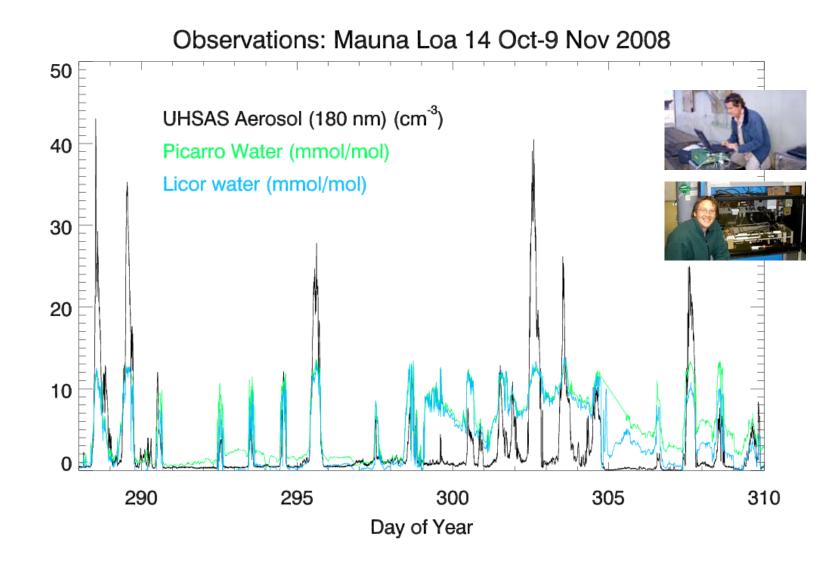
Detection of aerosol growth rate using a mass balance model constrained by water isotope measurements at Mauna Loa

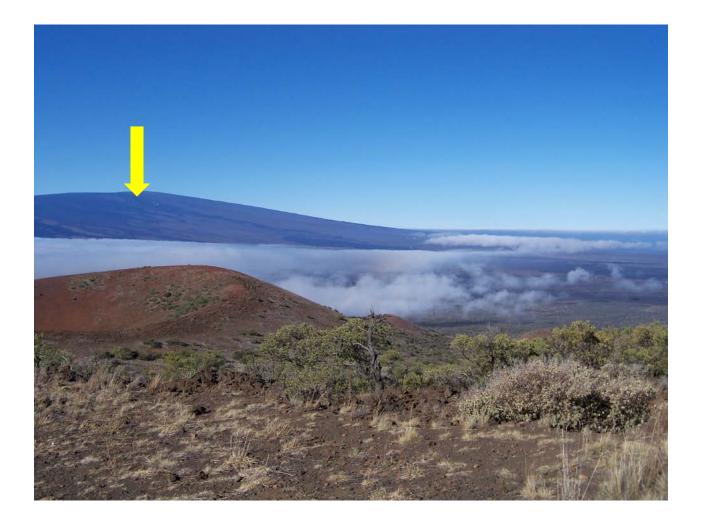
Adriana Bailey^{1,2}, David Noone^{1,2}, Darin Toohey¹ ¹Atmospheric and Oceanic Sciences, University of Colorado at Boulder ²Cooperative Institute for Research in Environmental Sciences



How well does water predict aerosols?



The marine boundary layer rises during the day



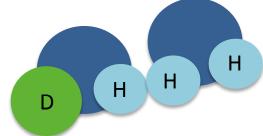
Yes, the cloud is massive

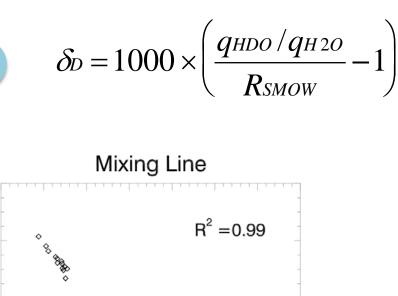


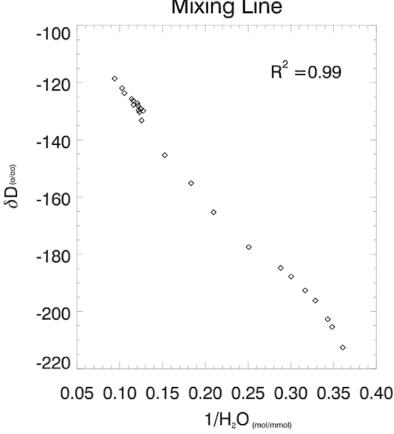
Transport also occurs within the free troposphere



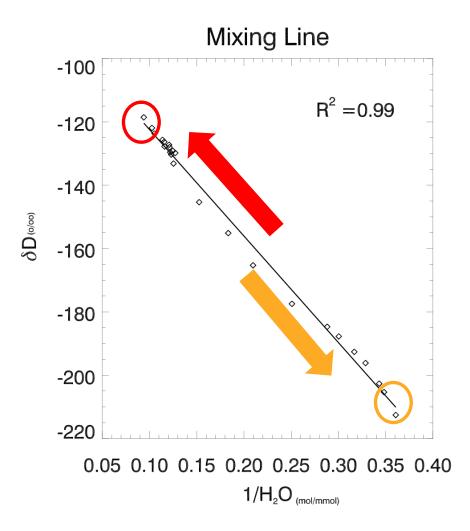
A built-in tracer: stable isotopes







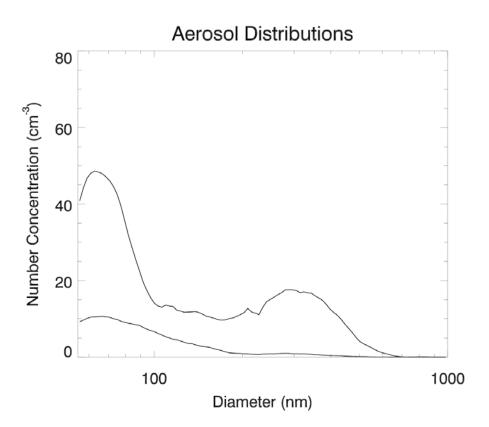
$$q_{next} = q_{end} - (q_{end} - q_{now})e^{-k\Delta t}$$
$$\underline{k} = -\frac{1}{\Delta t} \left(\frac{q_{end} - q_{next}}{q_{end} - q_{now}} \right)$$



$$q_{next} = q_{end} - (q_{end} - q_{now})e^{-k\Delta t}$$
$$k = -\frac{1}{\Delta t} \left(\frac{q_{end} - q_{next}}{q_{end} - q_{now}} \right)$$

No aerosol source/sink:

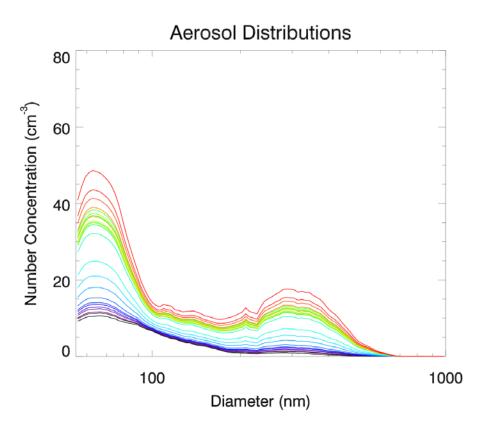
$$a_{next} = a_{end} - (a_{end} - a_{now})e^{-k\Delta t}$$



$$q_{next} = q_{end} - (q_{end} - q_{now})e^{-k\Delta t}$$
$$k = -\frac{1}{\Delta t} \left(\frac{q_{end} - q_{next}}{q_{end} - q_{now}} \right)$$

No aerosol source/sink:

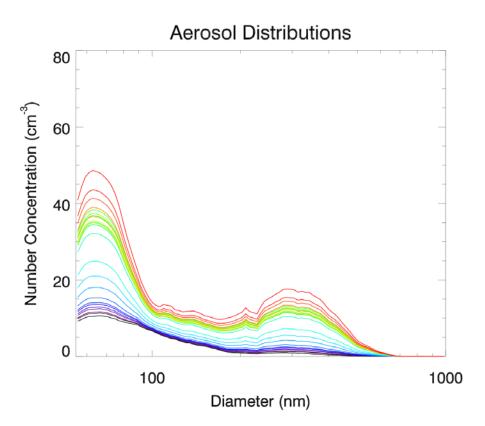
$$a_{next} = a_{end} - (a_{end} - a_{now})e^{-k\Delta t}$$



$$q_{next} = q_{end} - (q_{end} - q_{now})e^{-k\Delta t}$$
$$k = -\frac{1}{\Delta t} \left(\frac{q_{end} - q_{next}}{q_{end} - q_{now}} \right)$$

No aerosol source/sink:

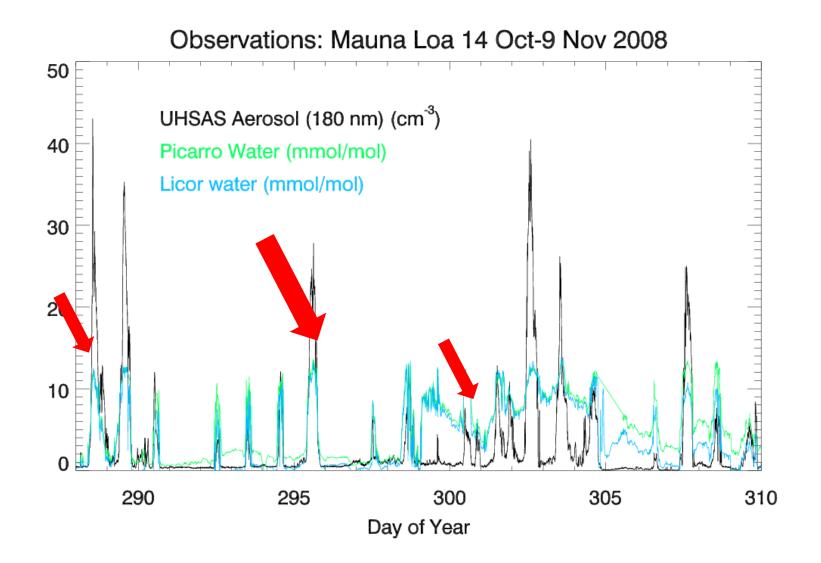
$$a_{next} = a_{end} - (a_{end} - a_{now})e^{-k\Delta t}$$



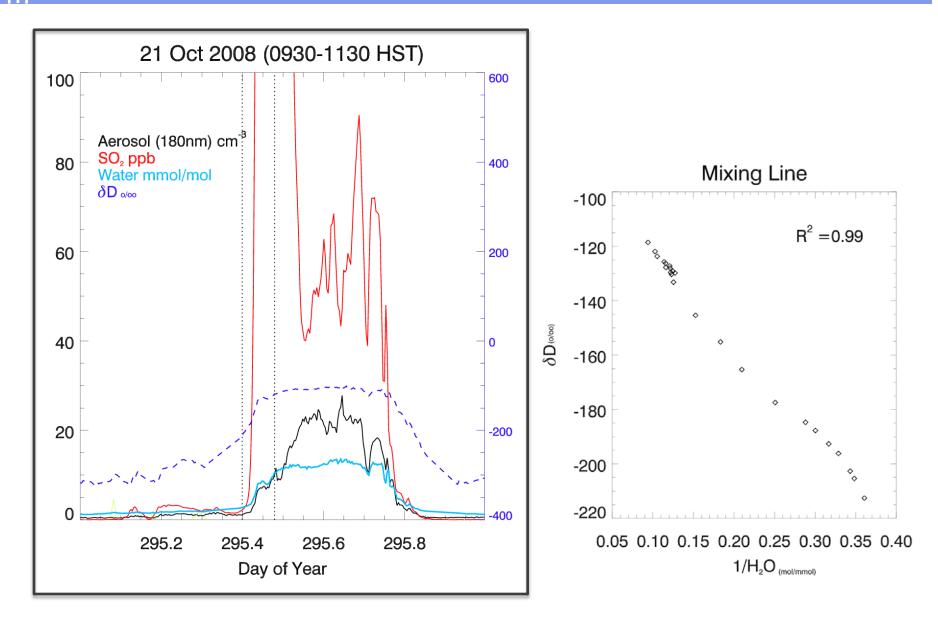
Explicit source term:

$$\frac{da}{dt} = k(a_{end} - a) + source$$

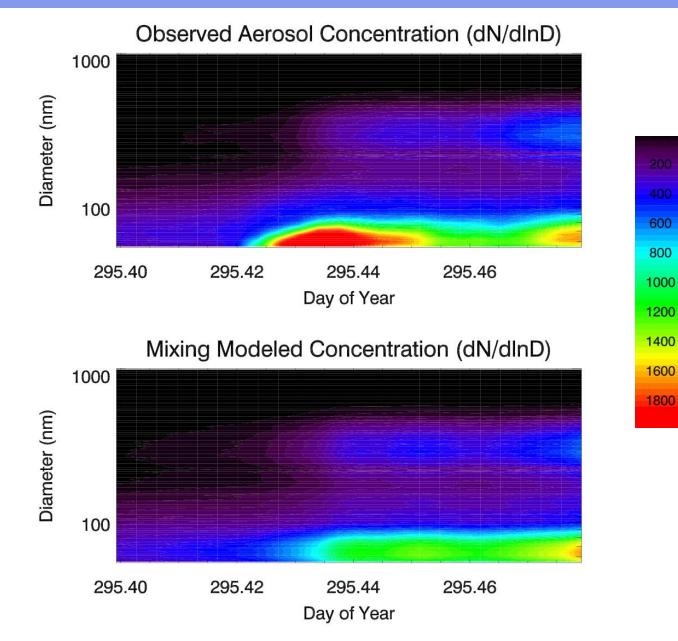
Putting the model to work...



Snapshot: 21 Oct

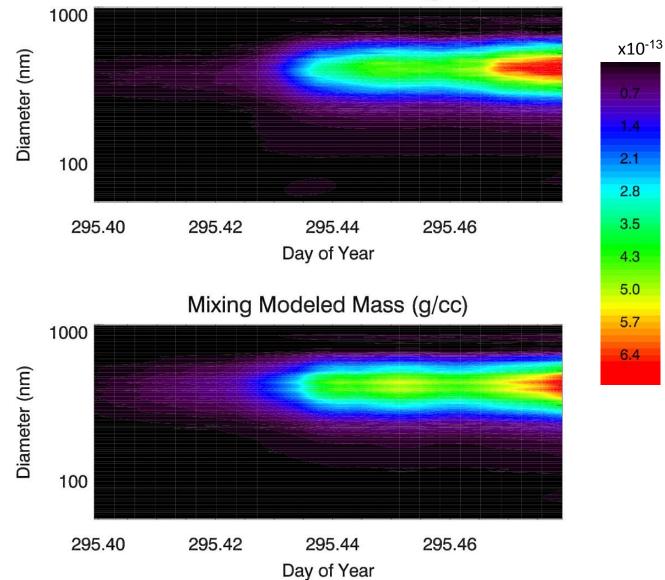


New particles form mid-morning

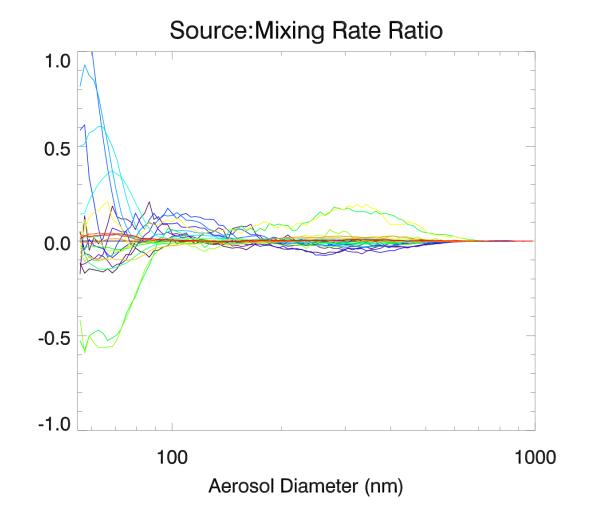


Growth occurs later in the morning

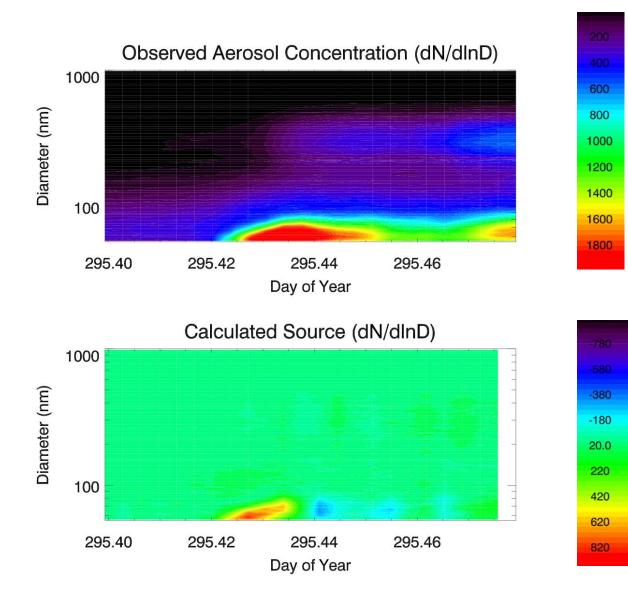
Observed Aerosol Mass (g/cc)



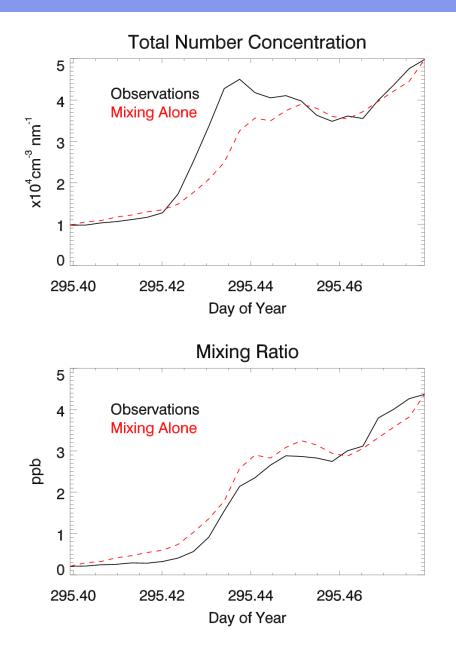
Source importance with time can be diagnosed with S/k



"Source" delineates a nucleation and growth event

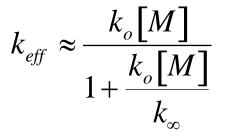


Is the source of the calculated source obvious?



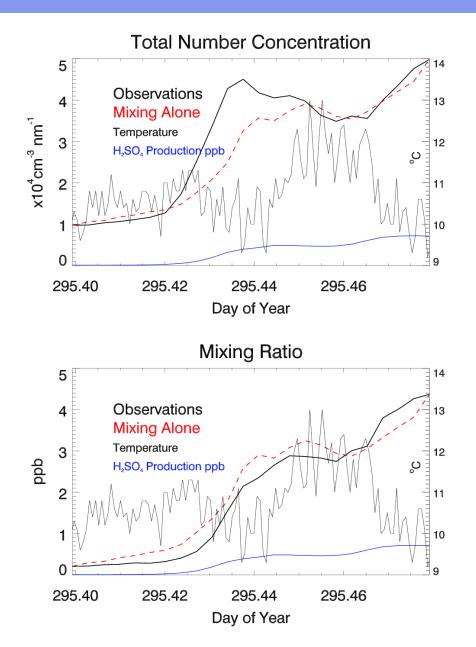
Back of the envelope:

$$\frac{d[H_2SO_4]}{dt} = k_{eff} [OH] [SO_2]$$

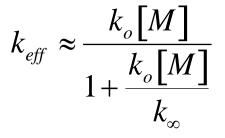


- k_o = low pressure rate constant
- k_{∞} = high pressure rate constant

H₂SO₄ production from SO₂ seems plausible!



Back of the envelope: $\frac{d[H_2SO_4]}{dt} = k_{eff} [OH] [SO_2]$

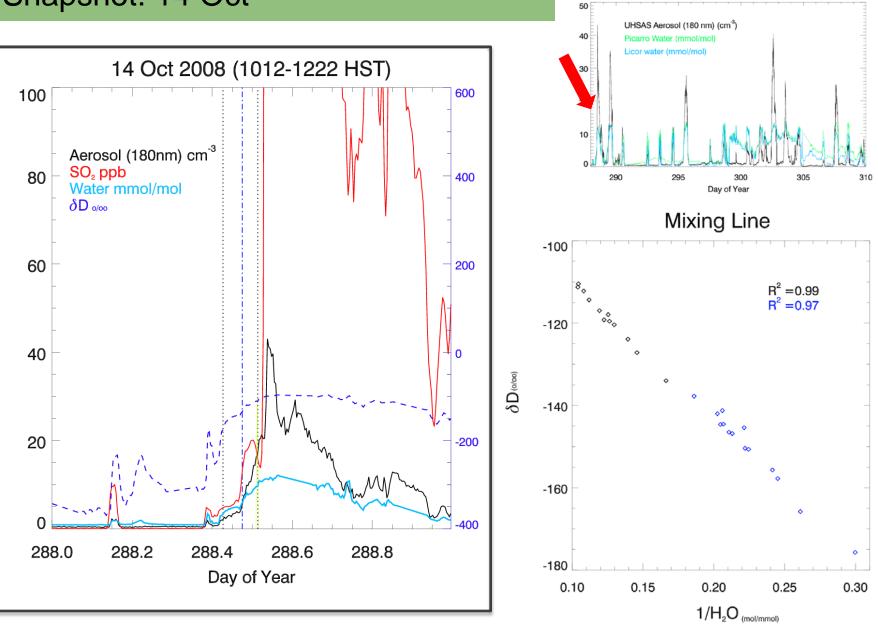


- k_o = low pressure rate constant
- k_∞ = high pressure rate constant

Water's isotopologues define mixing periods and can be combined with aerosol measurements to predict aerosol distribution changes due to air mass mixing and to back out the "true" particle source or sink (independent of mixing)

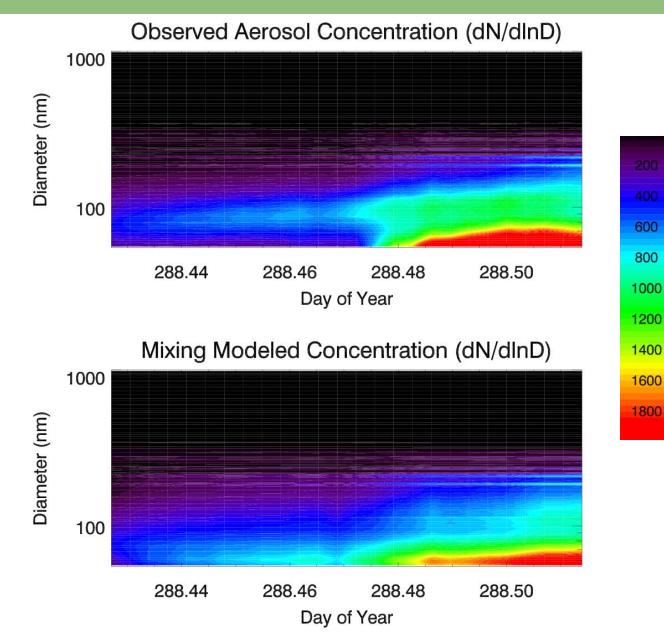
- Analyze additional mixing events + "stagnant" events
- Consider other aerosol sources such as upper troposphere

Snapshot: 14 Oct

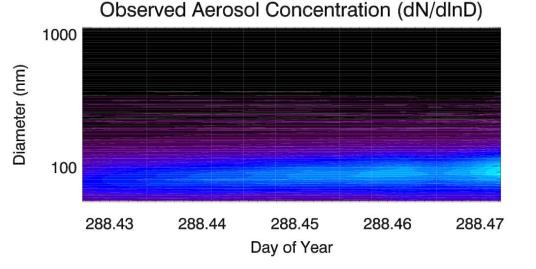


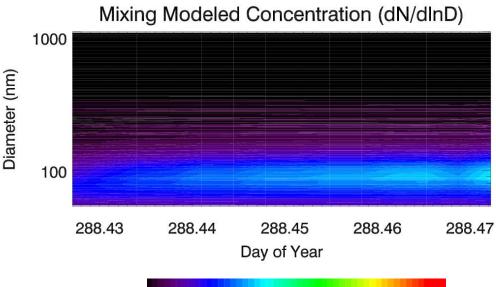
Observations: Mauna Loa 14 Oct-9 Nov 2008

Is there chemistry occurring too?

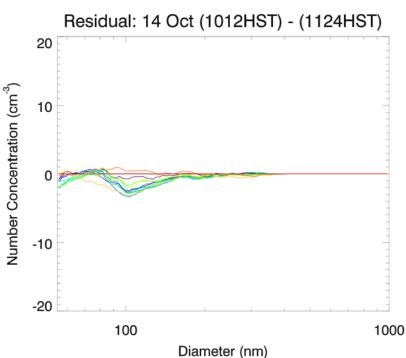


Mixing dominates earlier

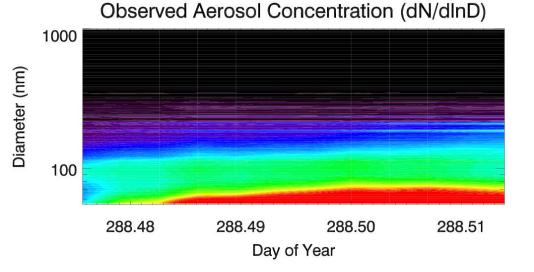


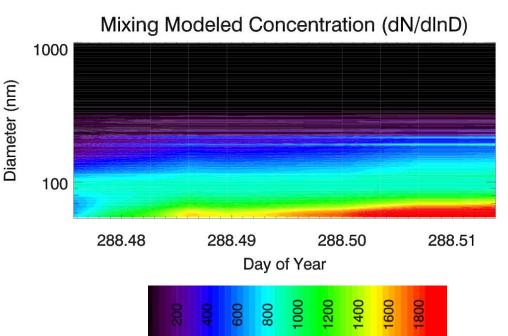


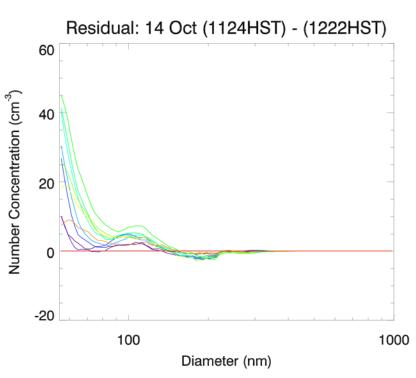




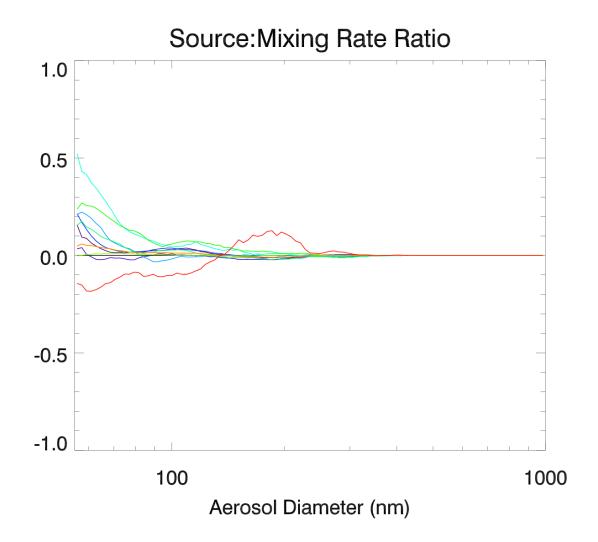
Particle formation appears to occur later on



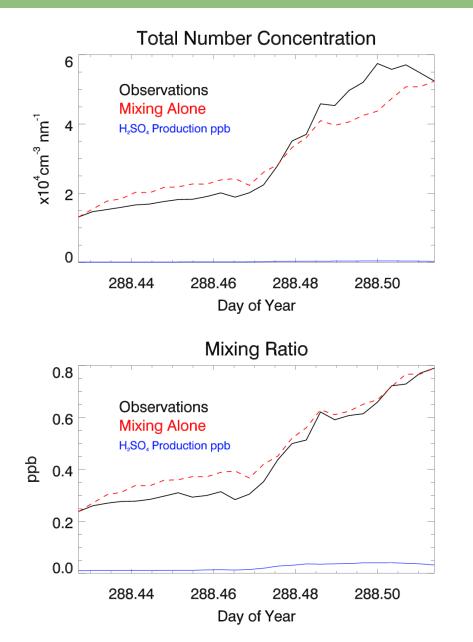




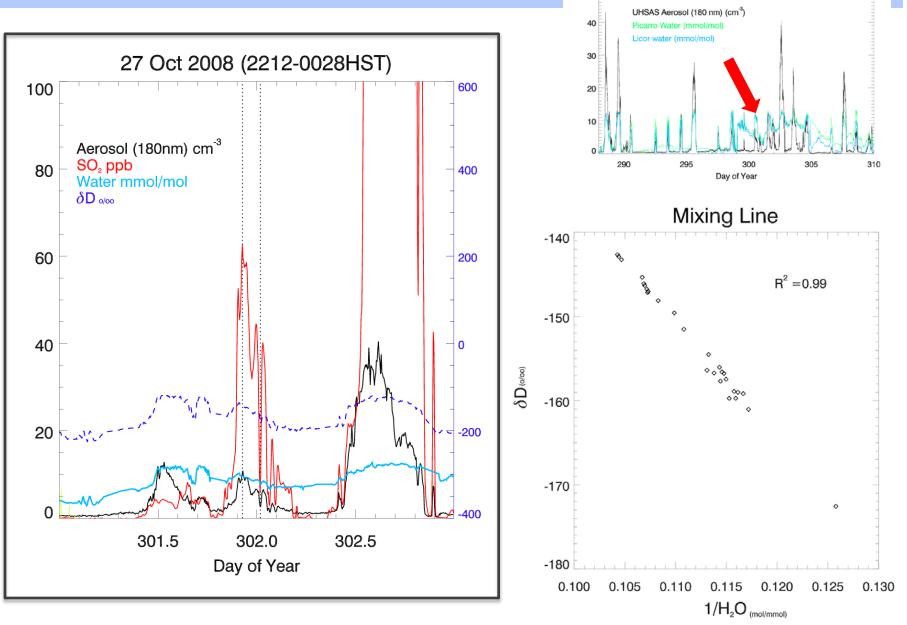
Calculated source important but smaller than expected?



H_2SO_4 may be a plausible source



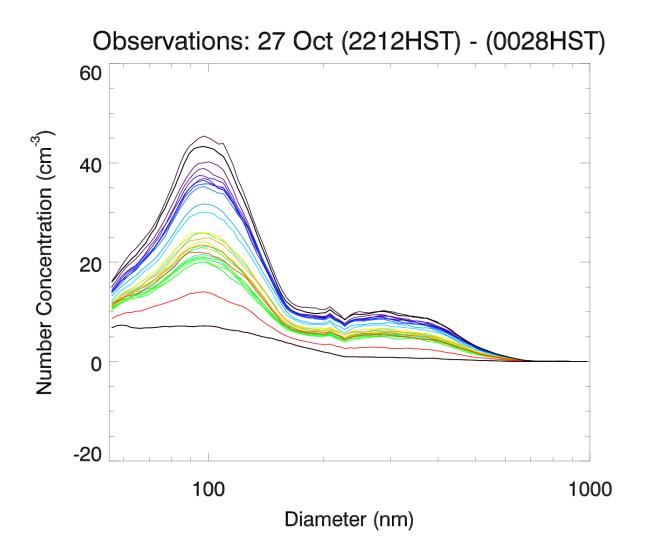
Snapshot: 27 Oct (Night)



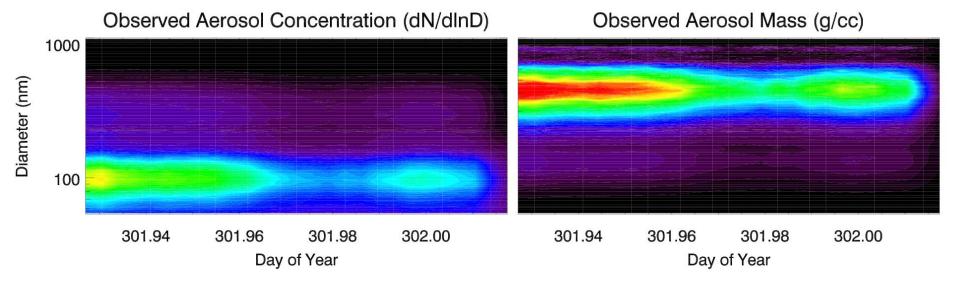
Observations: Mauna Loa 14 Oct-9 Nov 2008

50

Aerosols do not monotonically decline



The mixing model still captures the aerosol change



Mixing Modeled Concentration (dN/dlnD) Mixing Modeled Mass (g/cc) 1000 100 301.94 301.96 301.98 302.00 301.94 301.96 301.98 302.00 Day of Year Day of Year $\times 10^{-13}$ 1400 1600 1000 1200 1800 600 800 2.8 3.5 4.3 5.0 6.4 5.7 2.1

Diameter (nm)

...and constrains the aerosol loss

