

Profiles of the water vapor isotope composition for determining regional water sources and trace gas exchange in the boundary layer

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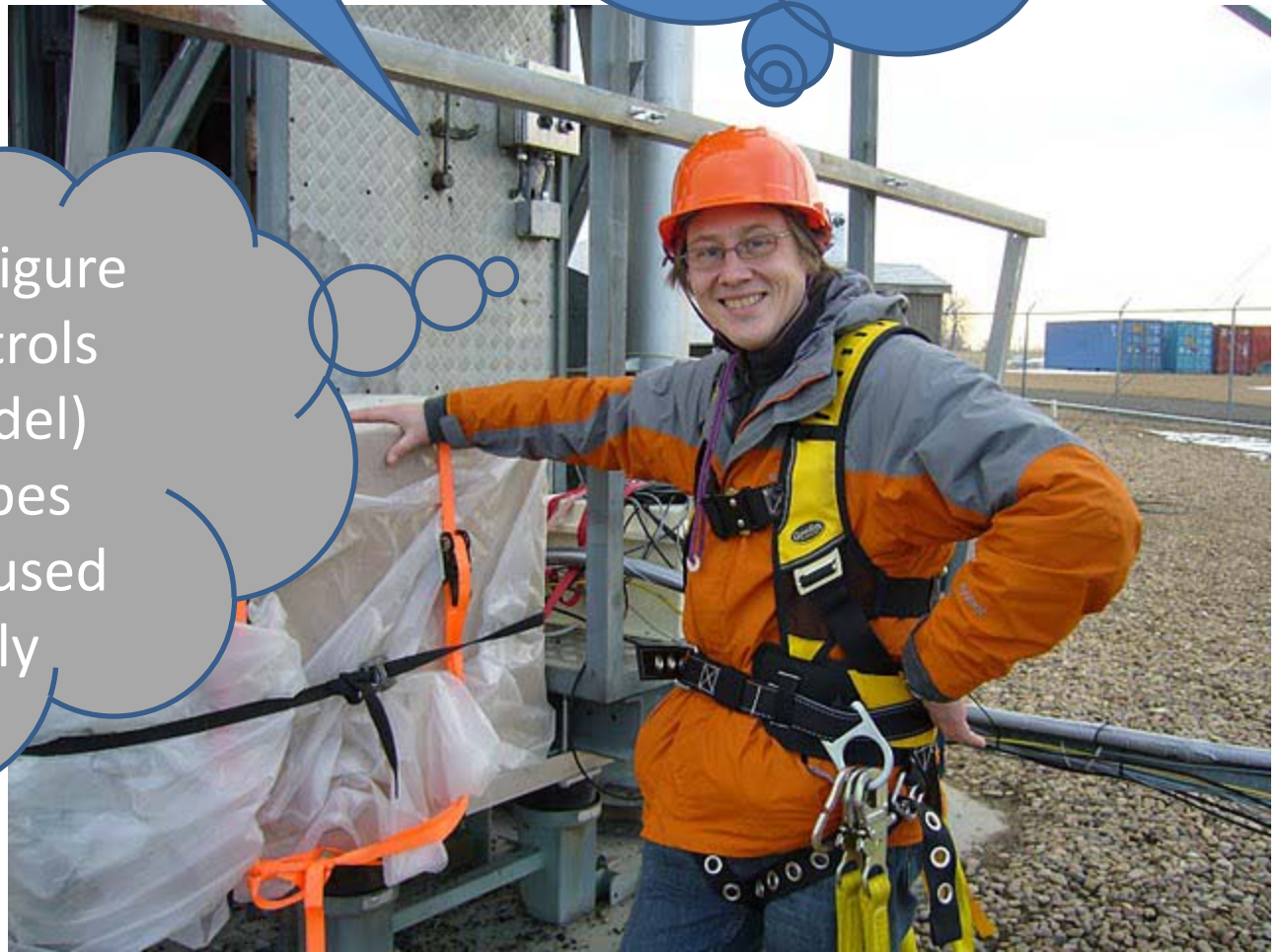
And thanks to Emily Graham

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I don't care about isotopes!

I care about the underlying physical system: water and energy balance, and isotopes help.

But, I need to figure out what controls isotopes (model) before isotopes (data) can be used meaningfully



The challenge...

- Trace gas exchange between the boundary layer and free troposphere is the leading uncertainty in estimating surface emissions from inversions (Transcom, Gurney et al. 2003)
- Boundary layer dynamics/mixing is poorly represented in models (certainly theoretical limitations associated with the treatment of stable layers, also resolution)
- ***Aim: use water vapor isotopic measurements to resolve PBL transport, then apply to other gases (say, CO₂, aerosols,...)***

Need detailed mechanistic studies.

Cases:

1. ***BAO/Niwot Ridge continental***
2. *Summit (AWO) Greenland, super stable*
3. ***Mauna Loa/marine boundary layer***

Conceptual basis

Water in the PBL
(non condensing case)

$$\frac{\partial q}{\partial t} = -\frac{\partial}{\partial z} \left(k \frac{\partial q}{\partial z} \right) + E$$

Isotopes (with HDO/H₂O ratio R)

$$\frac{\partial Rq}{\partial t} = -\frac{\partial}{\partial z} \left(k \frac{\partial Rq}{\partial z} \right) + R_E E$$

2 equations, 2 unknown (E and k).

Works exceedingly well, with hyperbolic relationship between q and R
(i.e., “Keeling plot”)

Trace gas “X”, known k find source (respiration, ... etc)

$$\frac{\partial X}{\partial t} = -\frac{\partial}{\partial z} \left(k \frac{\partial X}{\partial z} \right) + \text{source}$$

Experiment

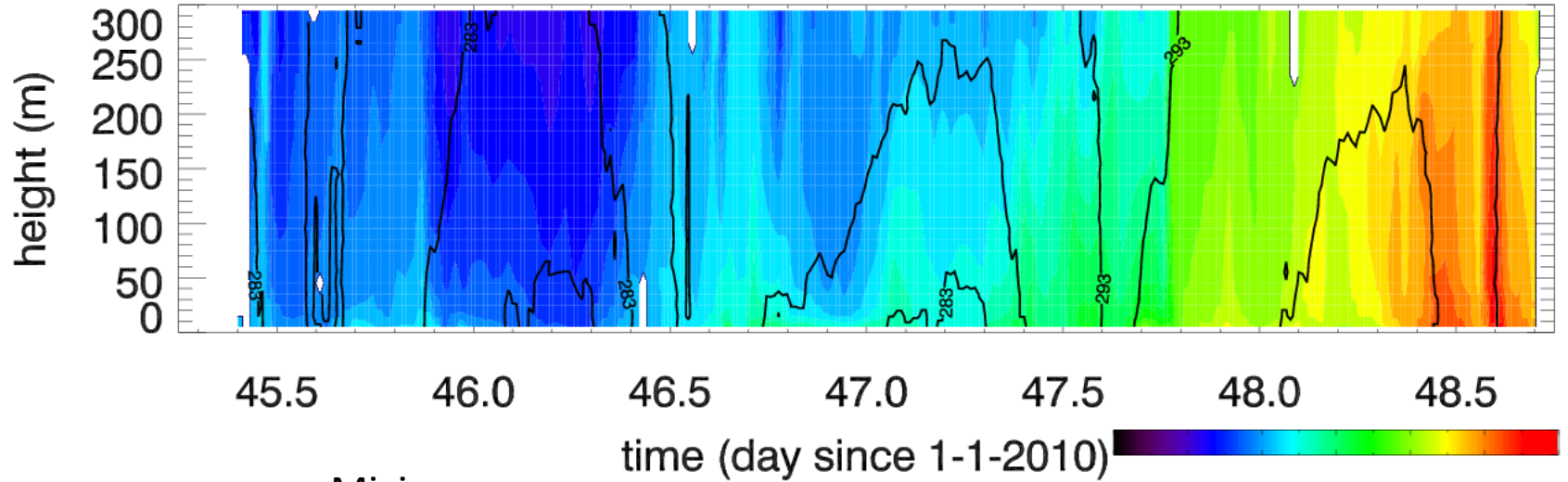
- Mounted a Picarro Water Vapor Isotope Analyzer on the BAO instrument carriage.
- Along with additional sensors
- LiCor open path CO₂/H₂O, temperature, pressure, sonic wind (Peter Blanken, Emily Graham and Dan Wolfe)
- Every 15 minutes, elevator went up or down for about 4 days (Feb 15-18, 2010)
- Ascent takes 8m50s, decent 8m30s
- Data mapped to high resolution profiles
- 312 profiles 0-300 meters with approximately 5-20 meter resolution (depending on instrument response)



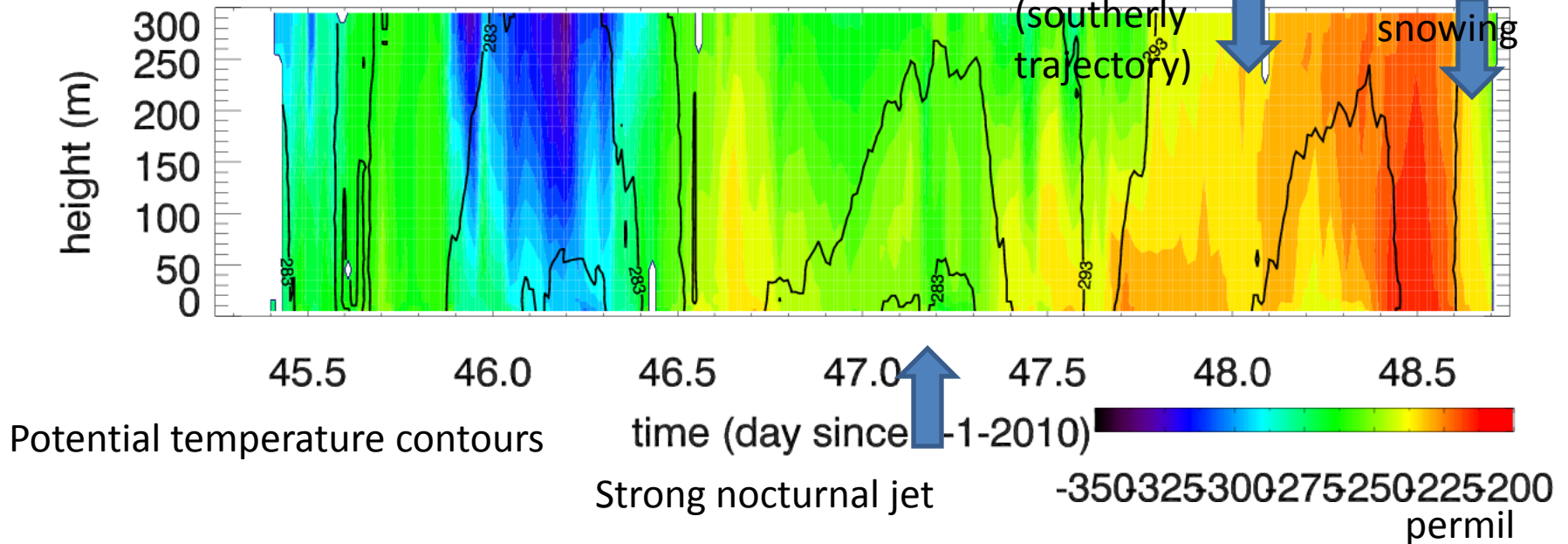
Cement
factory
(~20 km)



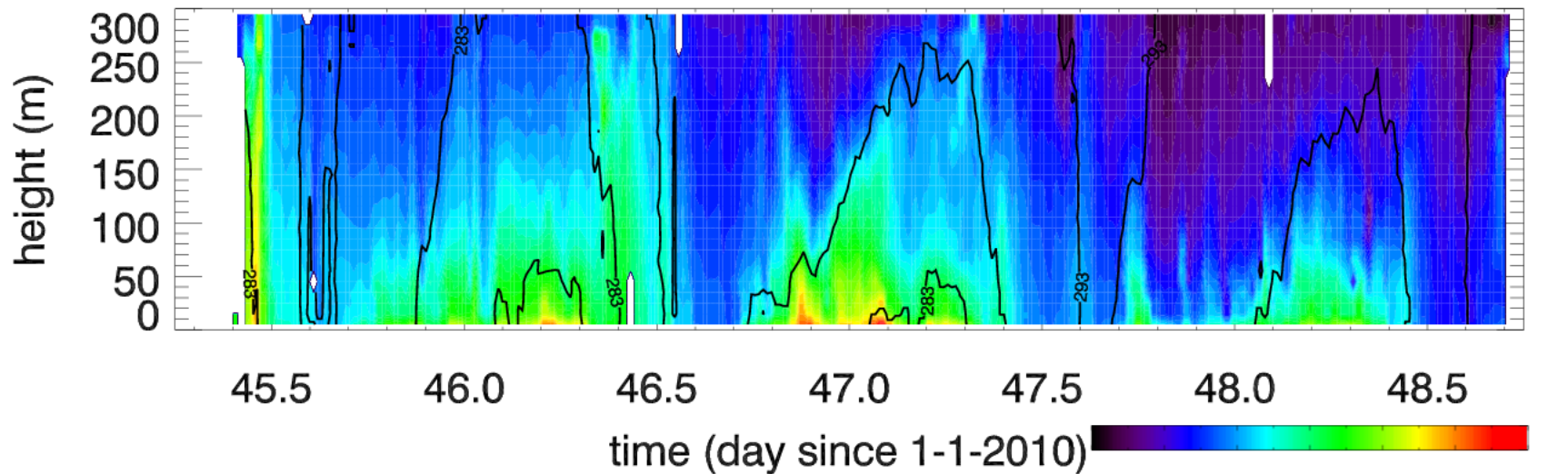
Picarro H2O



Picarro deID

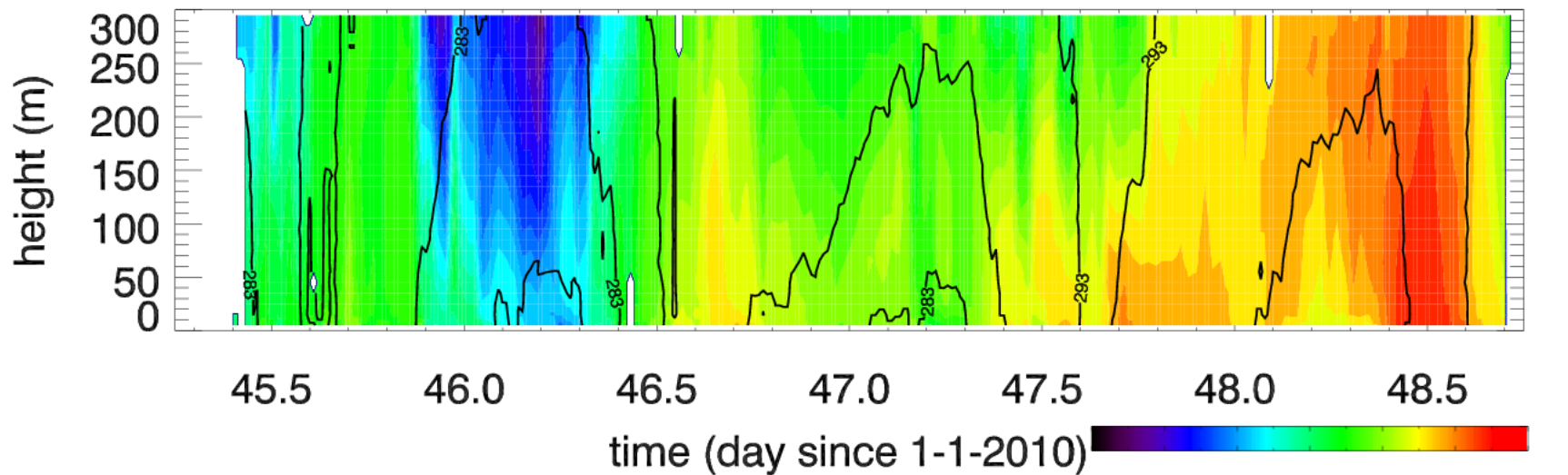


LiCor CO2



380394407421435448462

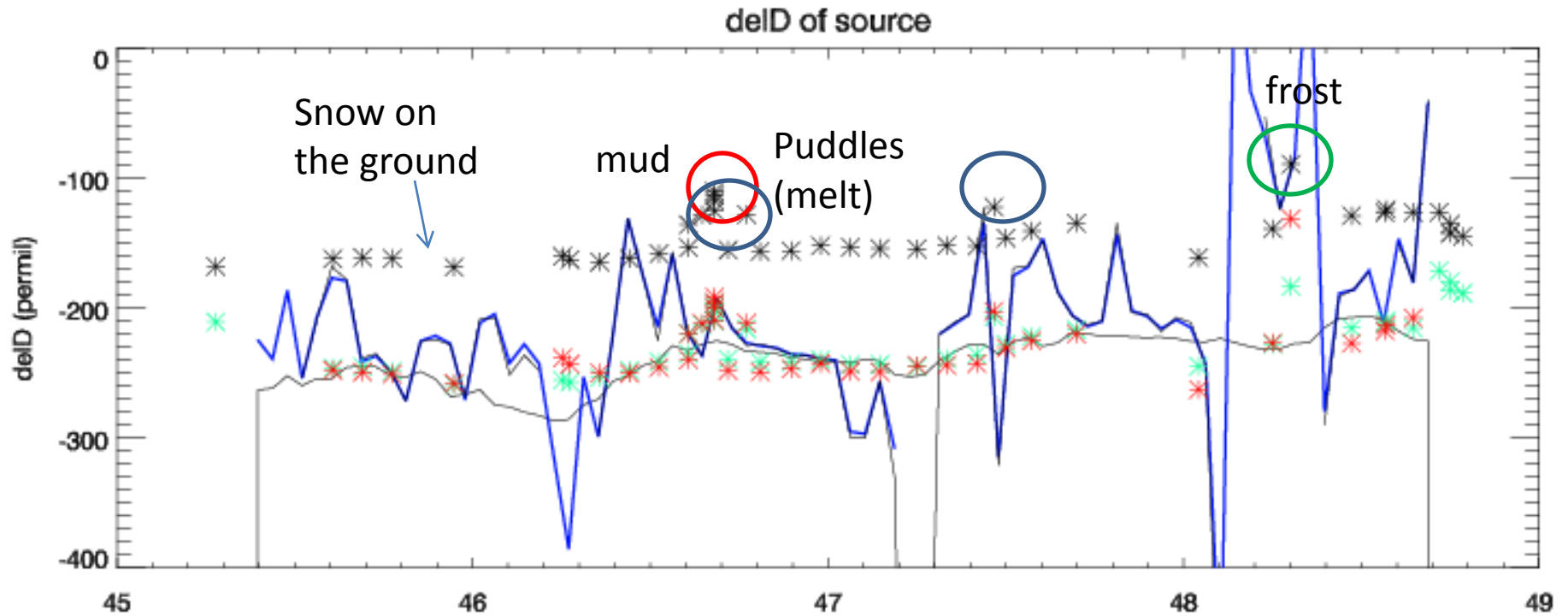
Picarro deID



-350325300275250225200

Constrain moisture source

Flux gradient/profile method, equivalent to “Keeling plot”



Blue curve: isotopic composition derive from profile

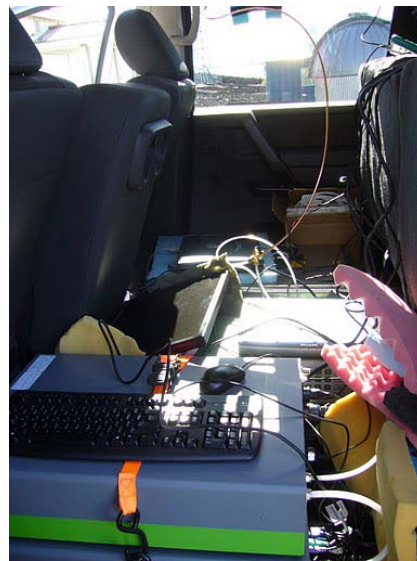
Red and cyan: isotopic composition of evaporation derived from measured surface water

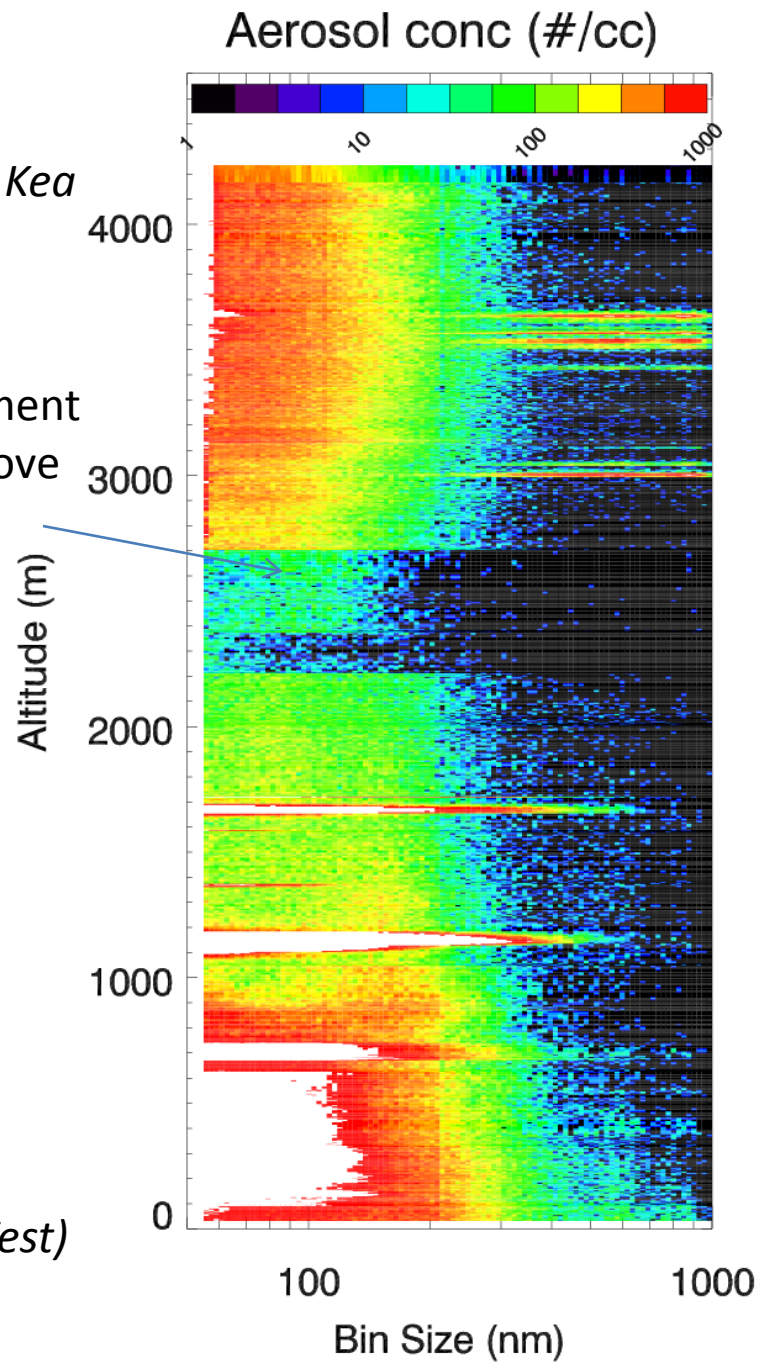
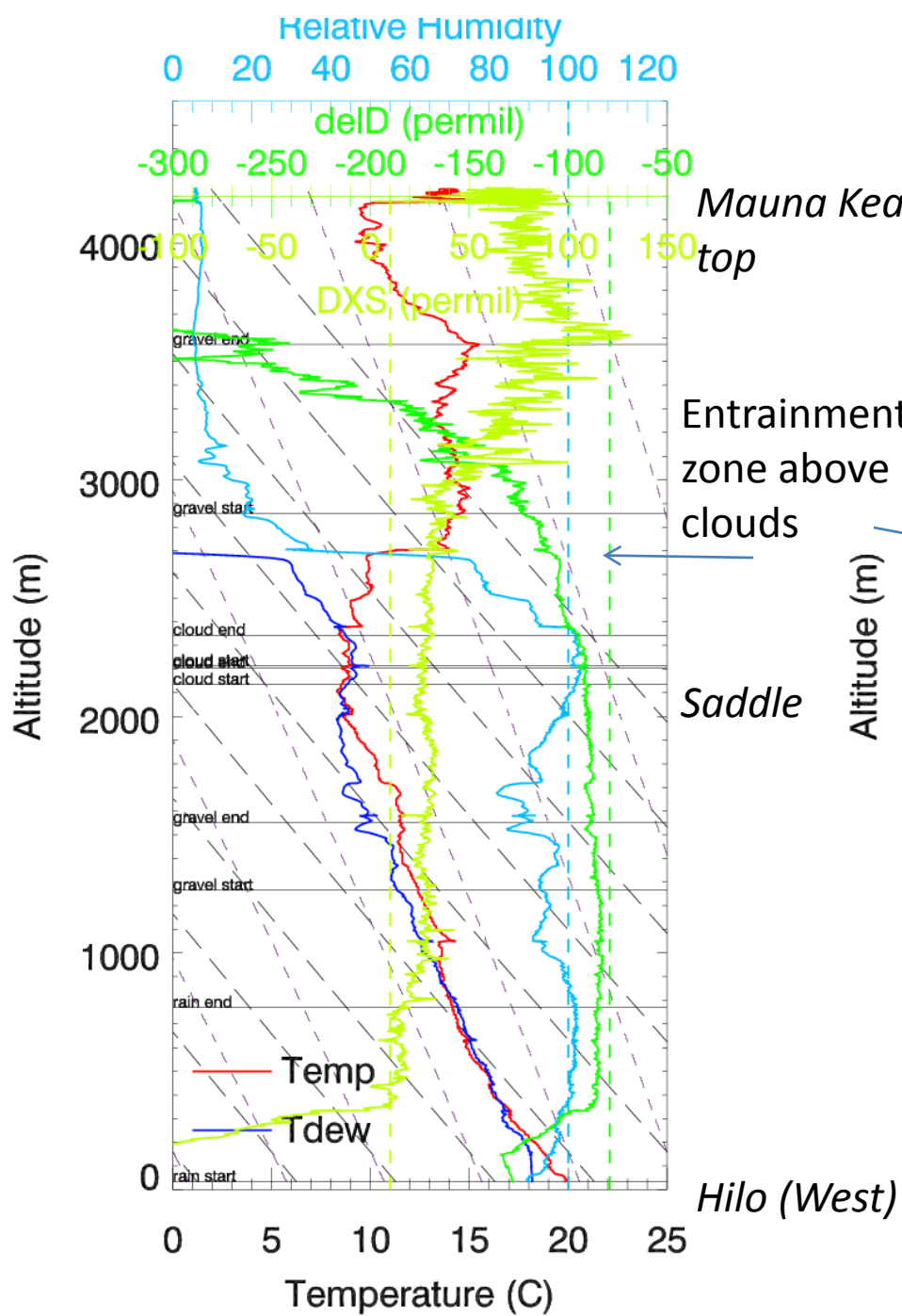
Sneak peek: new results from MLO

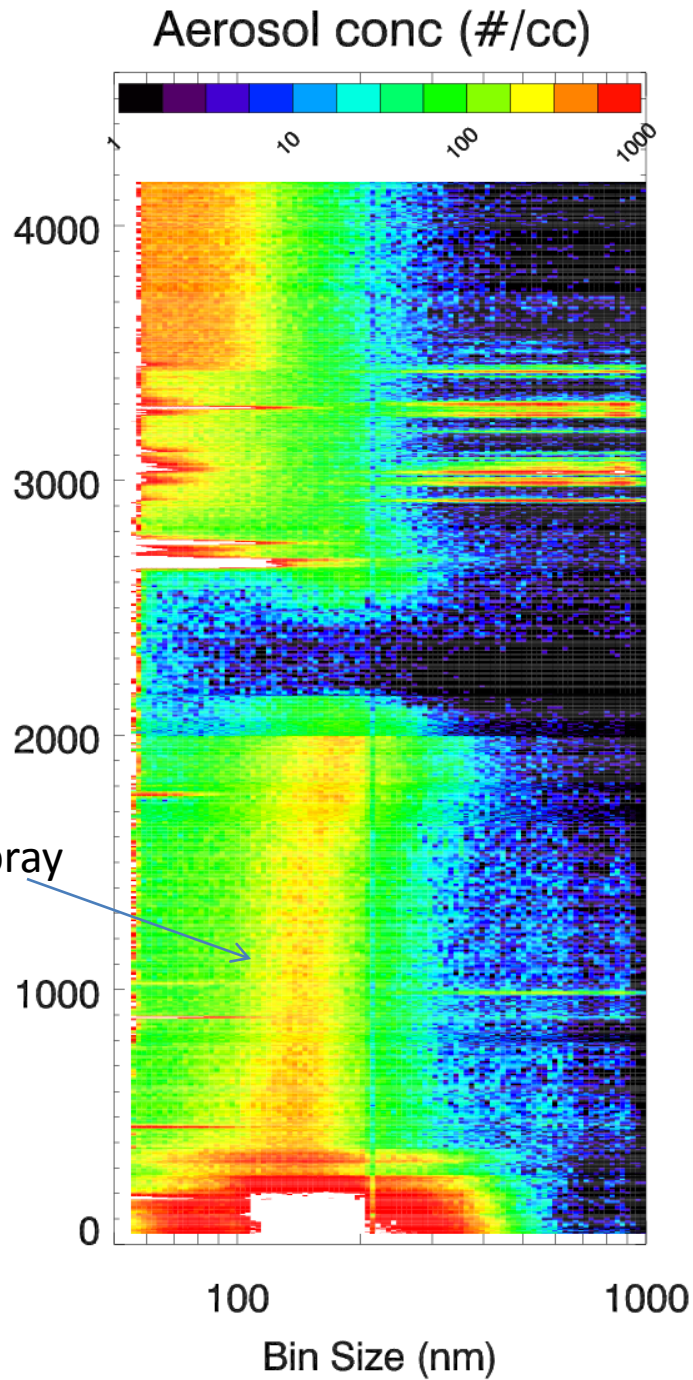
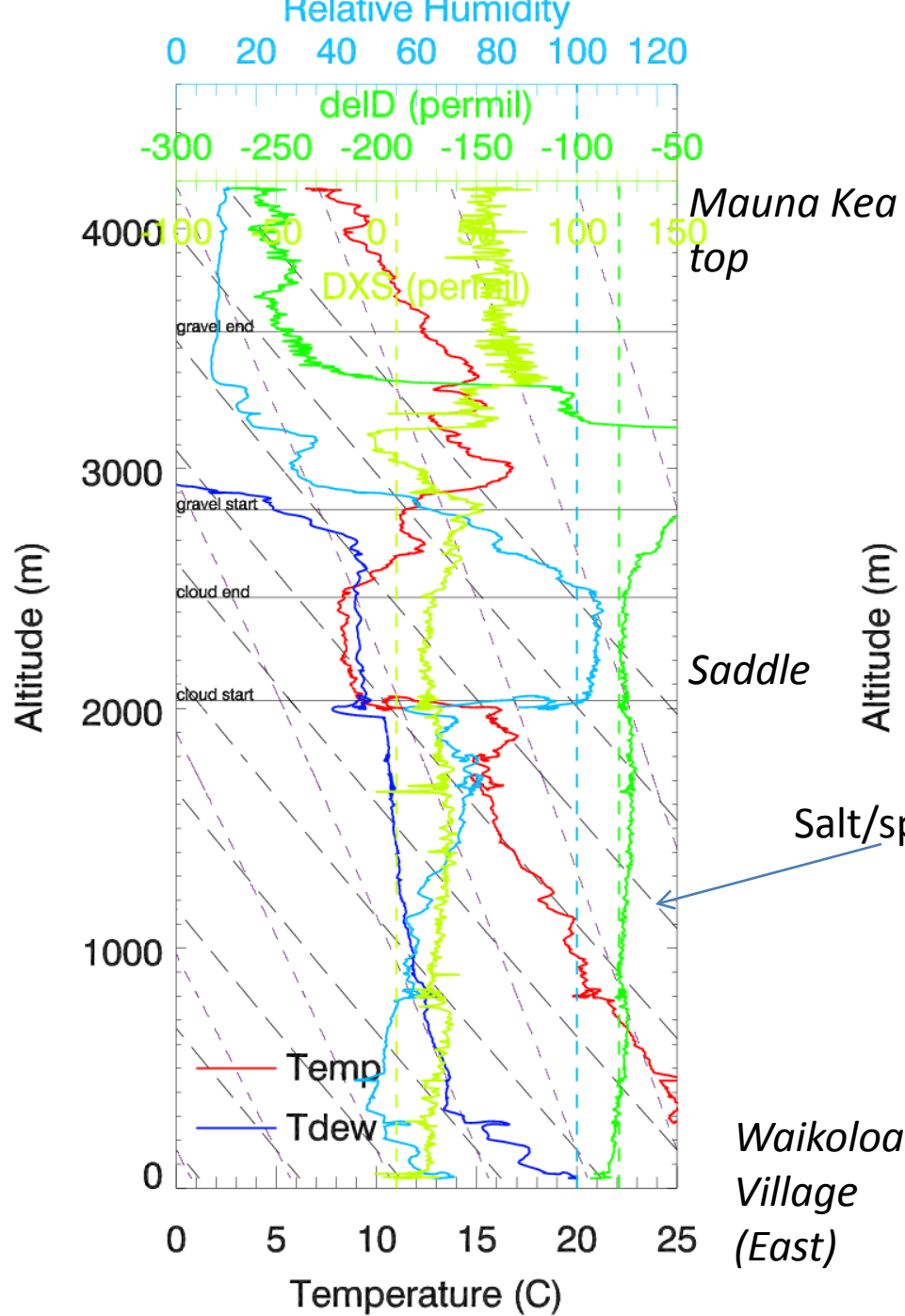
Similar approach, applied to the marine boundary layer.

Isotopes and aerosols at MLO provide metrics on vertical/local and long ranged air mass influence on baseline measurements

- *What is the degree of exchange between the MBL and the free troposphere, versus horizontal transport?*
- PBL profiles obtained by car mounted instruments
- Ascent from sea level to the top of Mauna Kea (4200 m)







Conclusions

- Combination of H₂O and isotopic measurements provides ***very clear signature of air mass mixing*** (vertical, and distinct lateral air masses)
- This can be used to identify strong transport conditions for other trace gases (here, CO₂ and aerosols...)
- Local water sources can be identified. (This study, snow melt, then evaporation from “puddles”)

- Expect growing season water source to be dominated by transpiration, this can tie to photosynthesis (knowing water use efficiency) to partition biospheric CO₂ source
- Study to continue at BAO (thanks to NSF)

