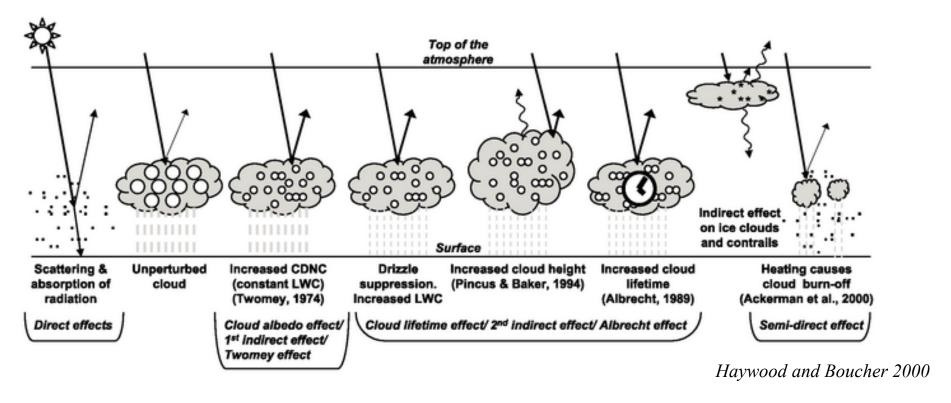


The aerosol influence on cloud radiative forcing (indirect forcing) is the largest uncertainty in current climate forcing models.



Which aerosols activate to form cloud droplets is an important factor in reducing the indirect forcing uncertainty.

Cloud Condensation Nuclei (CCN) activation to water droplets

Köhler Equation:

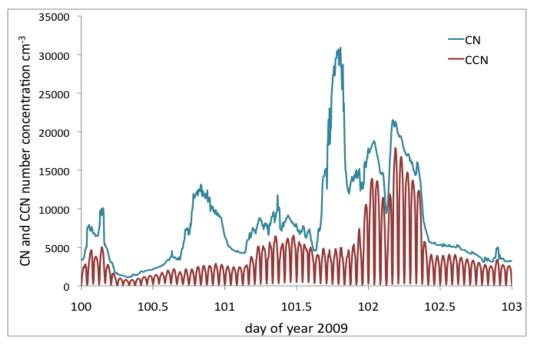
$$\ln\left(\frac{p_{w}(D)}{p^{o}}\right) = \ln(S) = \frac{4M_{w}\sigma_{w}}{RT\rho_{w}D} - \frac{6n_{s}M_{w}}{\pi\rho_{w}D^{3}}$$

$$size \quad chemistry$$

Need the size-dependent aerosol composition to determine the CCN activation properties

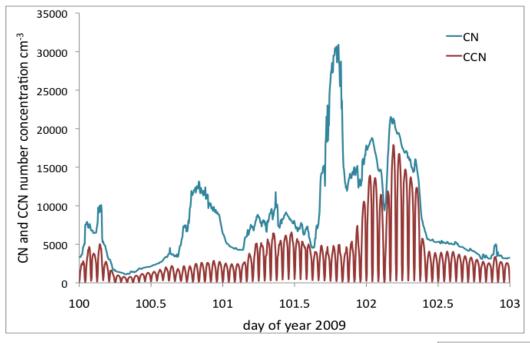
Single hygroscopicity parameter: *κ Petters and Kreidenweiss, ACP, 2007.*

Observations of CN, CCN and scattering coefficient



CN and CCN number concentrations often don't correlate well.

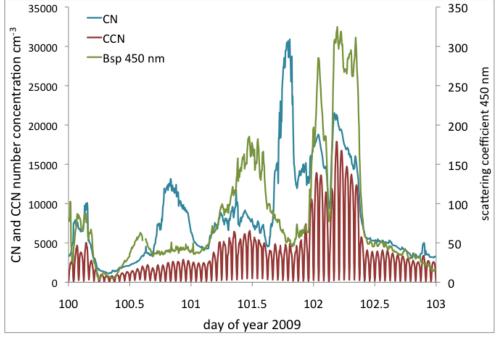
CN number concentrations are highest at the smallest particle sizes which also have the shortest lifetime and are least likely to live long enough to form CCN.



CN and CCN number concentrations often don't correlate well.

CN number concentrations are highest at the smallest particle sizes which also have the shortest lifetime and are least likely to live long enough to form CCN.

The aerosol scattering coefficient, which follows the aerosol surface area and larger accumulation mode particles correlates better with CCN number concentrations.

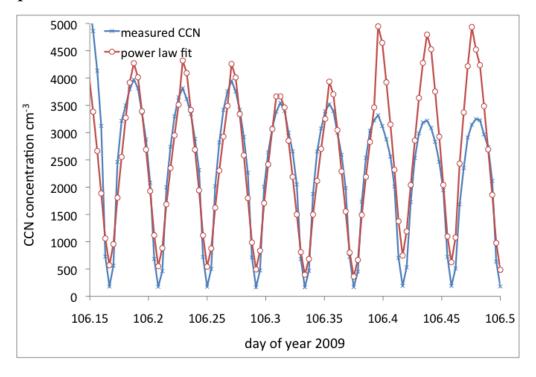


Empirical methods for estimating CCN concentrations

Twomey power law, Pure Appl. Geophys., 1959

$$CCN(\%ss) = C(\%ss)^k$$

"C" scales with CCN number concentration. "k" scales with the particle solubility such that high k values have low solubility and highly soluble particles like sea salt have low k values.

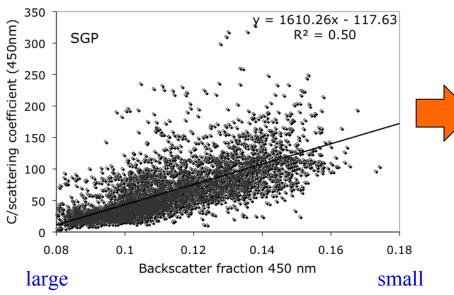


The power law fit doesn't work well at high %ss as k increases with %ss.

At low %ss values <0.1% the CCN instrument has kinetic limitations.

Model CCN over range of %ss from ~0.2 to 0.8%ss.

Model approach: correlate CCN power law fit parameters to aerosol optical properties



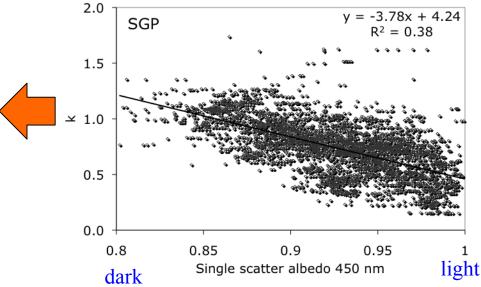
 $C/\sigma sp(450 \text{ nm}) = m \cdot BSF(450 \text{nm}) + b$

Normalize C to σ sp.

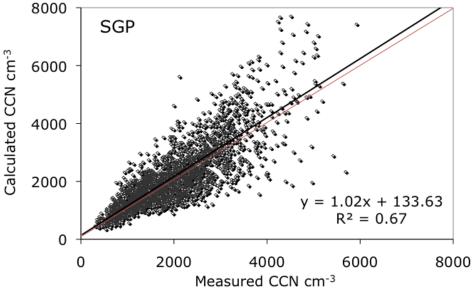
Use backscatter fraction (BSF) as proxy of the aerosol size

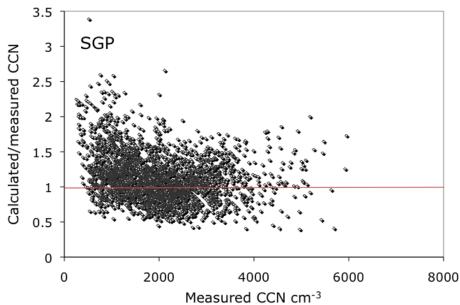
$$k = p \cdot SSA(450 \text{ nm}) + d$$

Use single scatter albedo (SSA) as a proxy of the aerosol composition



Model results for Oklahoma (SGP)





Model predicts the measured CCN values using only the measured optical properties.

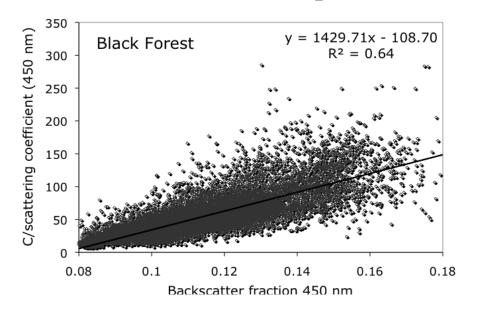
Uncertainty is highest at low CCN concentrations when uncertainty in optical data is is highest.

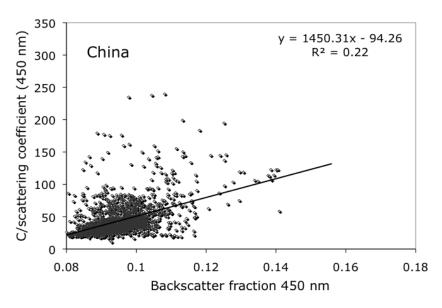
SGP is an agricultural region. The aerosol is mostly organic mixed with sulfate and nitrates.

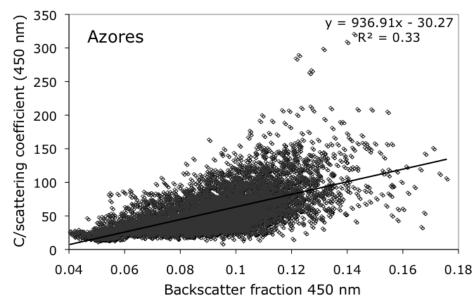
The aerosol is well behaved in that aerosol size and albedo are closely correlated.

Other regions?

Correlation of C parameter to backscatter fraction



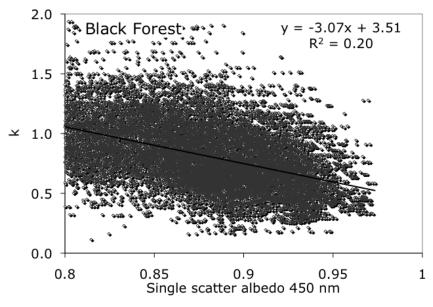


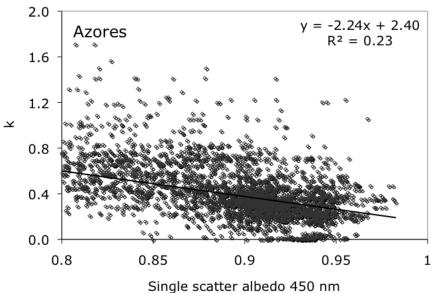


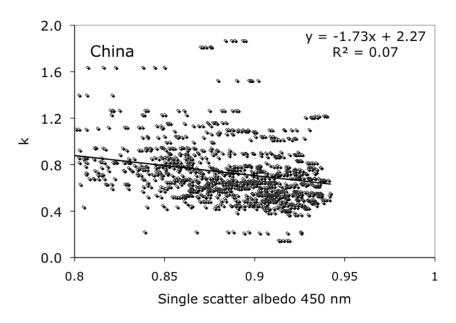
China has a high fraction of large dust particles

The slope of $C/\sigma sp$ vs BSF is significantly smaller for marine aerosol and indicates an aerosol with a high scattering efficiency.

Correlation of k parameter to single scatter albedo







Changes in the aerosol absorption or "effective black carbon" have a weak influence on CCN formation for dust and sea salt aerosol.

Location	k value (stdev)
Oklahoma	0.73 (0.26)
Black Forest	1.06 (0.65)
China	0.76 (0.33)
Azores	0.43 (0.42)

Model results from other sites

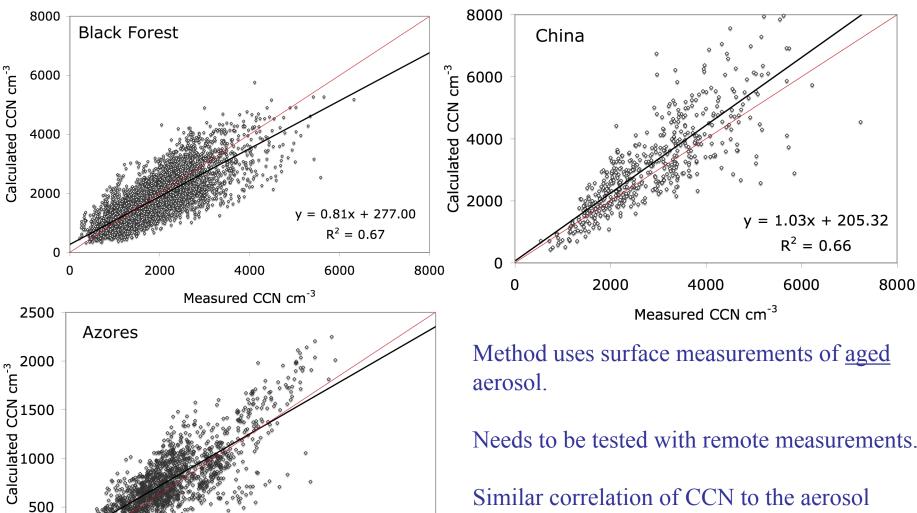
500

0

500

1000

Measured CCN cm⁻³



y = 0.88x + 159.65

 $R^2 = 0.59$

2000

1500

2500

Method uses surface measurements of aged aerosol.

Needs to be tested with remote measurements.

Similar correlation of CCN to the aerosol index (AI) has been observed with surface and HSRL lidar measurements.