

# Partitioning of terrestrial carbon sources using $^{14}\text{CO}_2$ : observations and modeling

Scott Lehman

*INSTAAR, University of Colorado at Boulder*

J. Miller, P. Tans, C. Sweeney, A. Andrews - *CIRES/NOAA/ESRL, Boulder*

**C. Wolak** - *INSTAAR, Boulder*

J. Southon - *UC Irvine, California*

B. LaFranchi, T. Guilderson - *LLNL/CAMS, California*

M. Fischer - *LBL, California*; J. Turnbull - *RRL, NZ*



# sensitivity to fossil fuel CO<sub>2</sub> (C<sub>ff</sub>)

$$1/\lambda_{^{14}\text{C}} = 8223 \text{ yr}$$

$$\Delta^{14}\text{C}\text{‰} = (^{14}\text{C}:\text{C}_{\text{sa}}/^{14}\text{C}:\text{C}_{\text{std}} - 1) * 1000$$

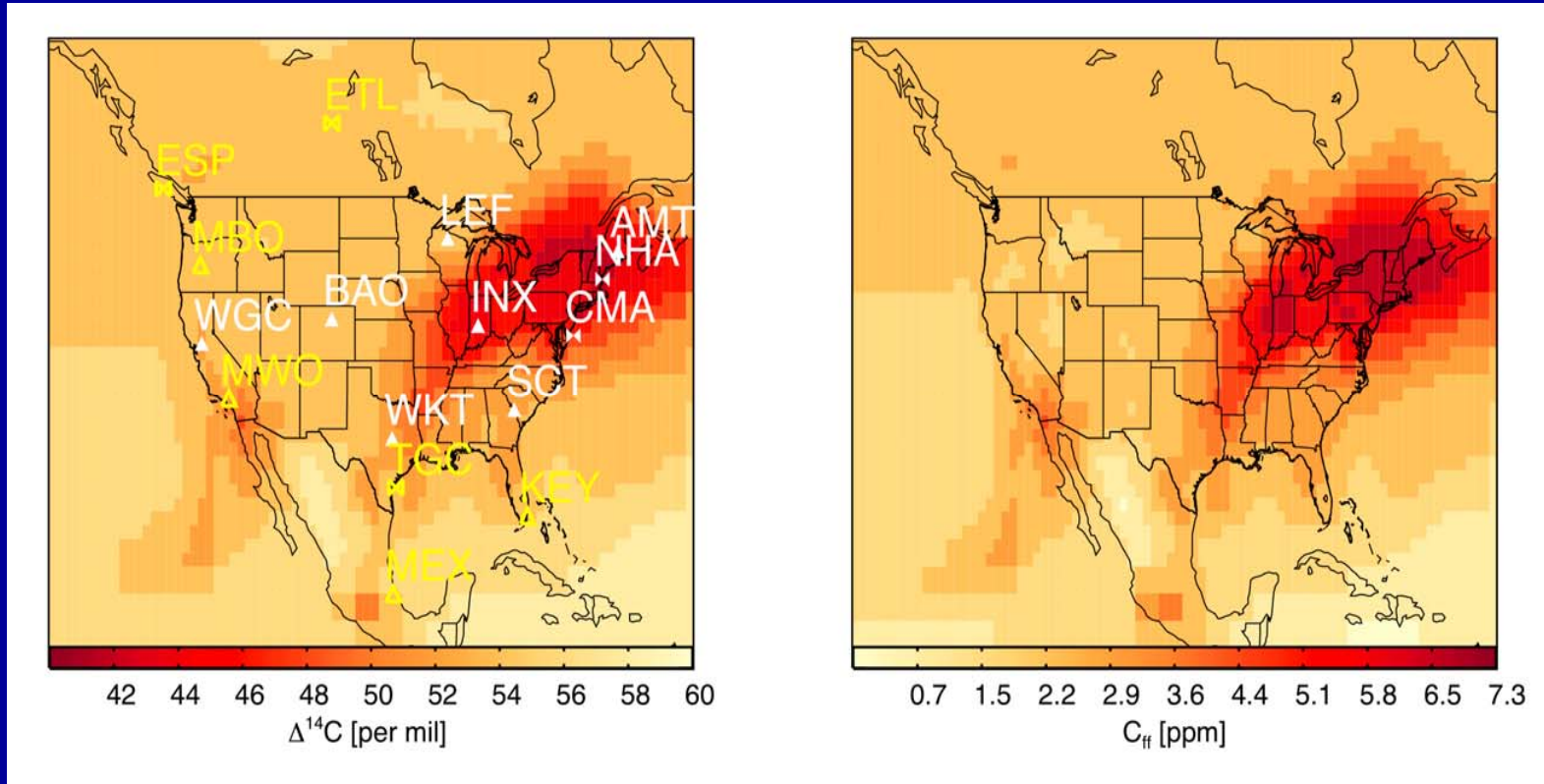
fossil fuel CO<sub>2</sub>  $\Delta = -1000\text{‰}$  (i.e. zero <sup>14</sup>C content)

ambient atm. (& other CO<sub>2</sub> sources)  $\Delta = \sim +50\text{‰}$

$$\frac{\Delta\Delta_{\text{ff-atm}}}{C_{\text{atm}}} = \frac{-1050\text{‰}}{390\text{ppm}} = \sim -2.7 \text{‰/ppm}$$

- detection of  $\sim 1$  ppm for recently-added fossil-fuel derived CO<sub>2</sub> (C<sub>ff</sub>) requires measurement precision of  $\sim 2\text{‰}$
- detection is unbiased if other contributions to tropospheric <sup>14</sup>C distribution over large land areas are small

# $\Delta^{14}\text{CO}_2$ vs. $C_{ff}$ in TM5



- distribution of  $C_{ff}$  dominates  $\Delta^{14}\text{CO}_2$  signal over NH land areas (figures scaled according to mass balance relation of  $-2.7\text{‰}/\text{ppm}$ )
- small differences primarily due to terrestrial disequilibrium flux of  $^{14}\text{C}$  (quantifiable)
- near-surface  $\Delta^{14}\text{C}$  gradients  $\sim 14\text{‰}$  for this week in Jan. 2006 ( $\sim 8\text{x}$  precision)

# quantification of $C_{ff}$

$$C_{obs} = C_{bg} + C_{ff} + C_r + C_p$$

$$\Delta_{obs} C_{obs} = \Delta_{bg} C_{bg} + \Delta_{ff} C_{ff} + \Delta_r C_r + \Delta_p C_p$$

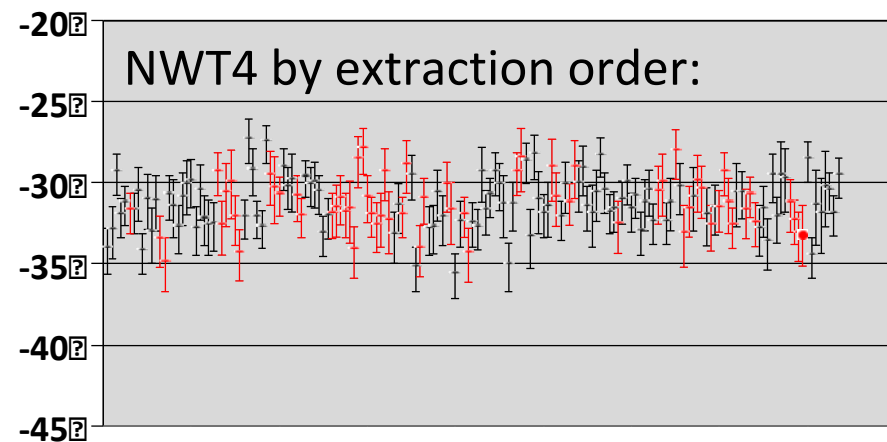
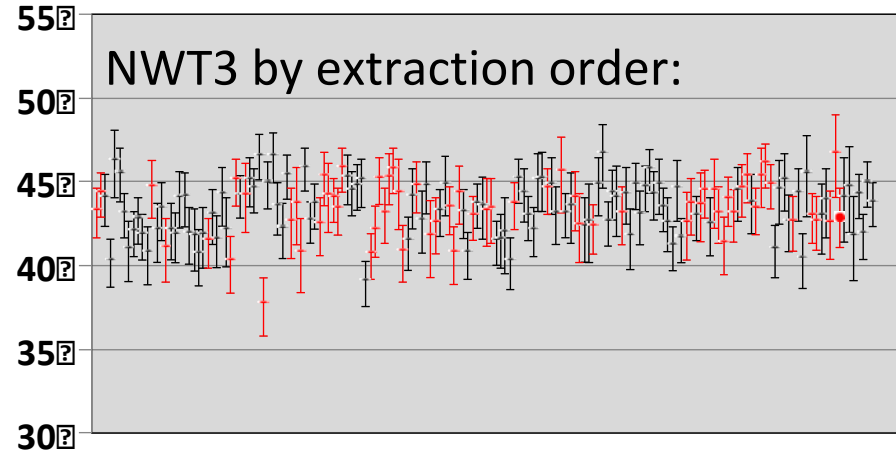
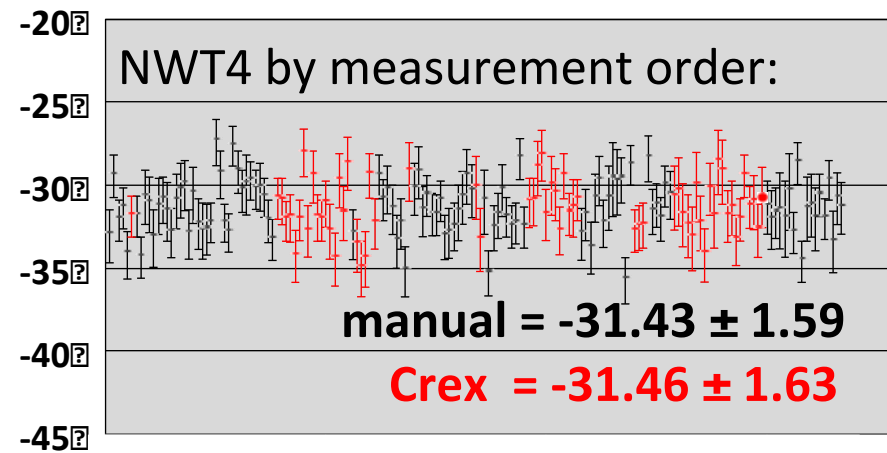
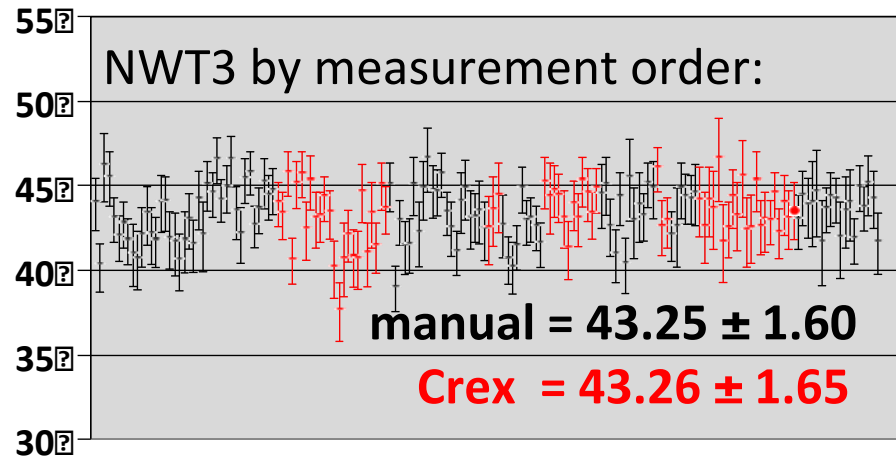
*and setting  $\Delta_p = \Delta_{bg}$ :*

$$C_{ff} = \frac{C_{obs}(\Delta_{obs} - \Delta_{bg})}{\Delta_{ff} - \Delta_{bg}} - \frac{C_r(\Delta_r - \Delta_{bg})}{\Delta_{ff} - \Delta_{bg}}$$

Turnbull et al. 2006

- right hand term is correction to  $C_{ff}$  of ~0.2 - 0.5 ppm (i.e.,  $C_r$  and  $\Delta_r$  can be independently quantified)
- $C_{ff}$  detection effectively limited by quadrature sum uncertainty of two  $^{14}\text{C}$  measurements ( $\Delta_{obs}$ ,  $\Delta_{bg}$ )
- isolation of  $C_{ff}$  delivers net  $C_{bio}$  (w/ same uncertainty as for  $C_{ff}$ )

# CU-UCI 1 $\sigma$ repeatability ( $\Delta^{14}\text{C}$ ‰)

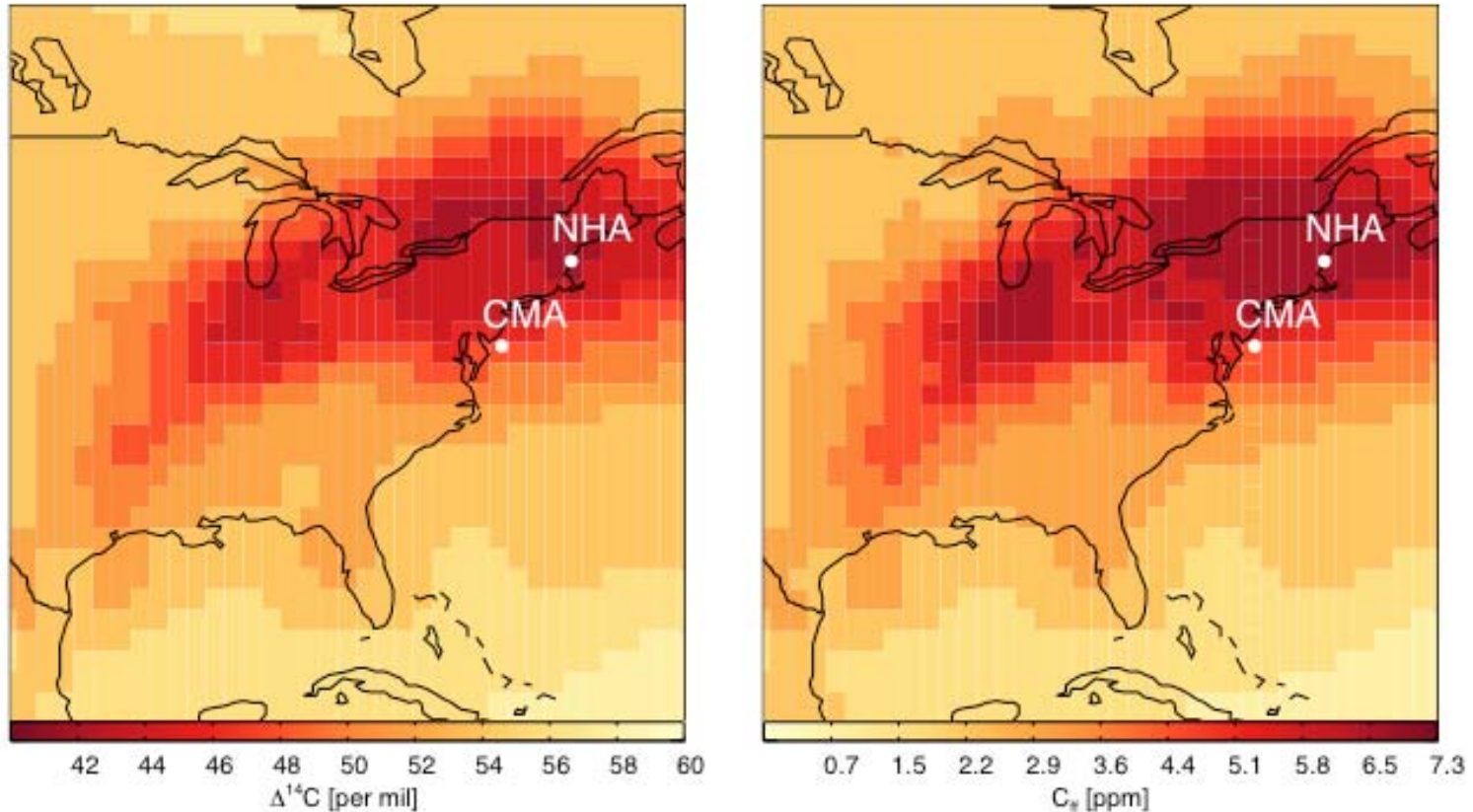


NWT3 and NWT4 replaced NWTstd (near exhaustion) in late 2009

NWTstd 1 $\sigma$  repeatability 2003-2009 = 1.8 ‰

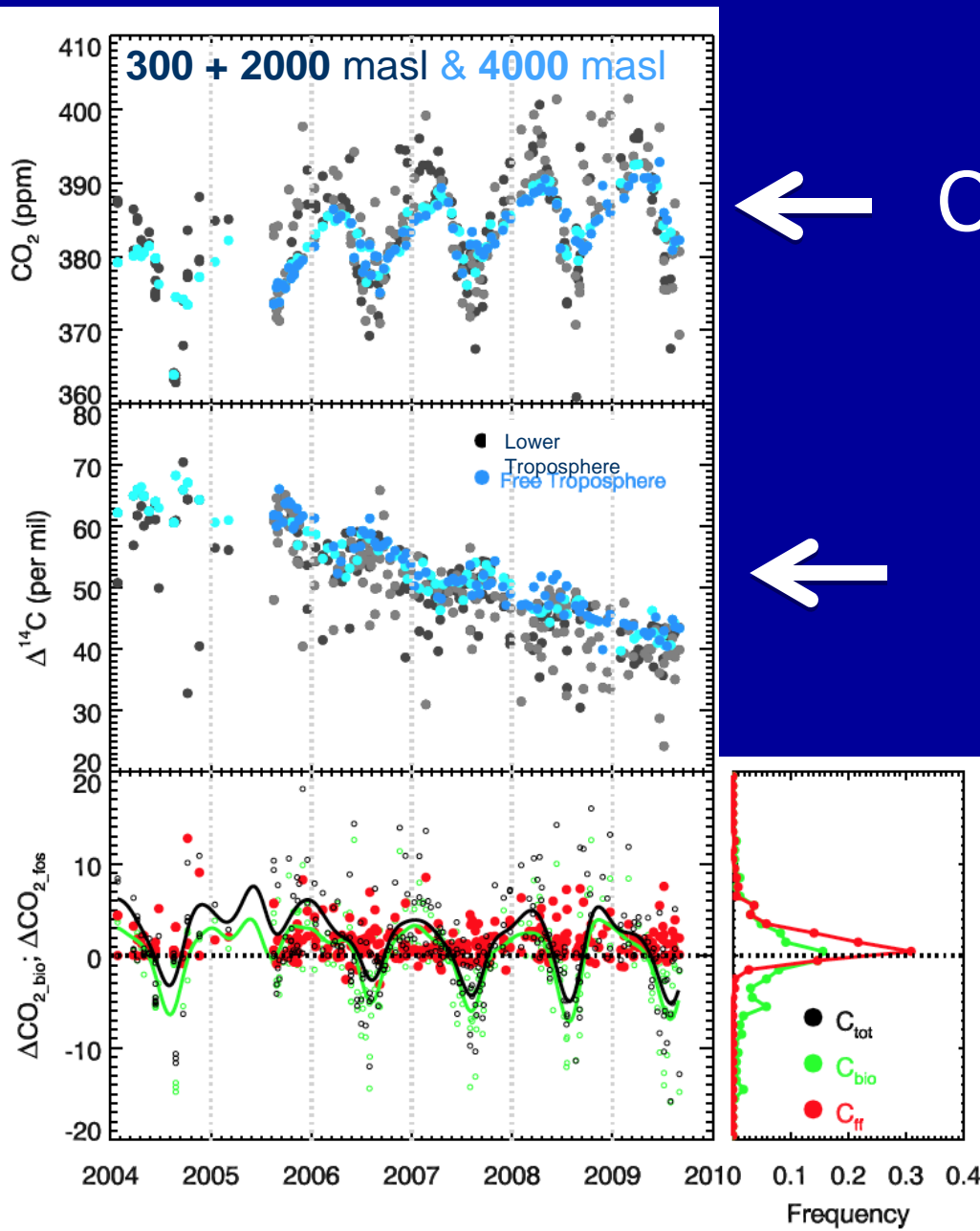
report 1 $\sigma$  repeatability or 1 $\sigma$  single sample precision, whichever is larger

# NE U.S. example [Miller et al. '12]



samples from 5-6 yrs of bi-weekly vertical aircraft profiles in area of significant regional emissions and outflow  
3-ht. sampling for  $^{14}\text{CO}_2$ , 9-ht. for all other gases

# isolation of $C_{ff}$ and $C_{bio}$ , CMA+NHA



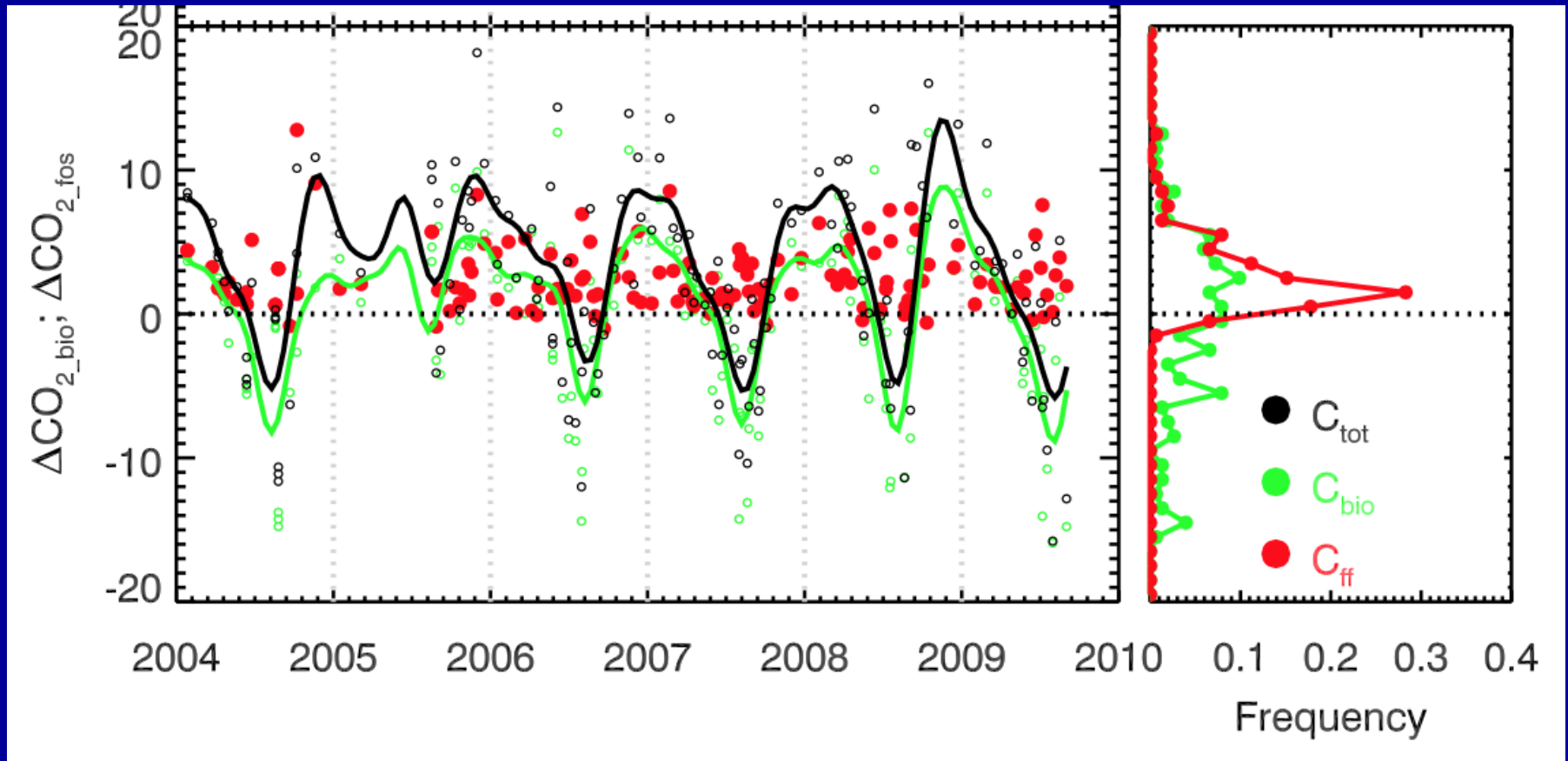
$$C_{obs} = C_{bg} + C_{ff} + C_{bio}$$

$$\Delta_{obs} = \Delta_{bg} + \Delta_{ff}$$

- obs fr. ~300 and ~2000 masl (vs. 4000 m 'bg' in FT)
- $C_{ff}$  detectable year round (1-10 ppm)



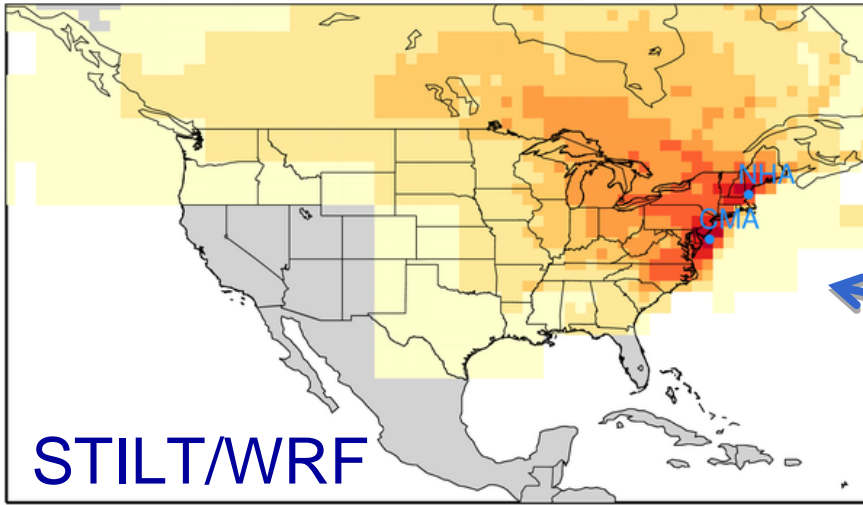
# 300 m obs only, CMA+NHA



- sharpens view of biospheric signal (vs. raw  $\text{CO}_2$  enhancement)
- $C_{\text{bio}}$  large even in winter (~60% of total winter-time enhancement) despite urban/industrial observational footprint
- $\text{CO}_2$ -only methods: can not assume urban enhancements are due to  $C_{\text{ff}}$
- $C_{\text{ff}}$  and  $C_{\text{bio}}$  independently useful

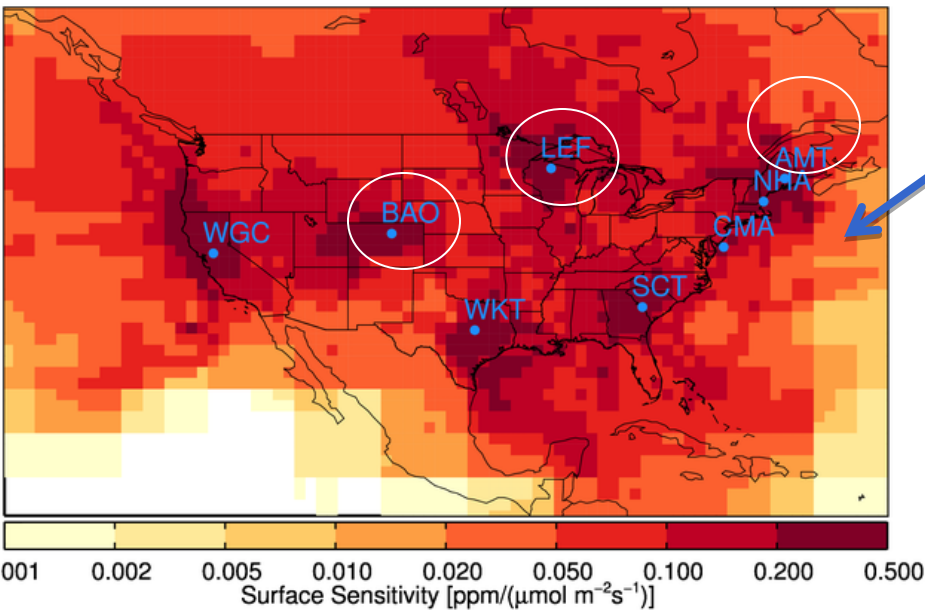


# growing observational footprint



annual average sensitivity for all PBL  $^{14}\text{C}$  obs

[Miller et al., 2012]

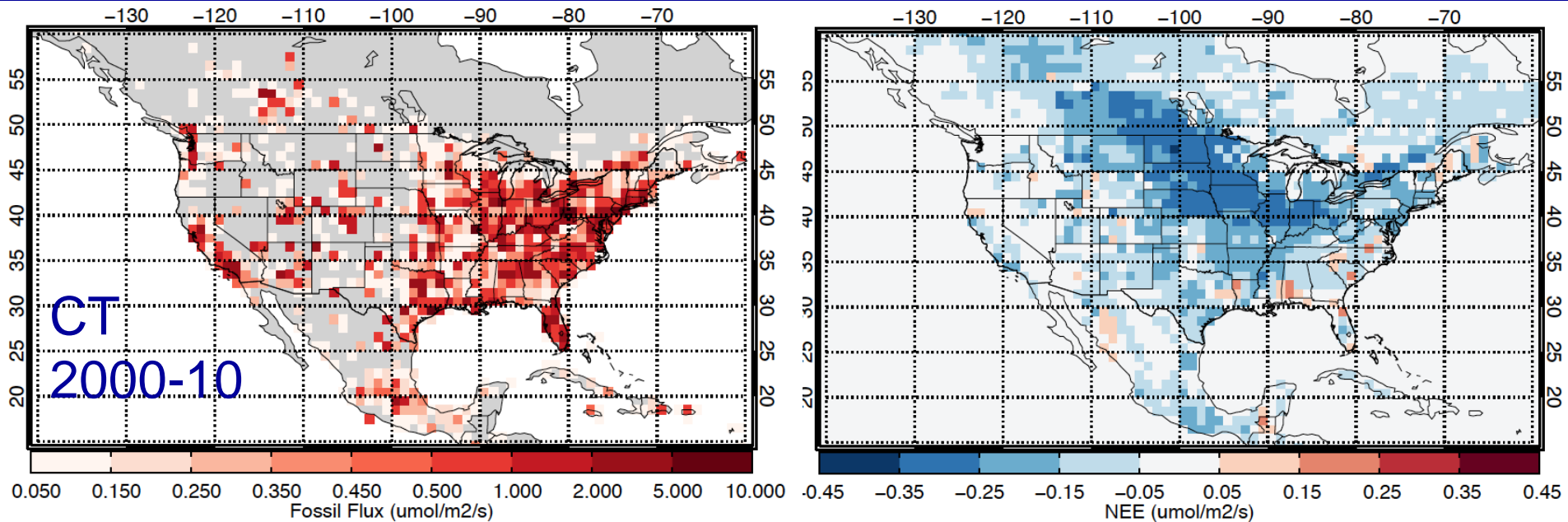


since late '09 from ~thrice weekly tower (PBL)  $^{14}\text{C}$  obs (excludes INX, MWO)

*$^{14}\text{C}$  footprint sufficient for meaningful guidance of CarbonTracker, using both  $^{14}\text{C}$  and  $\text{CO}_2$  as obs constraints*

circled sites: collaborative w/ LLNL/CAMS

# use of $^{14}\text{C}$ + $\text{CO}_2$ in CT to improve NEE



$F_{\text{ff}}$  *prior* (given 0 uncertainty)

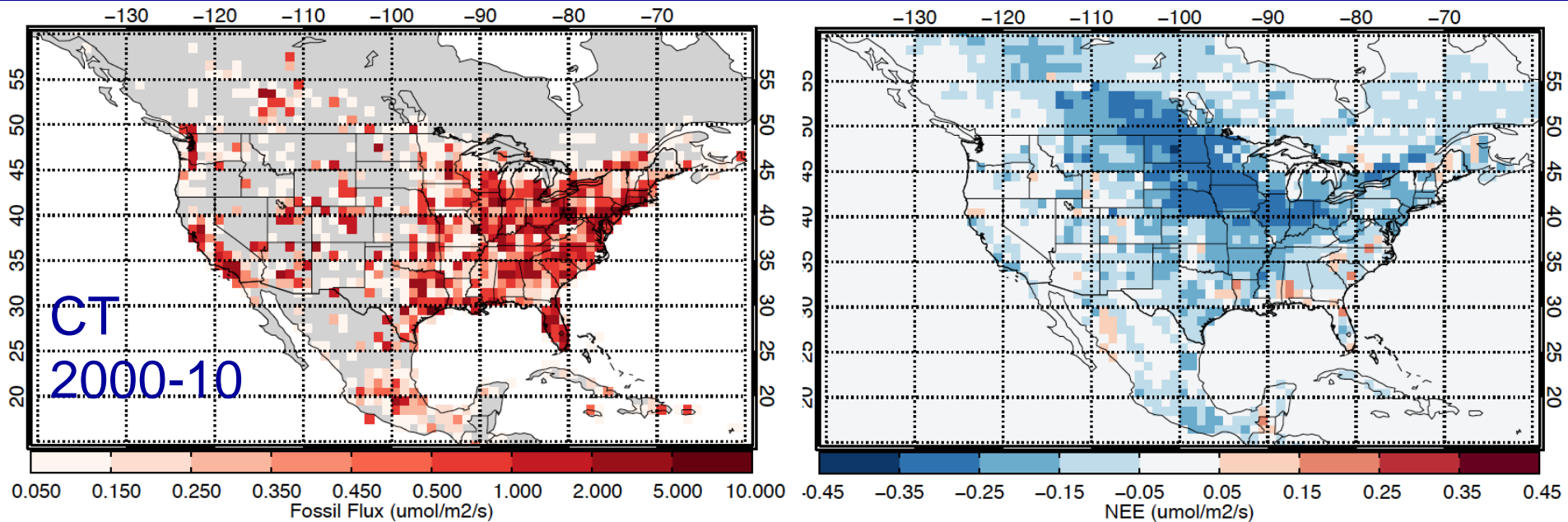
NEE *posterior* retrieval

deviation  $F_{\text{ff}}$  prior from actuals will lead directly to bias in retrieved  $F_{\text{bio}}$  (NEE) from inversion of  $C_{\text{obs}}$

$$dC_{\text{obs}}/dt = F_{\text{ff}} + F_{\text{bio}} + F_{\text{fire}}$$

$F_{\text{ff}}$  is large *w.r.t* net annual  $F_{\text{bio}}$ , and.. extrapolation of  $F_{\text{ff}}$  inventories will not capture  $F_{\text{ff}}$  anomalies associated with sustained heat and cold waves

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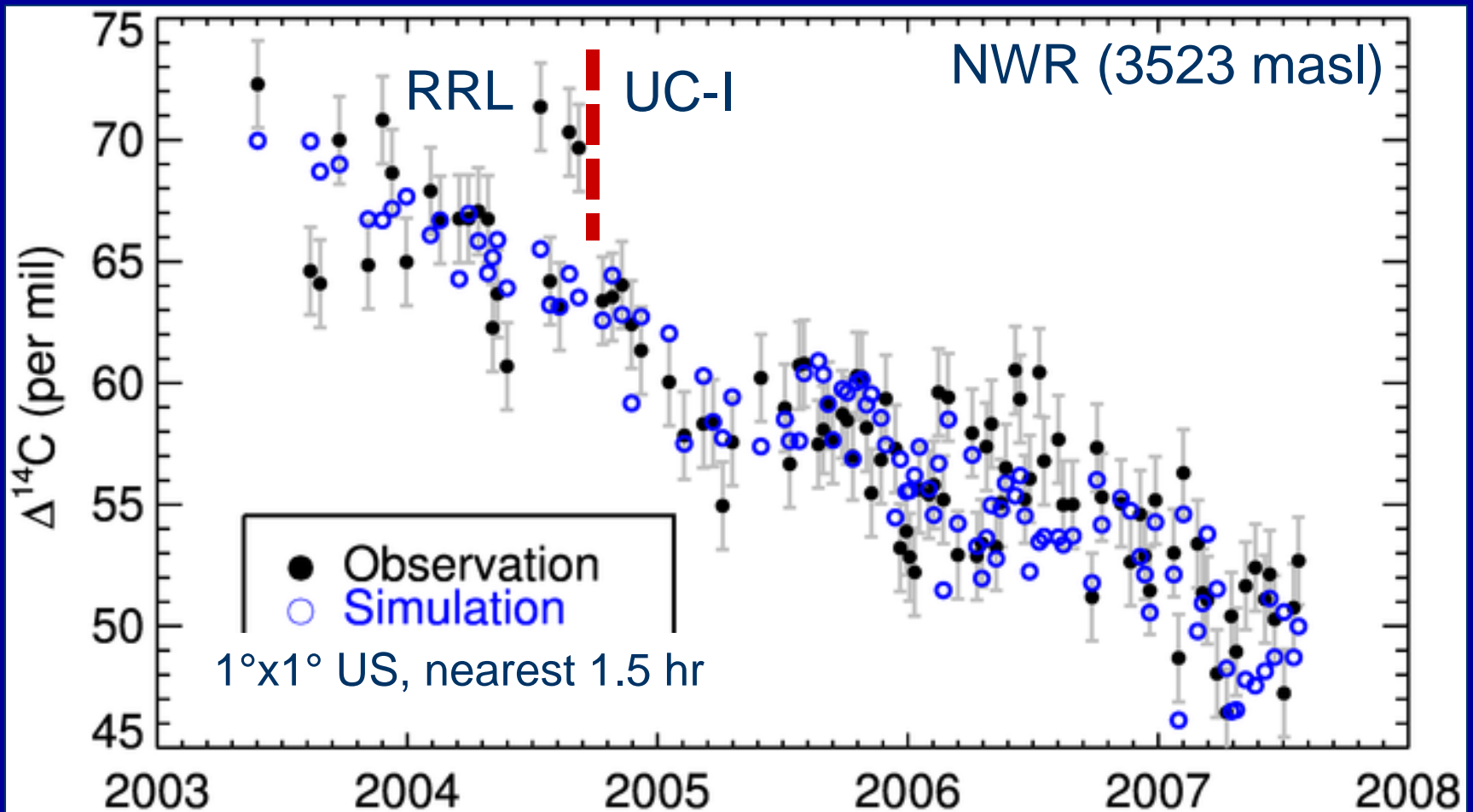
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$$dC_{\text{obs}}/dt = F_{\text{ff}} + F_{\text{bio}} + F_{\text{fire}}$$

relax  $F_{\text{ff}}$  prior uncertainty to permit guidance by  $^{14}\text{C}$  obs where available

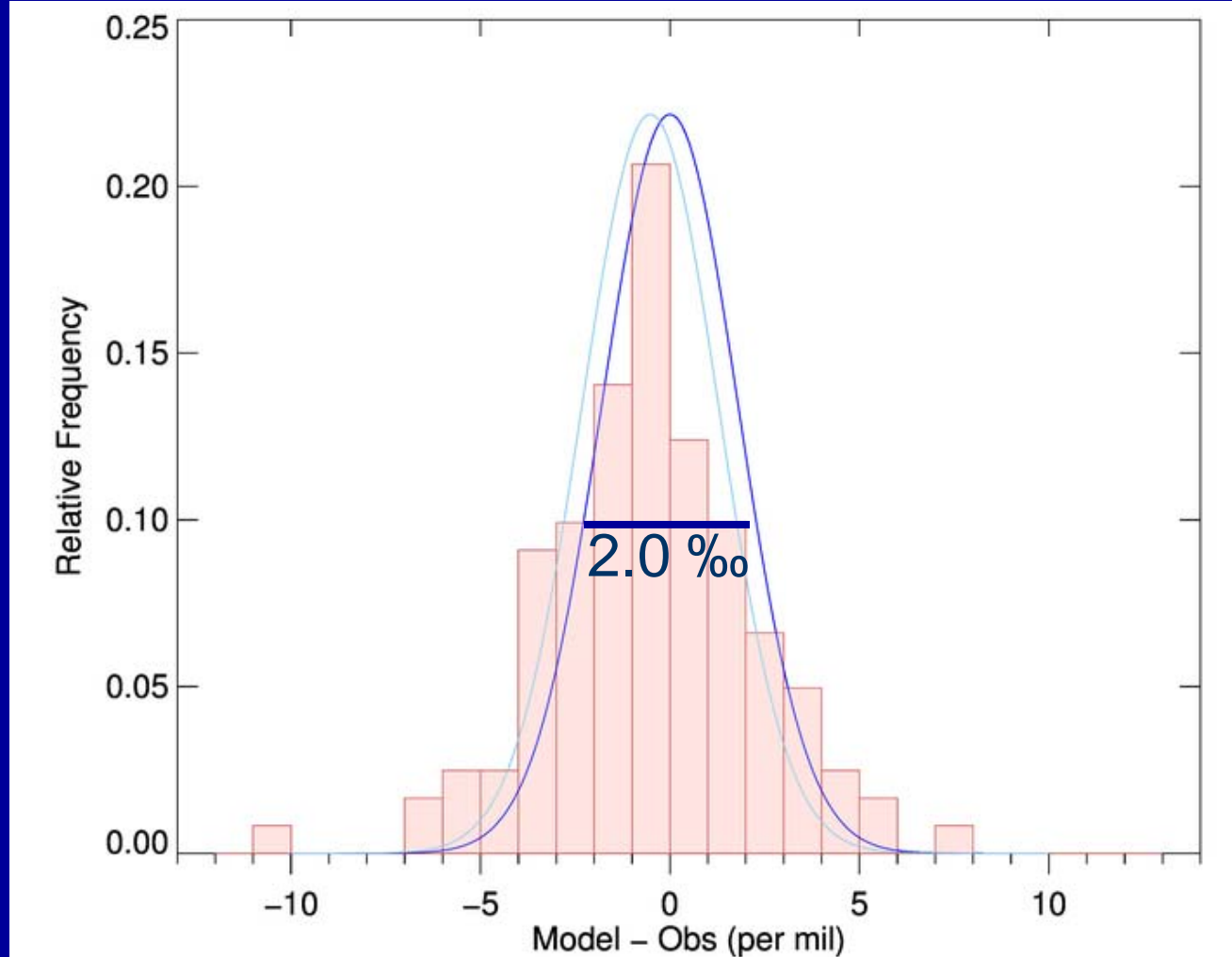
# TM5-<sup>14</sup>C vs. NWR (FT) obs



$$C_{atm} \frac{d\Delta}{dt} = (\Delta_{foss} - \Delta_{atm}) F_{fos} + \Delta_{dis} F_{surf-gross} + iso F_{cosm} + iso F_{nuc}$$

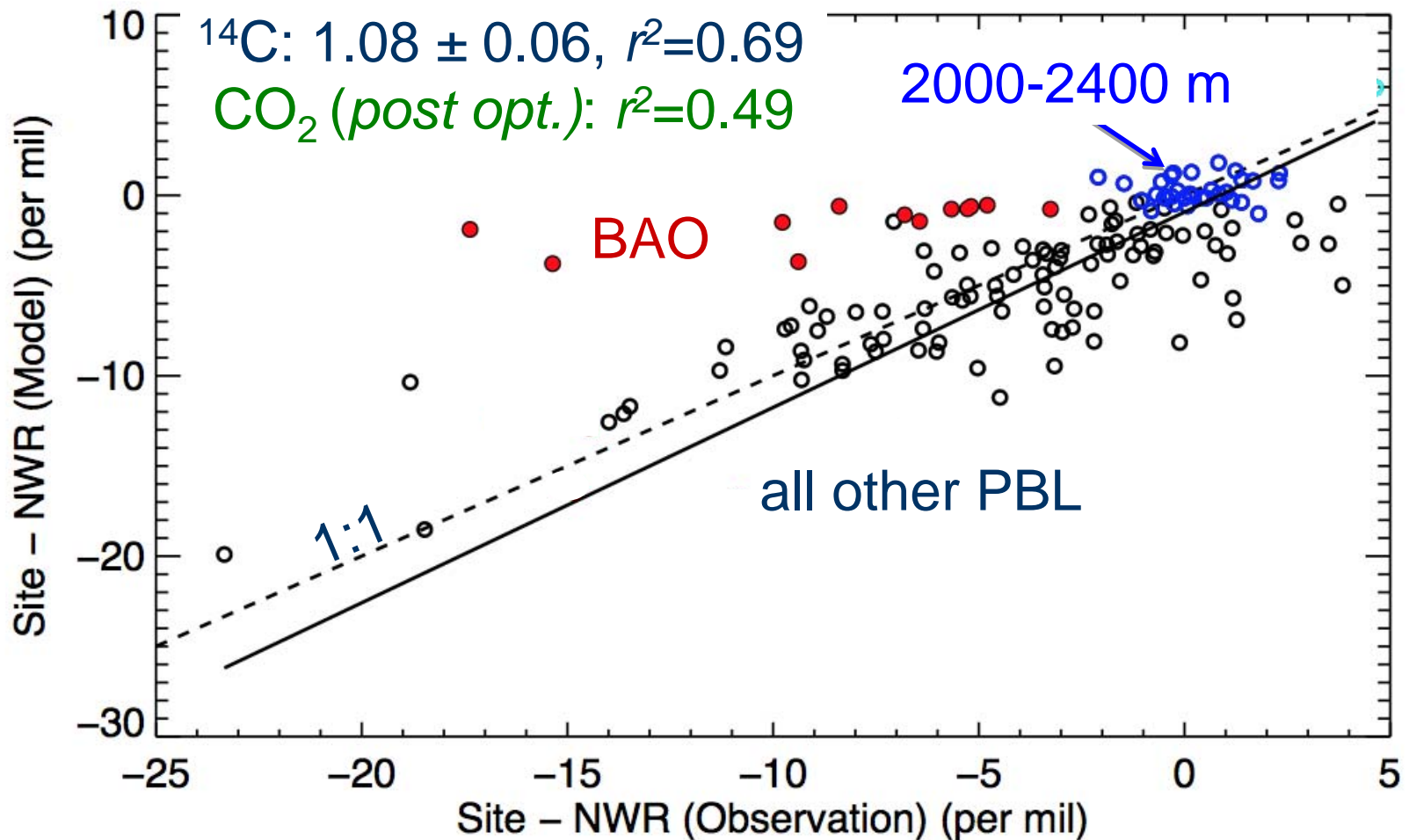
<sup>14</sup>C fluxes from *a priori* geophysical estimates & CT FF  
 (no tuning): *budget terms properly balanced, if not correct*

# residuals, NWR



90% within measurement error, bias = -0.5 per mil

# TM5-<sup>14</sup>C vs. Lower Trop. obs, US



monthly mean Site-NWR gradients,  $F_{\text{foss}}$  from CT  
(climatological m\_m obs = 132 site\_mos fr. ~2500 obs)

# summary

- developed scientifically meaningful  $^{14}\text{C}$  measurement capability
- expanded observational footprint to US national scale (and elsewhere)
- propose moving forward w/  $^{14}\text{C}$  and  $\text{CO}_2$  as dual observational constraints in CarbonTracker (for NEE,  $\text{C}_{\text{ff}}$  emissions verification)
- evaluating tracer: $\text{C}_{\text{ff}}$  emissions ratios [Miller *et al.*, JGR-A, 2012] nationally, updating absolute emissions estimates for correlate gases