

Pan-Arctic surface radiation measurements capture atmospheric preconditioning of sea ice melt season

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I. Abstract

Recent changes in the Arctic climate involve surface-atmosphere energy exchange processes and feedbacks associated with clouds, surface albedo, and the atmospheric state. The Arctic is comprised of regional climate regimes, which exhibit unique sensitivities and responses to climate change.

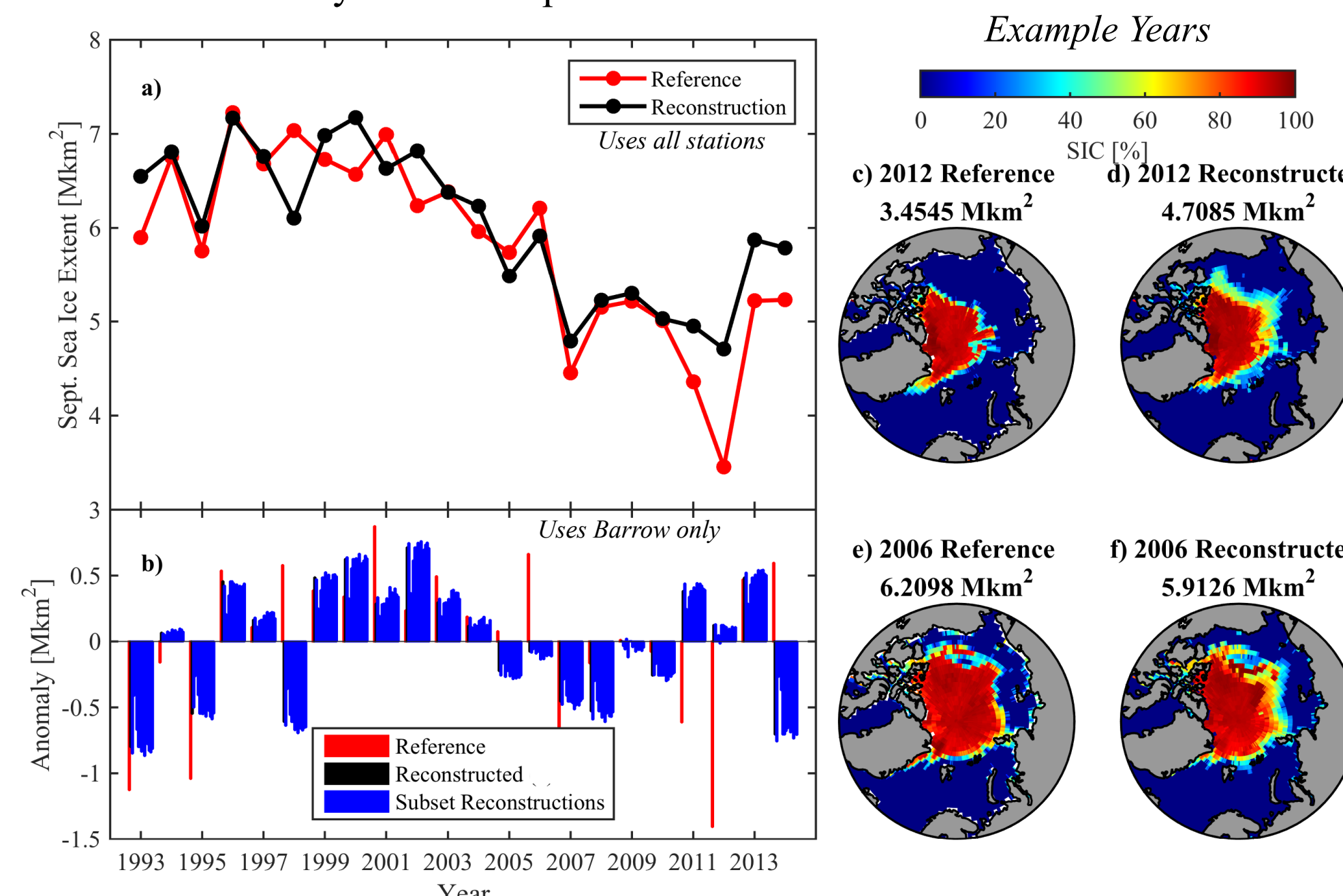
The surface radiation budget is monitored from several stations in the pan-Arctic region. Thus far, most studies have focused on individual locations, reporting significant changes in the surface radiation budget. Spatial analyses are lacking and a coordinated analysis of these measurements is needed to improve understanding of the processes involved in the changing Arctic climate. The International Arctic Systems for Observing the Atmosphere (IASOA) (<http://www.iasoa.org>) Radiation Working Group is collectively analyzing these observations. The focus locations are Tiksi (Siberia), Ny-Ålesund (Svalbard), Barrow (Alaska), and Alert (Canadian Archipelago). These stations have long records of quality measurements, enabling analyses to be conducted focusing on variability in the surface radiation budget over the past 10 to 20 years, a time period during which the Arctic has experienced dramatic changes.

In this study, radiative fluxes and cloud properties in spring (April – June) are investigated to improve understanding of atmospheric preconditioning of the sea ice melt season. In addition to direct observations from broadband radiometers, the Radiative Flux Analysis (RFA) value-added product is used. The RFA provides quality control and higher order metrics, such as cloud radiative forcing, cloud fraction, and optical depth.

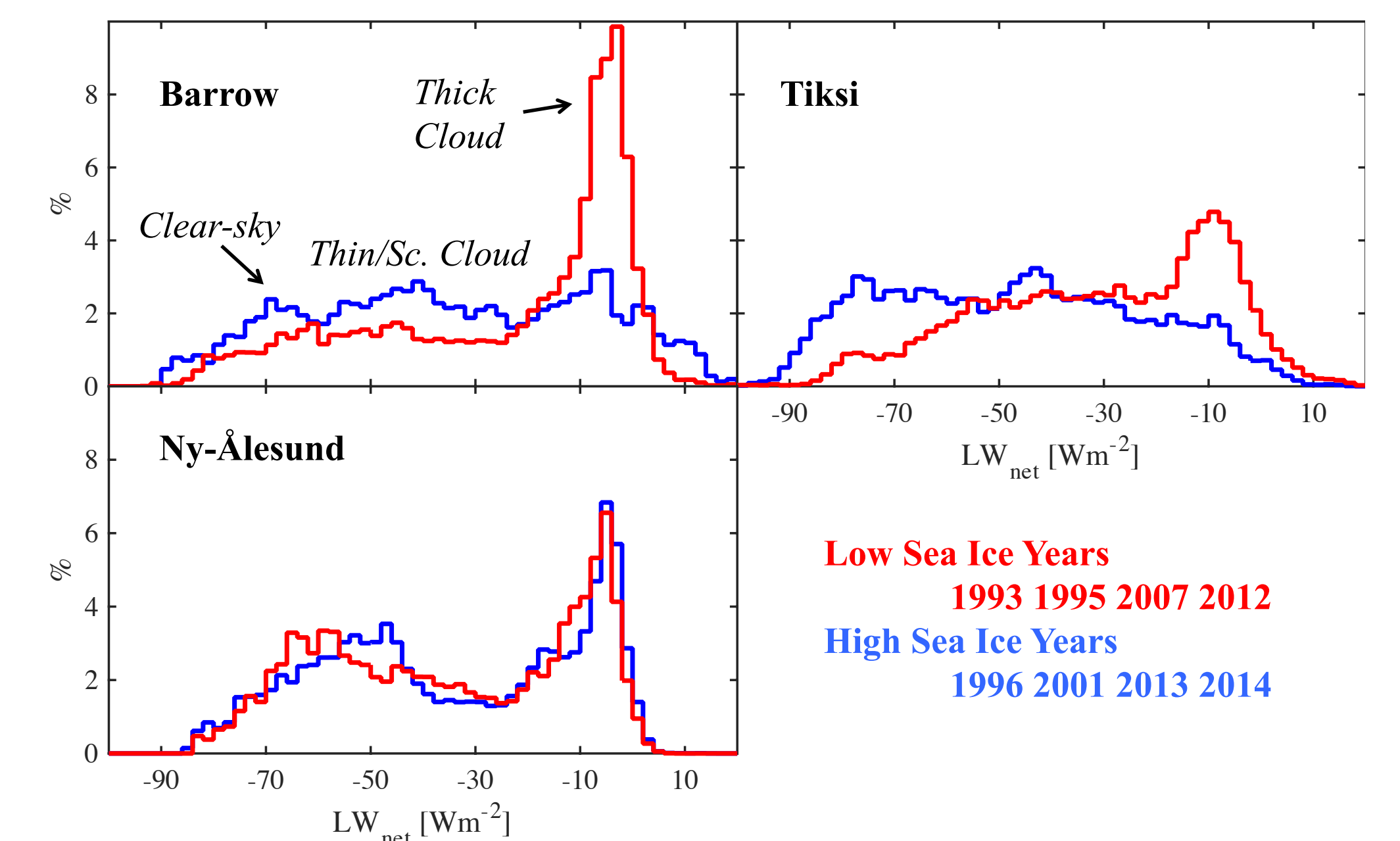


III. Towards a seasonal sea ice forecast

- Based on (II), observations of fluxes from stations may be suitable for seasonal sea ice forecasting of the pan-Arctic and sub-regions (a,c,d,e,f)
- Detrended anomalies similar to (a), but uses Barrow only. Shows consistency in reconstruction of all years developed from subsets of Barrow data record



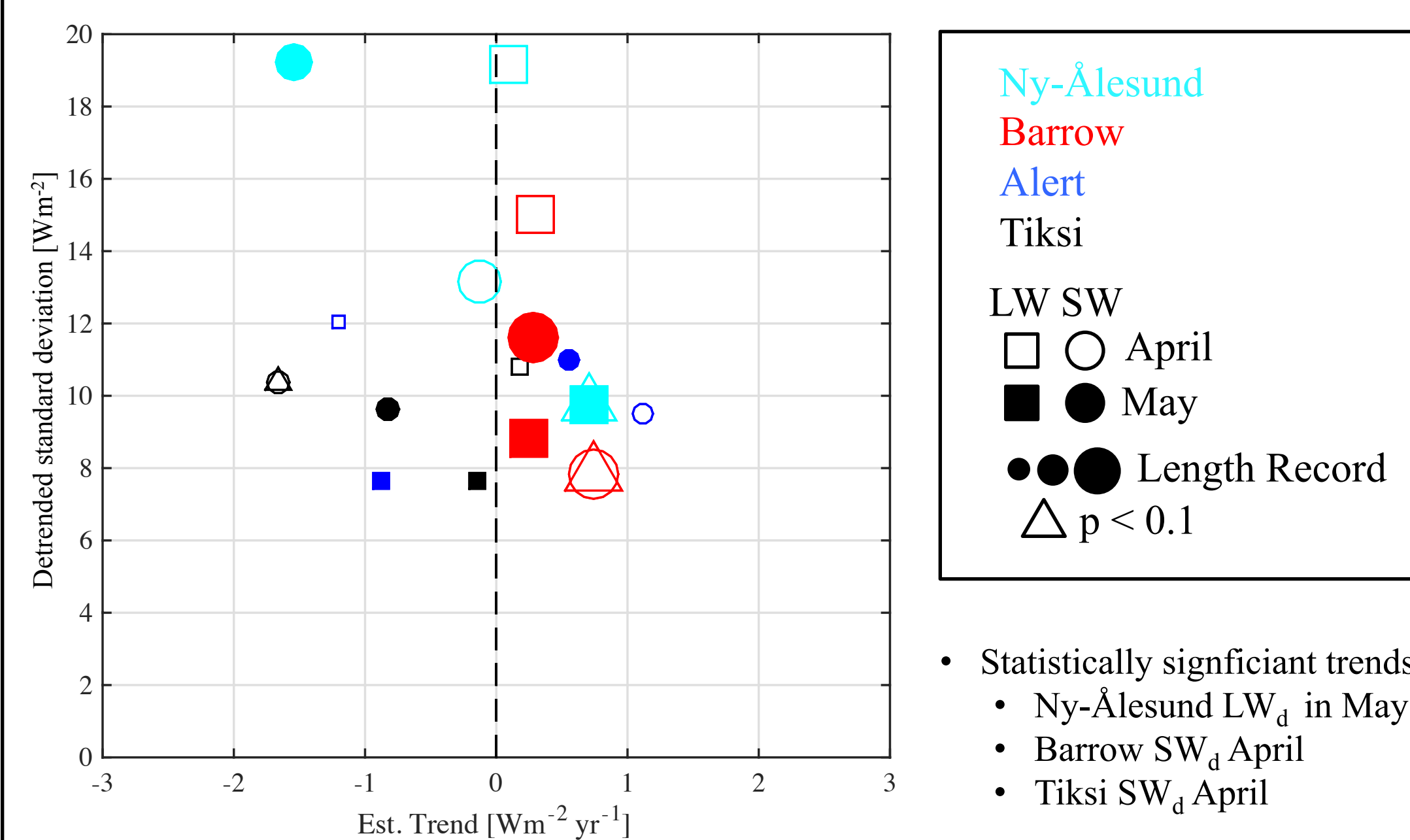
VI. LW_{net} mid to late April



- Bimodal distributions in LW_{net} (clear and cloudy modes) common in the Arctic
- April cloud anomalies in low sea ice years reduce surface cooling
- Consistent with (II), signal is seen at Barrow and Tiksi, but not Ny-Ålesund

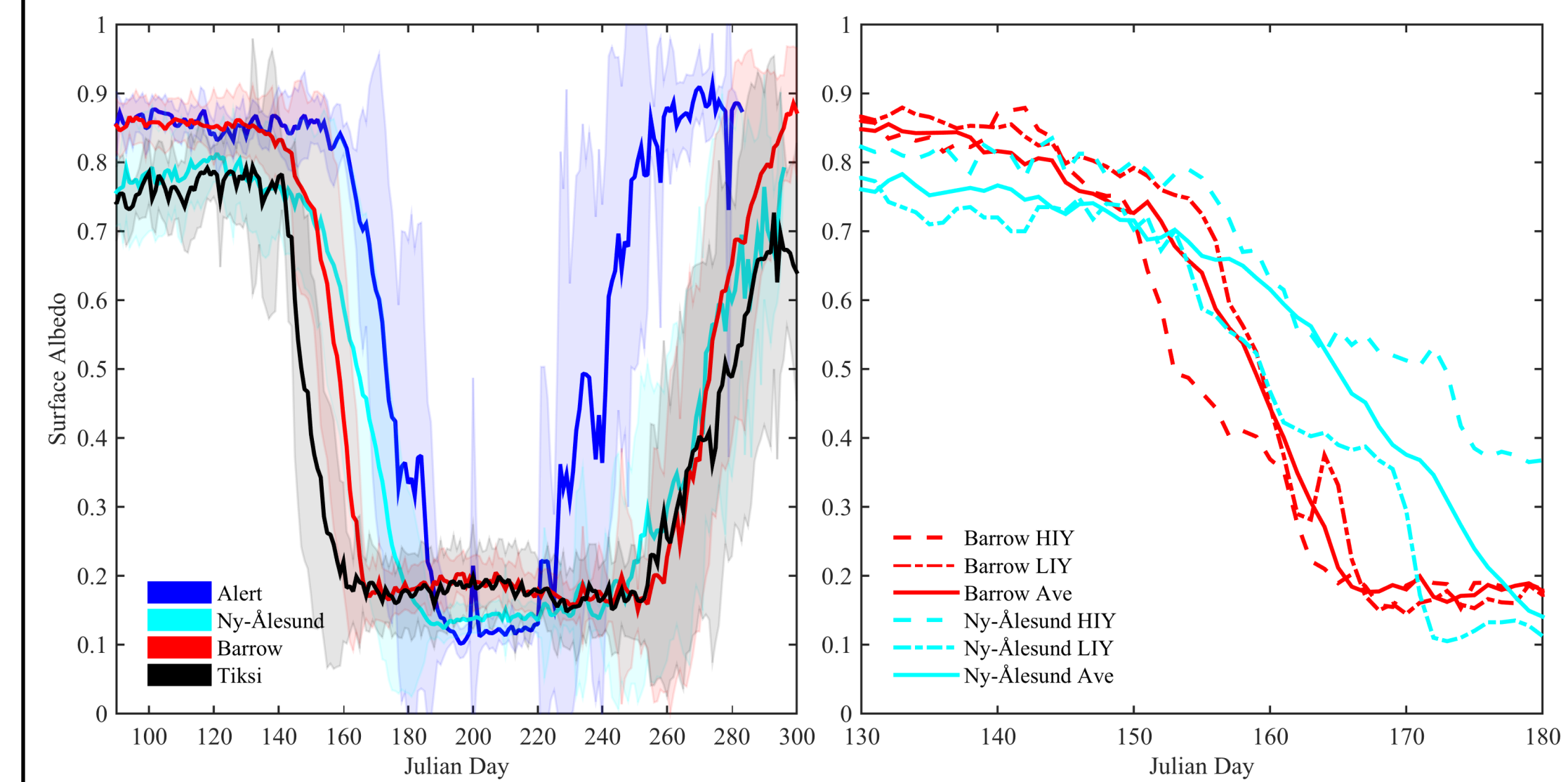
IV. April & May SW_d LW_d Trends

- Interannual variability is generally large compared to trends



- Statistically significant trends:
 - Ny-Ålesund LW_d in May
 - Barrow SW_d April
 - Tiksi SW_d April

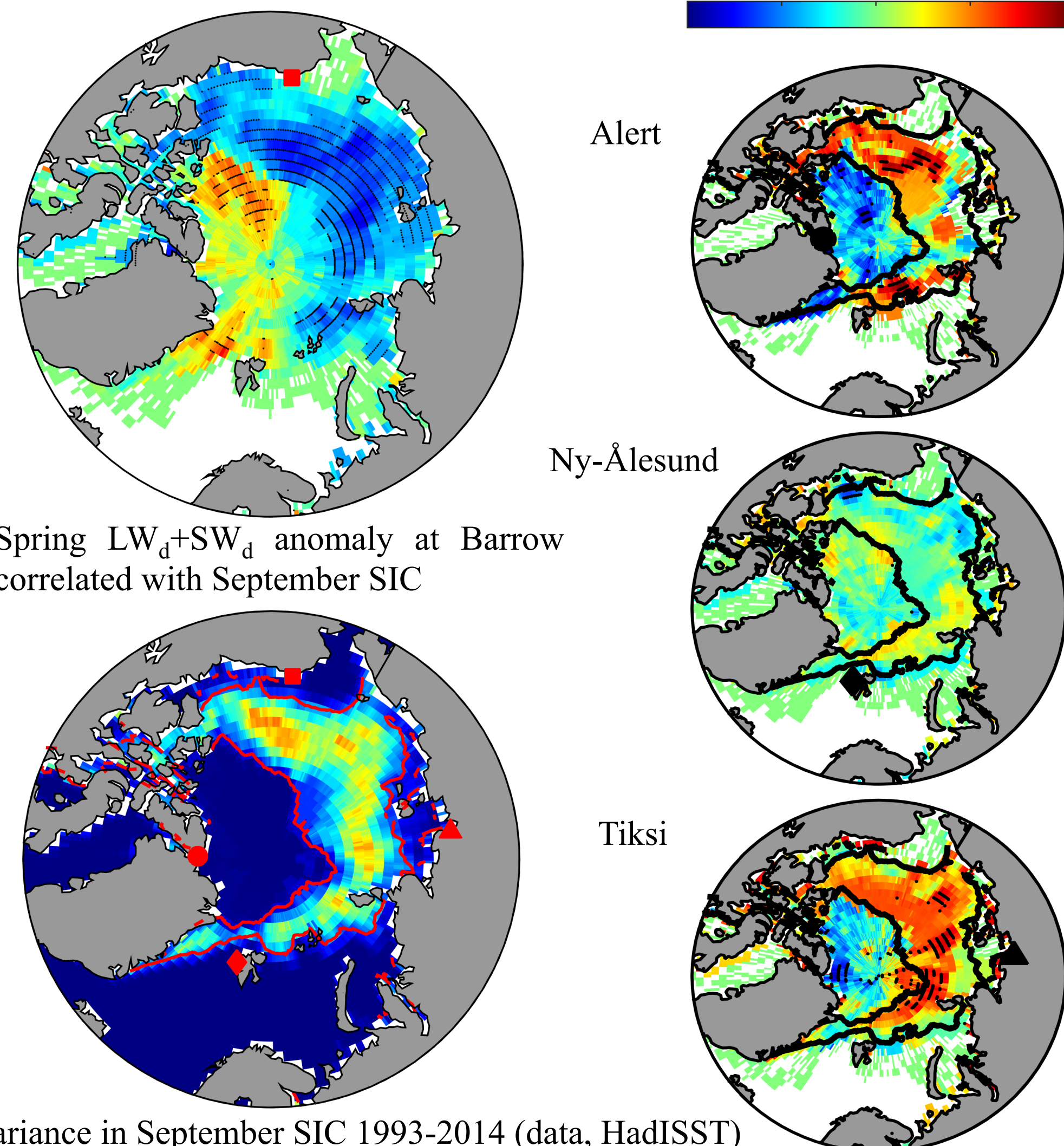
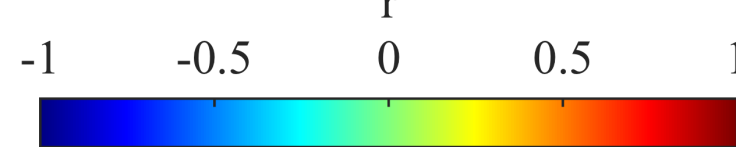
VII. Snowmelt Dates



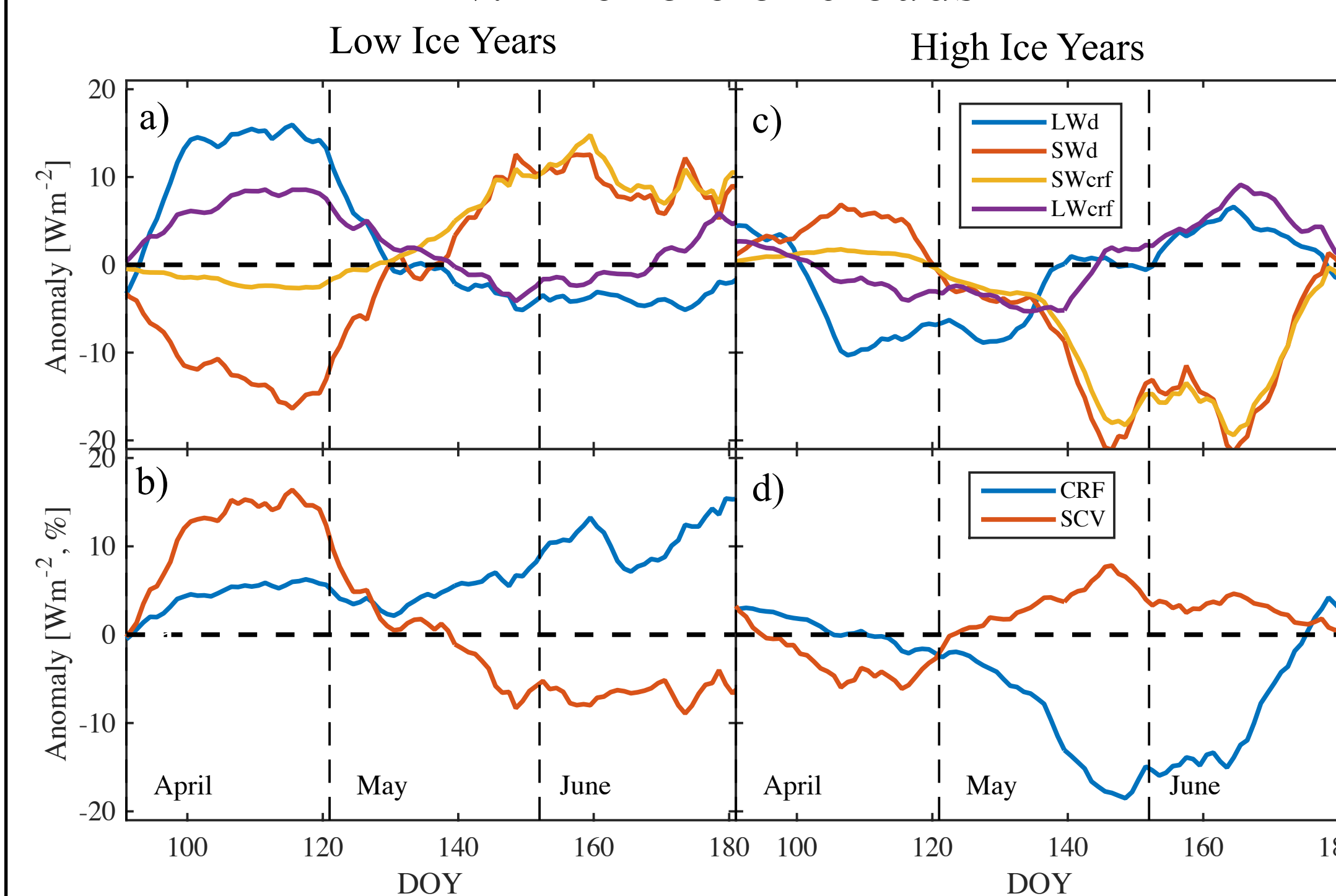
- Different locations have different snow-in, snow-out dates
- Snow melt in low ice years (LIY) is later at Barrow and earlier at Ny-Ålesund
- Barrow: also later snow onset in autumn in LIY (not shown)

II. Spatial representation of observatories: Lagged-correlative links to sea ice

- April-May downwelling shortwave (SW_d) and downwelling longwave (LW_d) radiation anomalies correlated with September sea ice concentrations (SIC) in region of interannual variability in sea ice
- The signal of atmospheric preconditioning^[1-3]



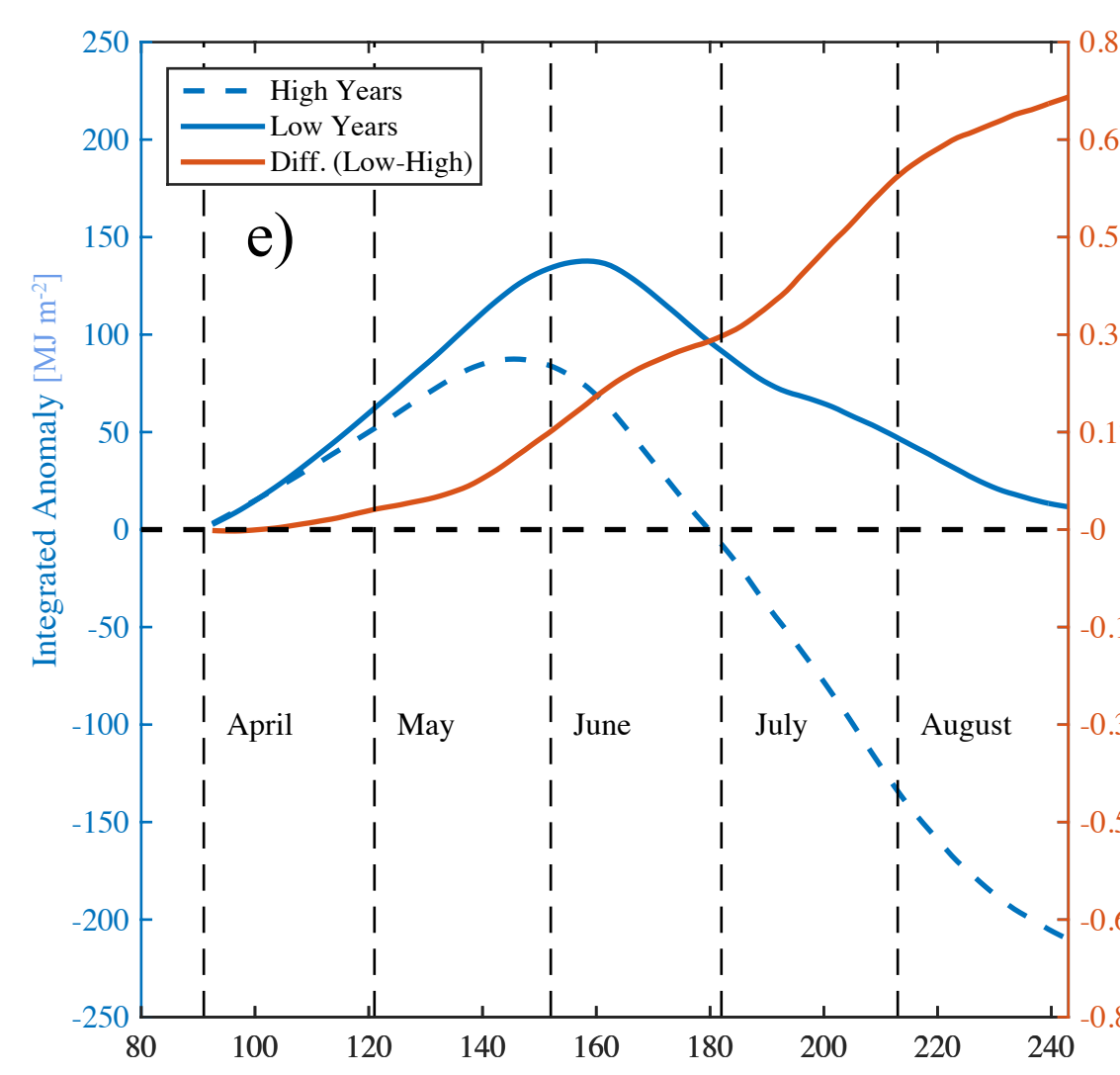
V. The role of clouds



- Cloudier early spring (April) followed clearer late spring (May-June) in low sea ice years (a,b). The opposite is true in high sea ice years (c,d), though the signal is weaker.

- In low years (a,b), cloud radiative forcing maximized: increases LW warming before seasonal cycle in SW dominates and minimizes SW cooling after.

- (e) From April 1, clouds support additional melted ice in low ice years vs. high ice years of 0.16 m (by June 1) and 0.74 m (by Sept 1).



VIII. Conclusions

- Radiative fluxes observed at coastal Arctic land stations carry information relevant to the sea ice zone where climate monitoring from the surface is impractical.
- Anomalies in SW_d and LW_d radiative fluxes in spring capture atmospheric preconditioning of sea ice – a seasonal scale forecast is possible.
- Trends in the fluxes in spring are found, but are small compared to interannual variability.
- Cloud radiative forcing is transitioning in spring from dominant LW warming (winter) to SW cooling (summer)^[4]. Clouds enhance melt season when positive cloud anomalies appear in April followed by negative anomalies in May/June. This maximizes longwave warming and minimizes shortwave cooling at critical times.
- The cloud anomaly in April substantially reduces LW cooling of the surface in early spring during low sea ice years.
- The spring snowmelt season is also influenced by these anomalies, though in more complex ways.

References & Acknowledgements

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